

Tenerife, March 2018

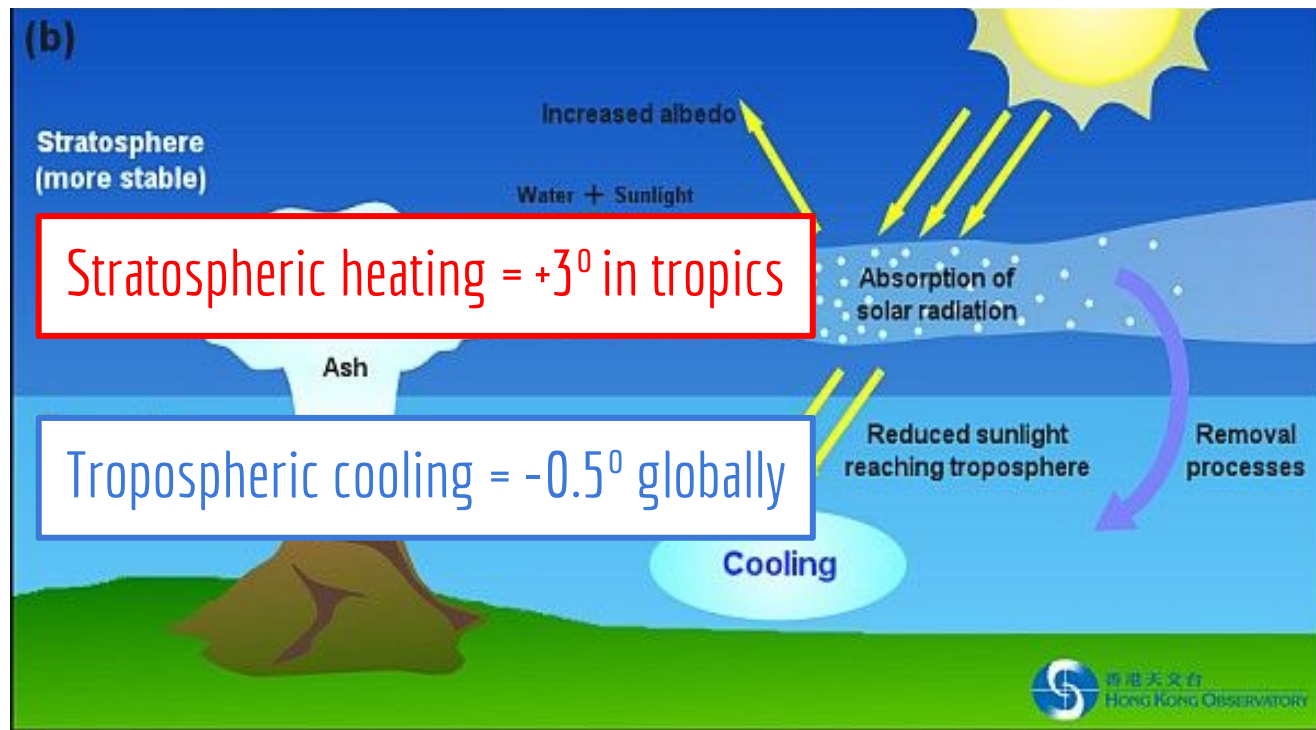
Stratospheric heating and surface cooling caused by volcanic aerosols in CM2.1 model

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Suleiman Mostamandy, Alexander Ukhov
and Georgiy L Stenchikov



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King Abdullah University of
Science and Technology

Volcanic Impact on Climate



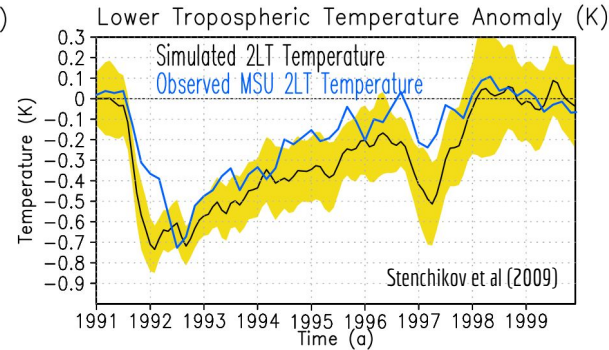
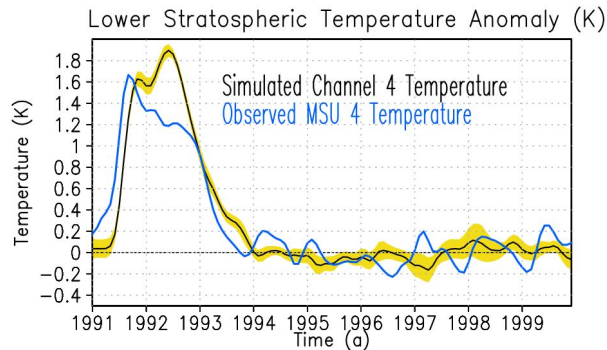
Volcanic aerosols in models

