



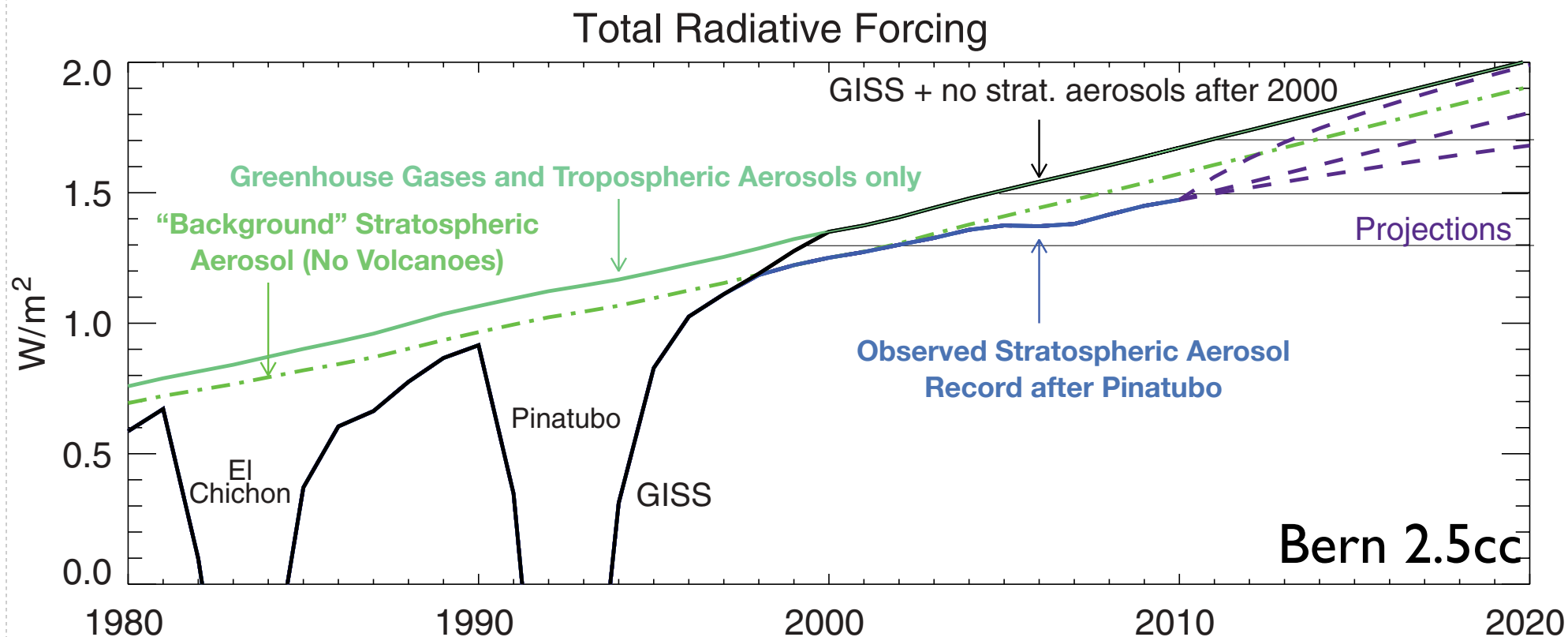
# The Contribution of Anthropogenic SO<sub>2</sub> Emissions to the Asian Tropopause Aerosol Layer

**Ryan R. Neely III (NCAR/ASP),** , P. Yu, K. H. Rosenlof, O. Brian Toon, Susan Solomon, J. M. English, J. S. Daniel, H. L. Miller

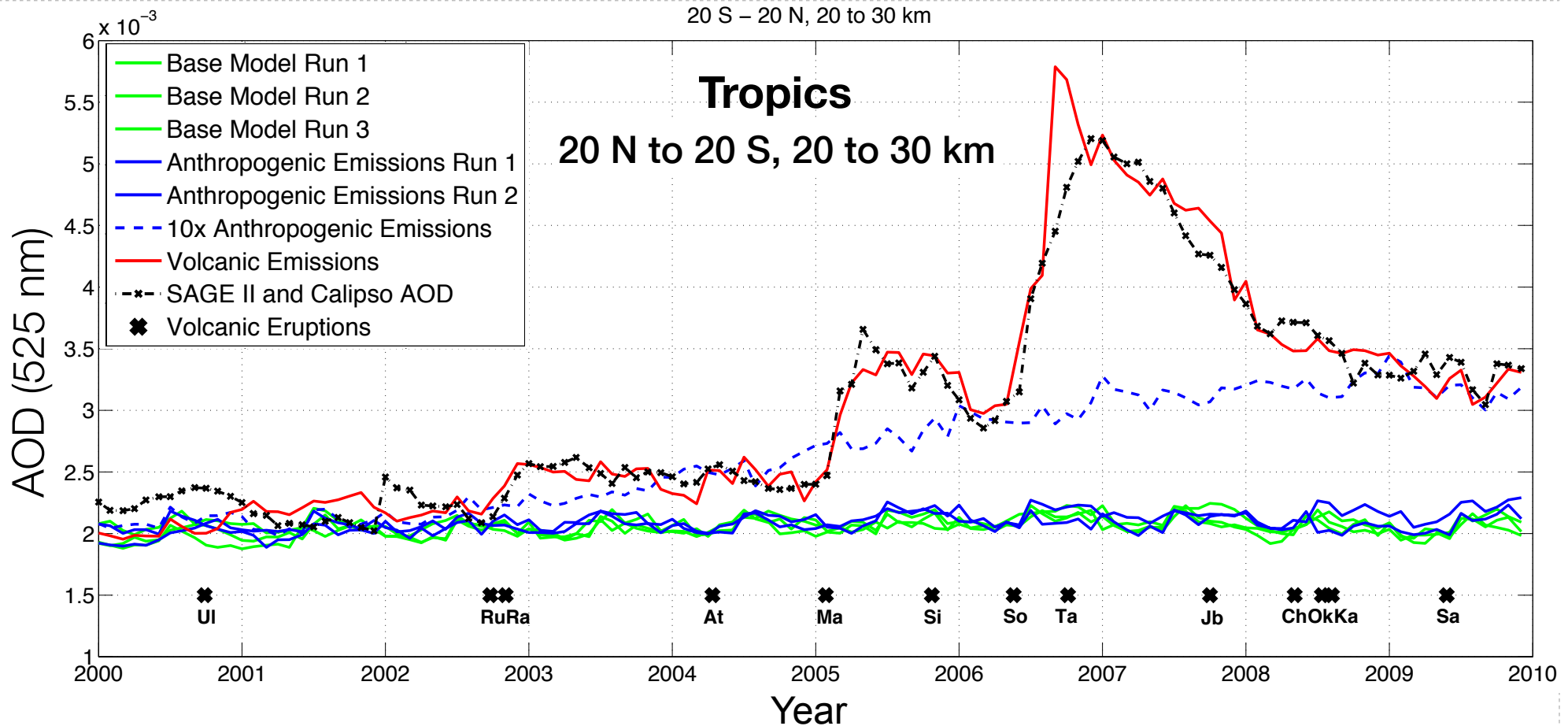
Adapted from: Neely et al., (2013), Regional Contributions of Anthropogenic SO<sub>2</sub> Emissions to the Asian Tropopause Aerosol Layer, Geophys. Res. Lett., *in review*.