Ways to account for the volcanic aerosols in models:

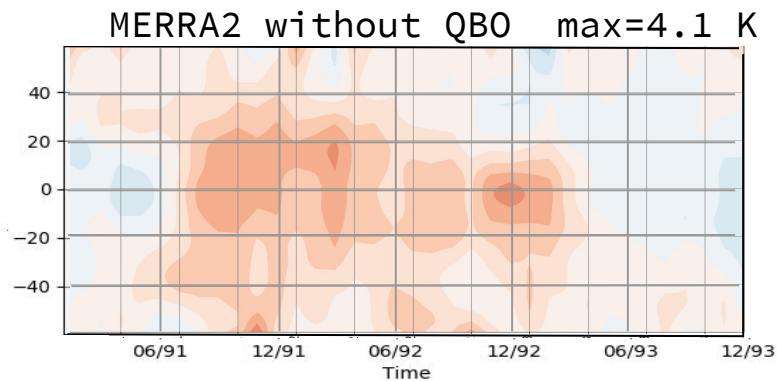
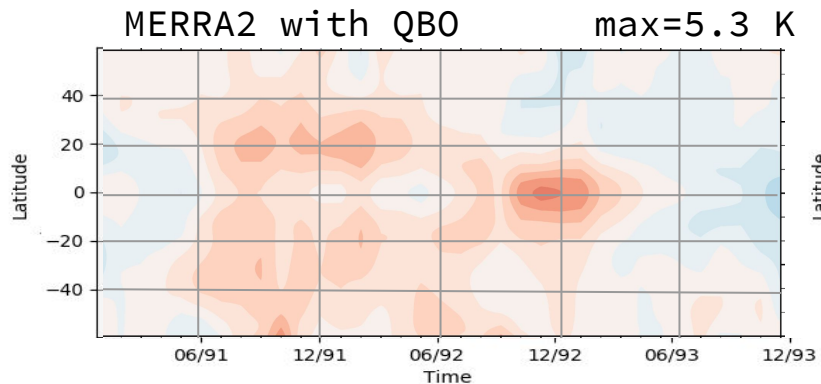
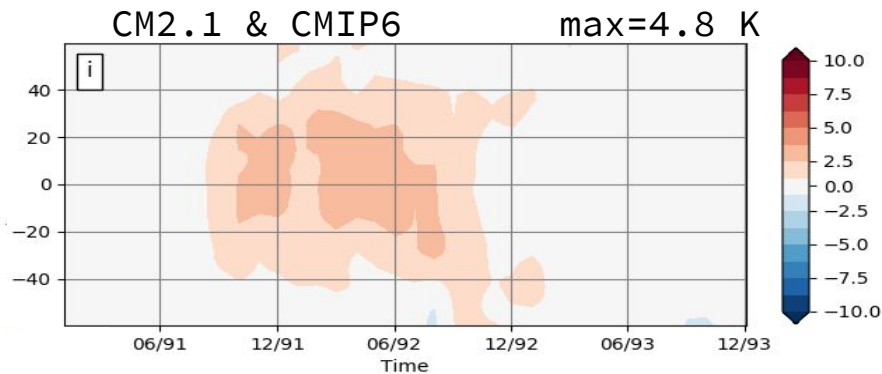
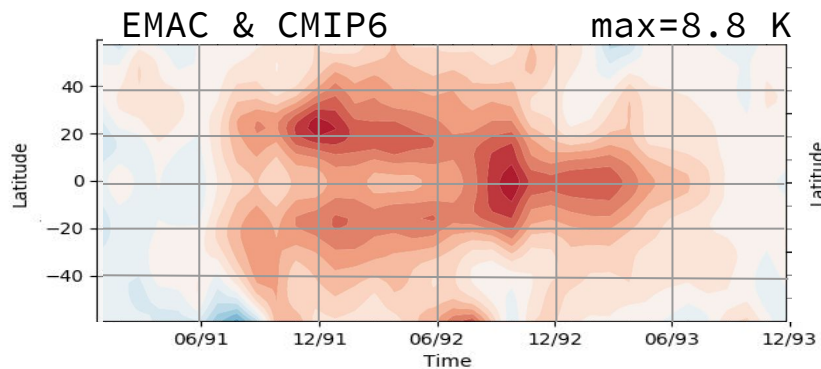
Interactive simulation
of the plume

Prescribed aerosol characteristics:

- aerosol extinction
- single scattering albedo
- asymmetry parameter



50 hPa temperature anomaly, Pinatubo 1991



Radiative impact

Radiative impact is the changes in:

1. **TOA radiative fluxes**
[the energy balance change of the entire system]
2. **Atmospheric heating rates**
[stratospheric heating]
3. **Surface fluxes**
[surface cooling]