# Variability in stratospheric aerosol impacts global radiative forcing

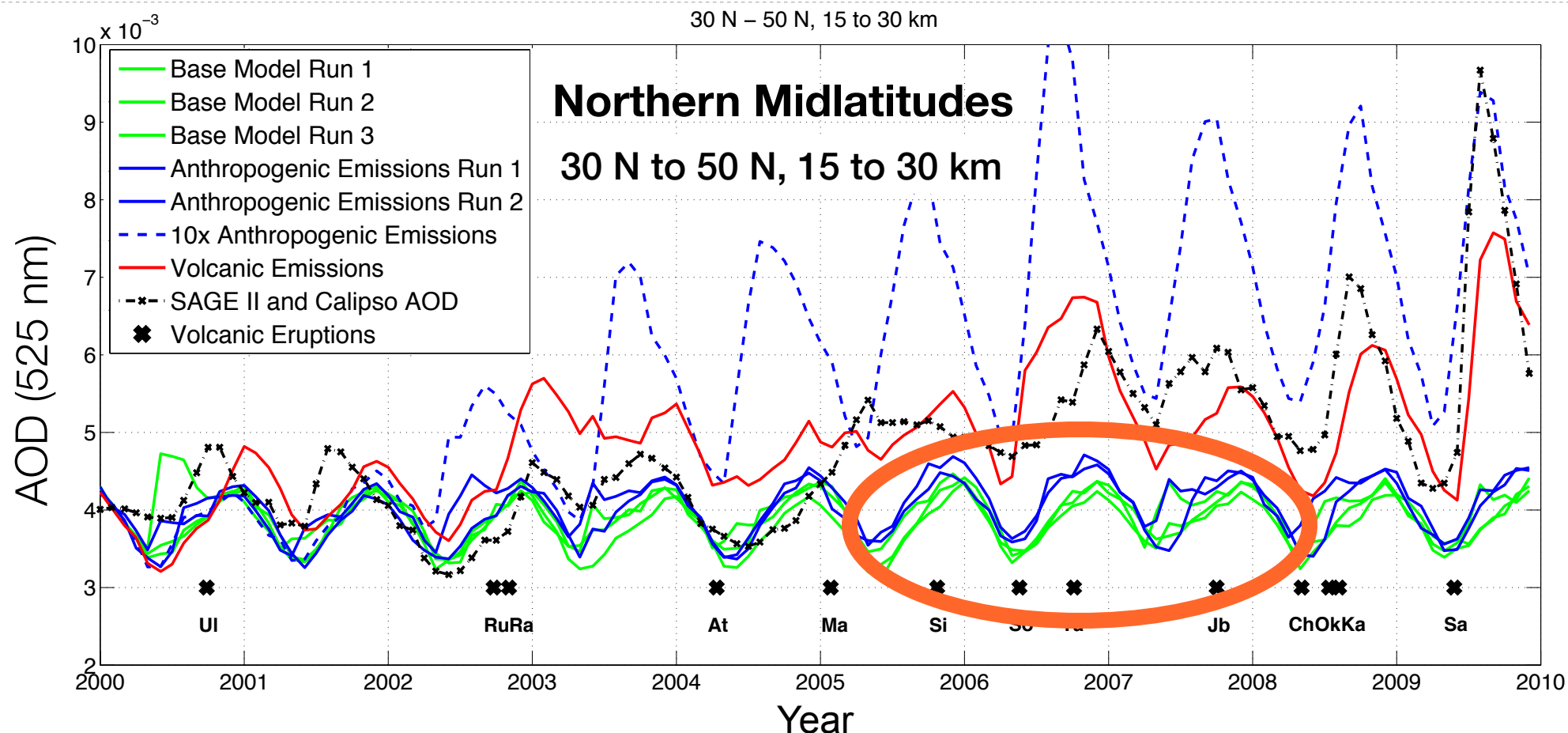


# Volcanoes drive stratospheric aerosol variability



Adapted from Neely, R. R., III et al. (2013), Recent anthropogenic increases in SO<sub>2</sub> from Asia have minimal impact on stratospheric aerosol, *Geophys. Res. Lett.*, n/a–n/a, doi:10.1002/grl.50263.