Sensitive to:

Radiative scheme

Model dynamics

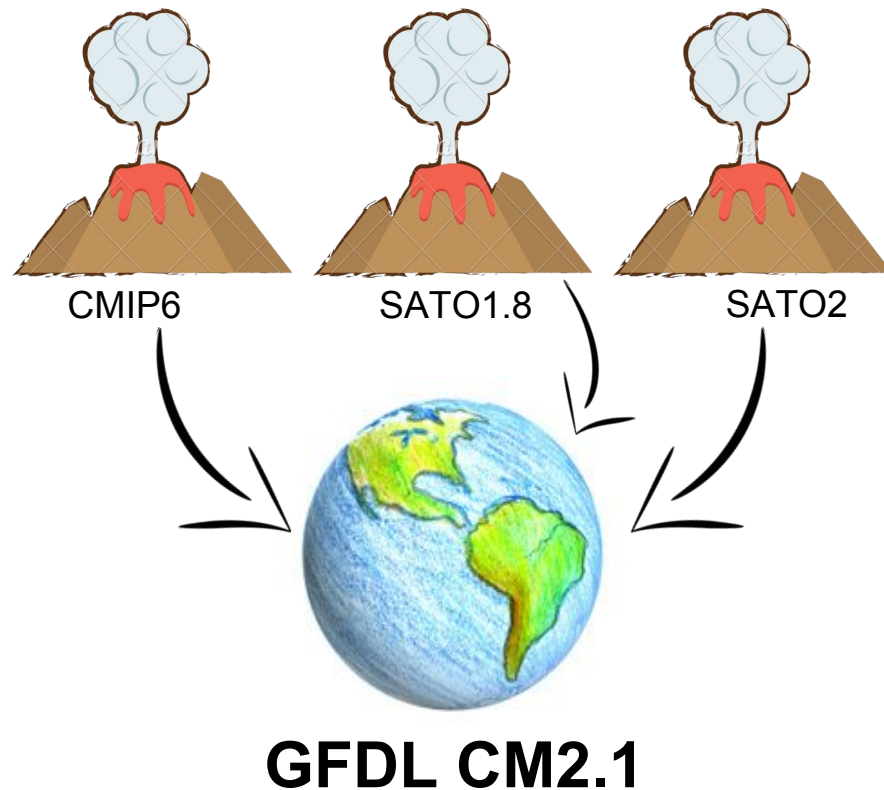
Aerosol optical
properties

Objective

**Evaluate
different volcanic datasets**

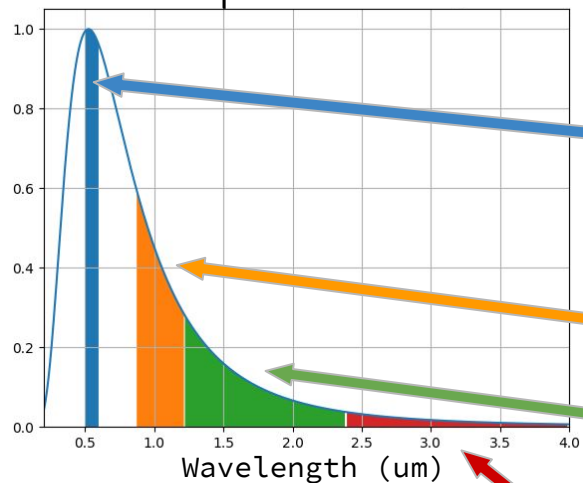
Questions:

1. What are the most important characteristics of a dataset that define the radiative impact?
2. How does the model reproduce the observed stratospheric heating and tropospheric cooling?



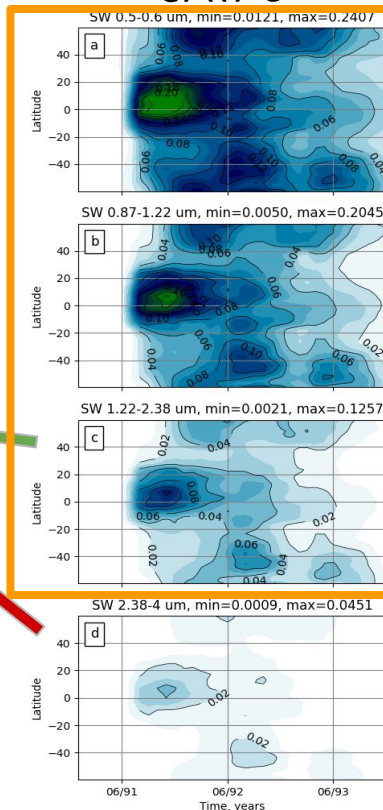
Extinction optical depth

Solar spectral irradiance

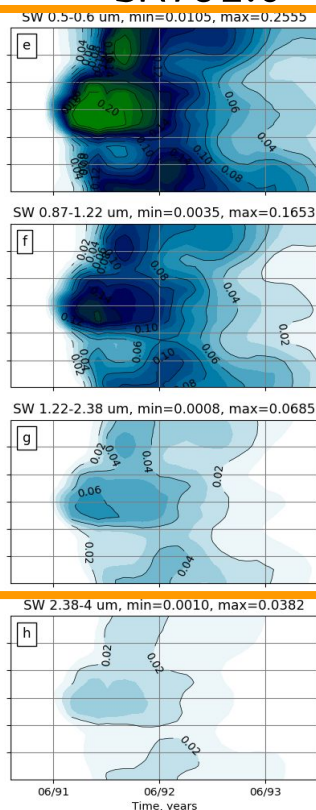


EAOD defines how much of solar radiation was prevented from reaching the surface by scattering and absorption.

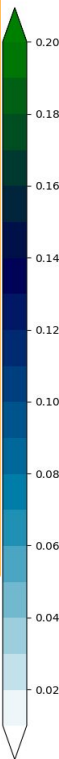
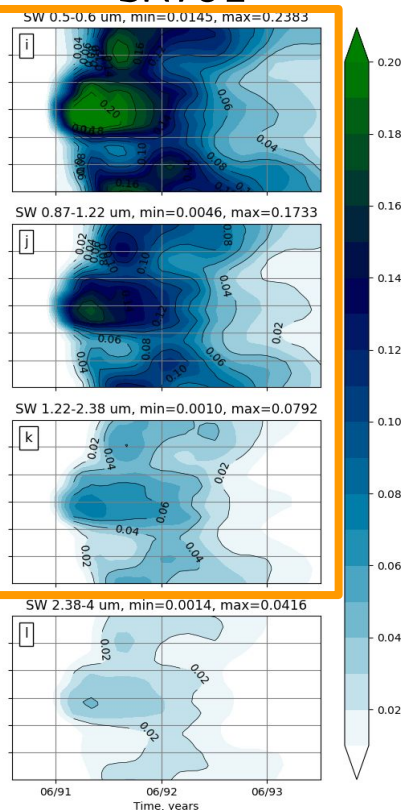
CMIP6



SAT01.8

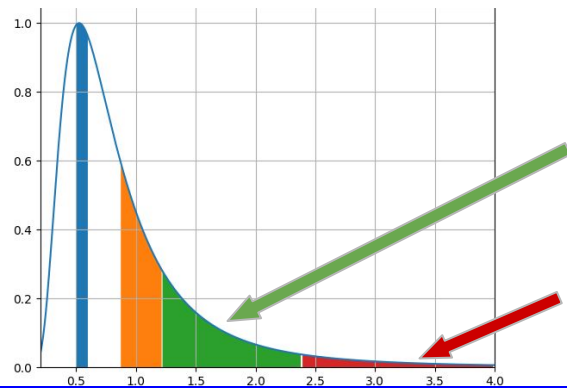


SAT02

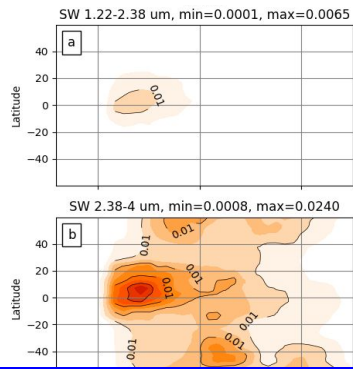


Absorption optical depth

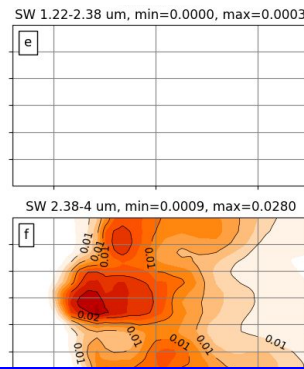
Solar spectral irradiance



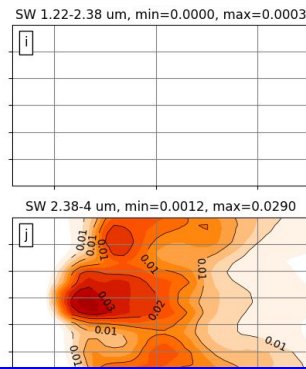
CMIP6



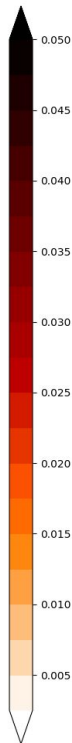
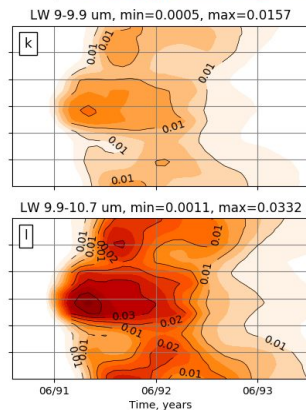
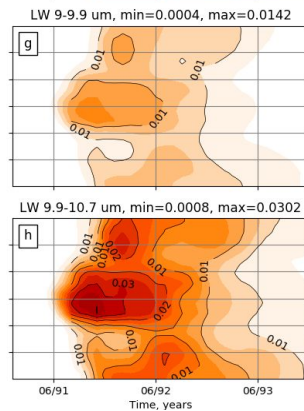
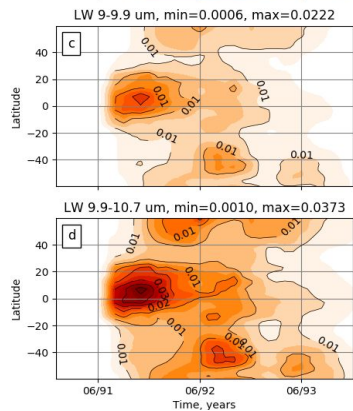
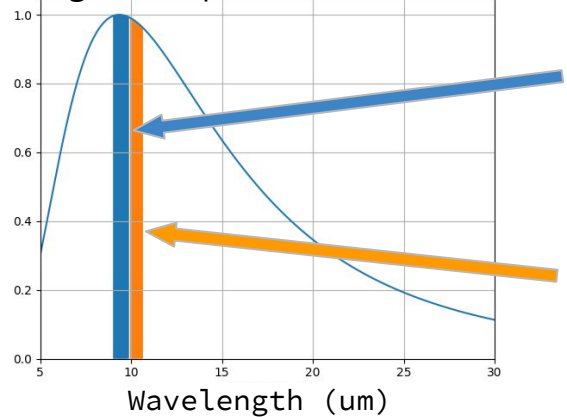
SAT01.8



SAT02

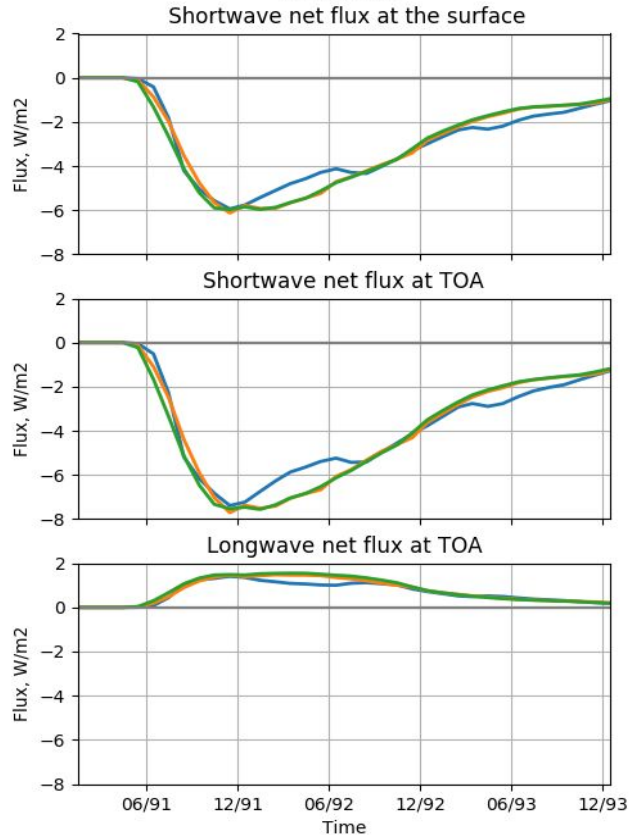


Longwave spectral irradiance

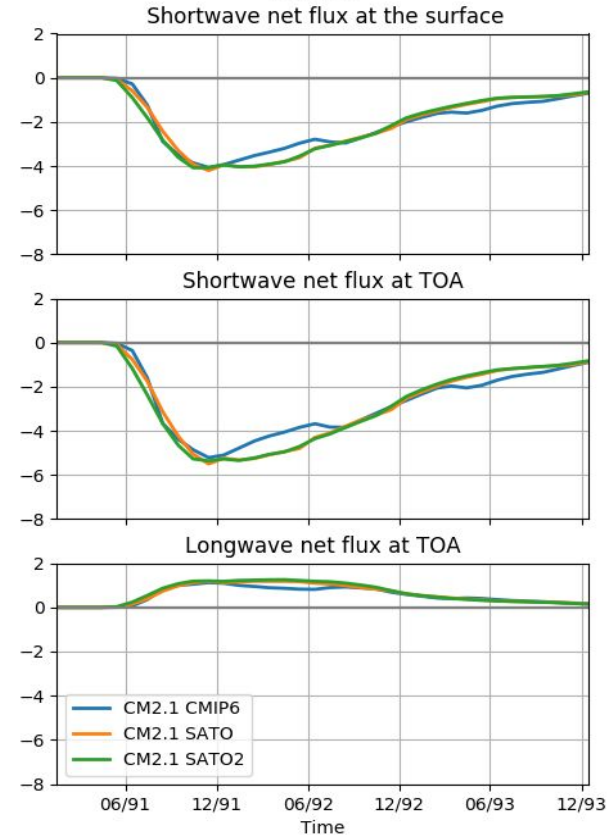


Radiative forcing at TOA and BOA

CLEAR SKY

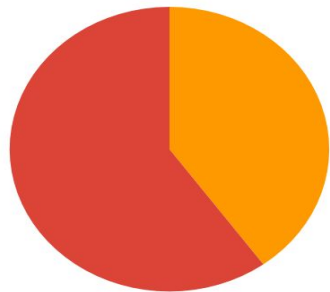


ALL SKY

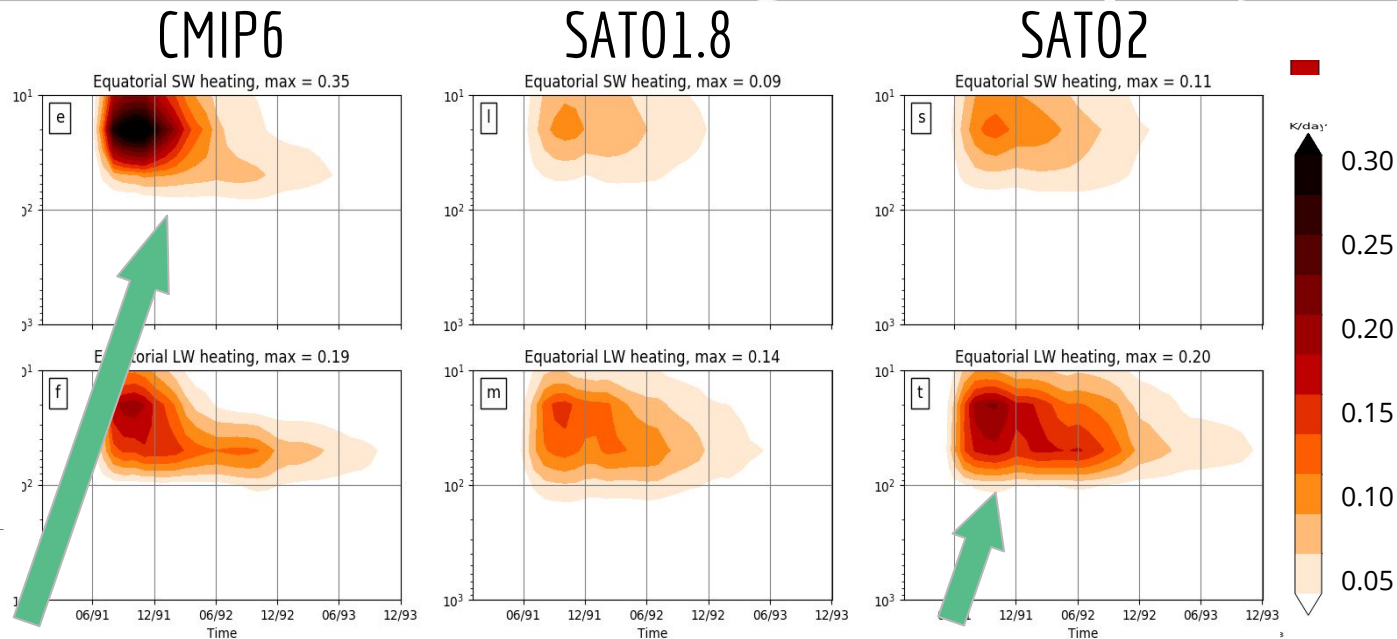


Equatorial SW and LW heating rates (K/day)

Contribution of
● Shortwave ● Longwave



Stenchikov et.al 1998



CMIP6:

1. Overestimates SW absorption
2. SW heating rate is 3x larger than in SAT0 or SAT02.

SAT02:

1. Due to the larger number of coarse particles, LW heating is enlarged

Temperature response

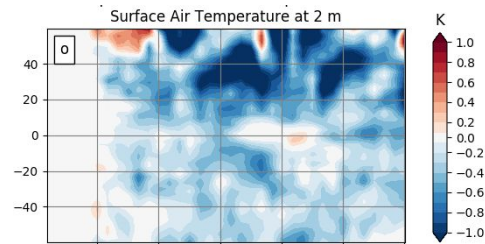
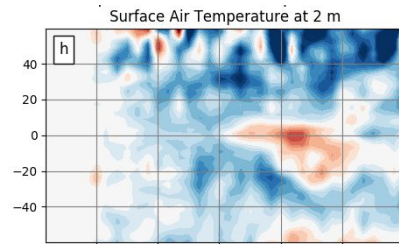
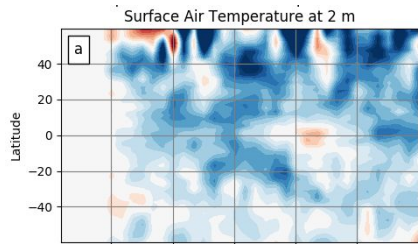
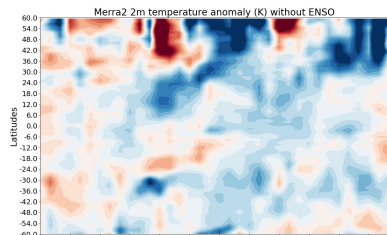
MERRA2

CMIP6

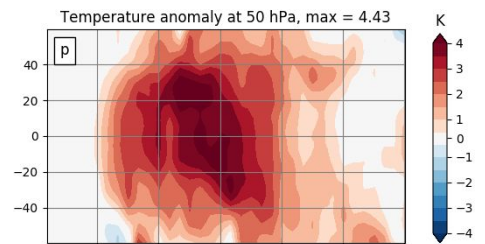
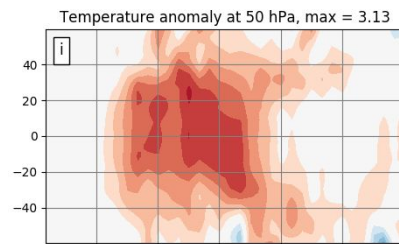
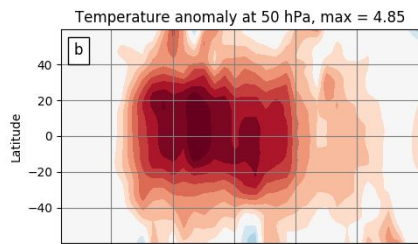
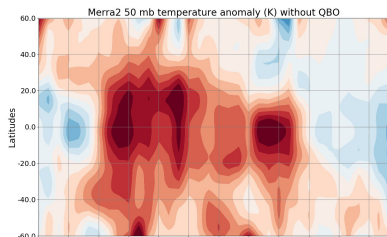
SAT01.8

SAT02

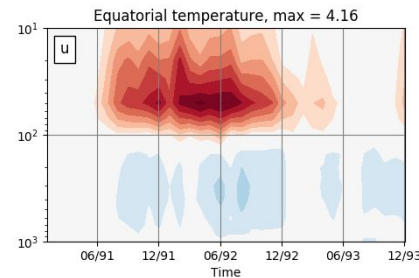
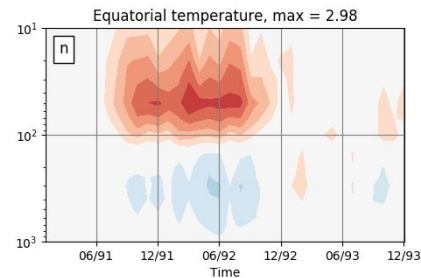
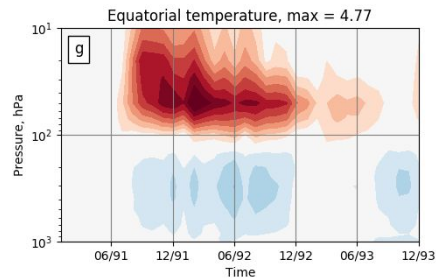
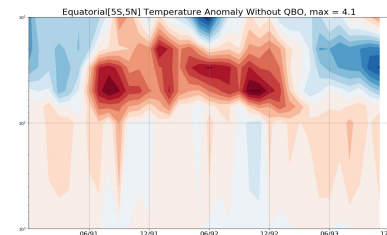
Surface



50 hPa



Equatorial

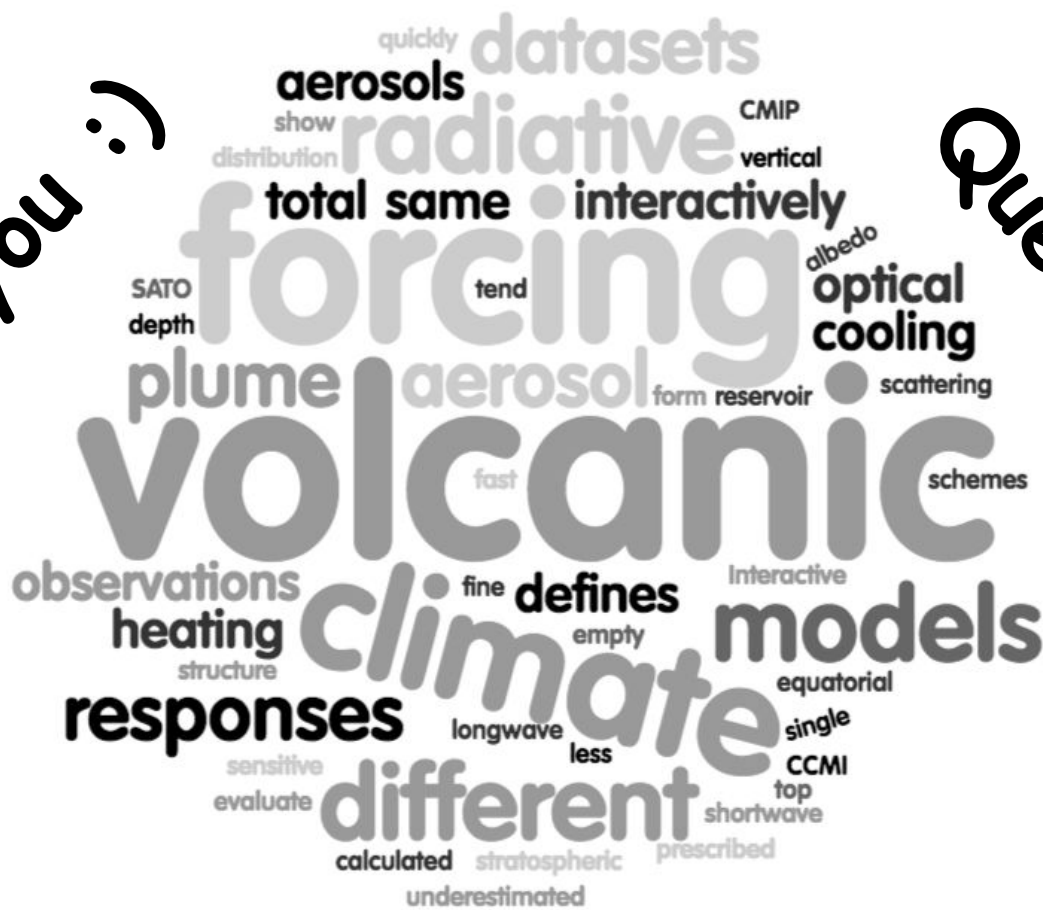


Conclusions

1. All datasets produce comparable forcing and climate impacts
2. Extinction and absorption optical depth are the most important aerosol characteristics to be preserved during processing
3. Stratospheric heating response is sensitive to SW and LW aerosol absorption. LW absorption is more important for stratospheric heating than the SW
4. CMIP6 overestimates solar heating rates as it absorbs starting from 1.22 μm
5. LW absorption is sensitive to the width of aerosol size distribution
6. Tropospheric cooling is sensitive to solely SW optical depth

Thank You :)

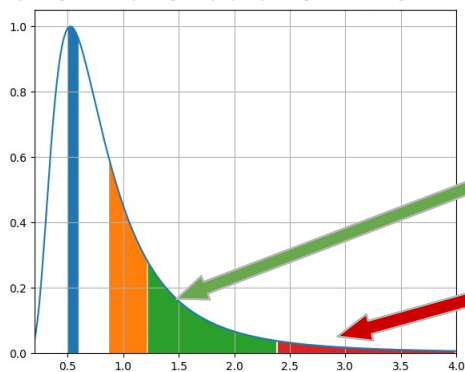
Questions?



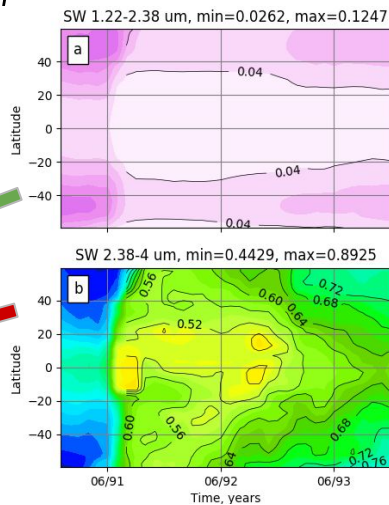
Evgeniya.Predybaylo@kaust.edu.sa

$A = 1 - \text{SSA (absorption)}$

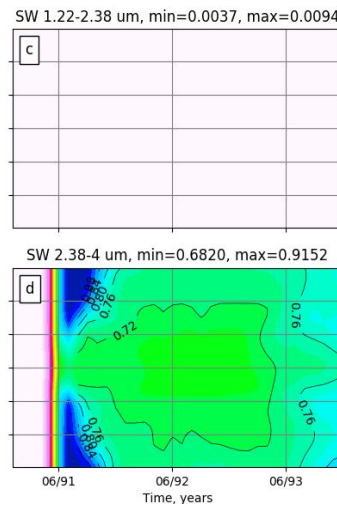
Shortwave radiation intensity



CMIP6



SAT01.8



SAT02