# Anthropogenic emissions may have some influence

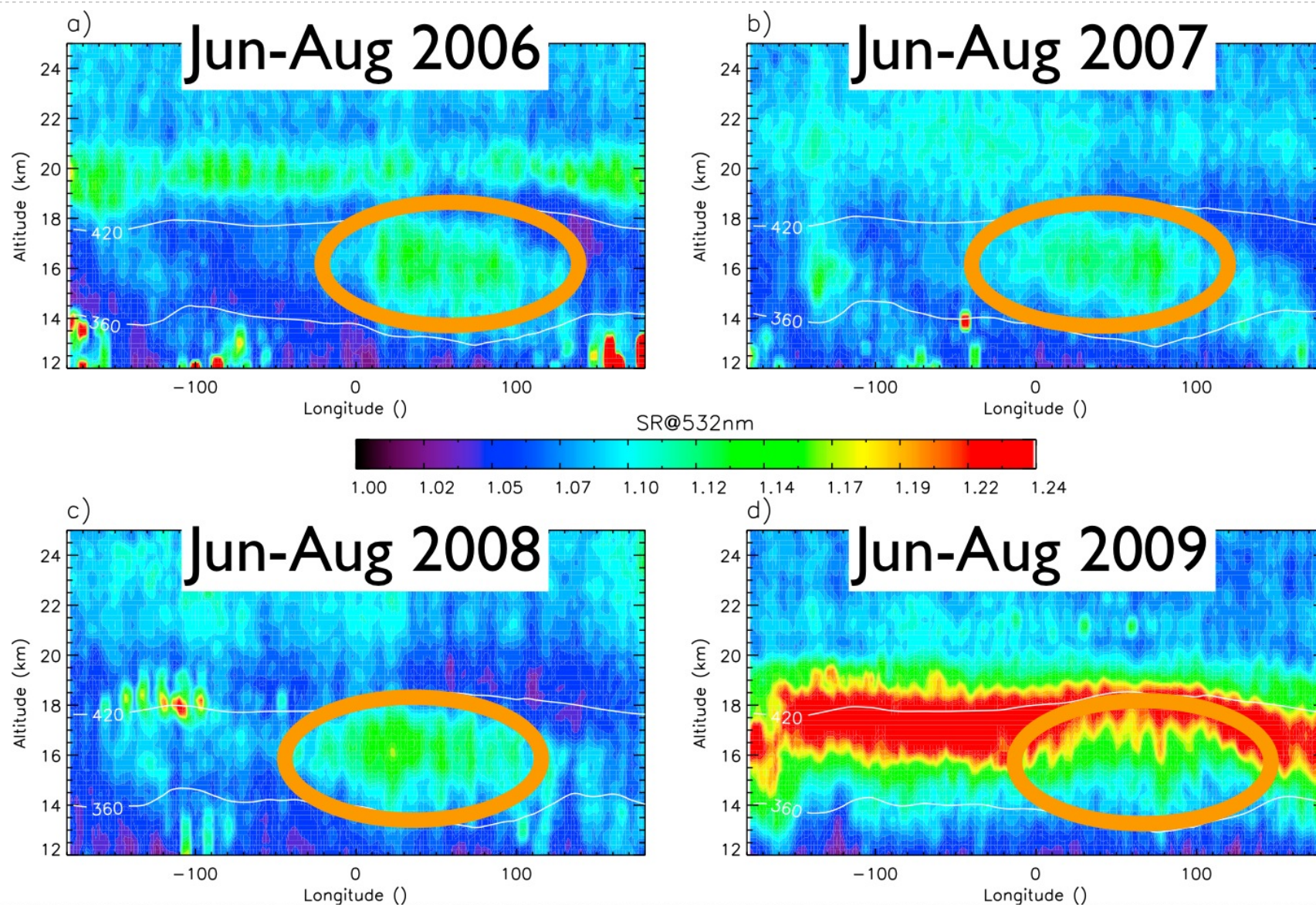


Recent variability in stratospheric aerosol is mostly due to volcanic eruptions but influences from anthropogenic emissions may have an impact on the lower troposphere (Neely et al. GRL 2013)



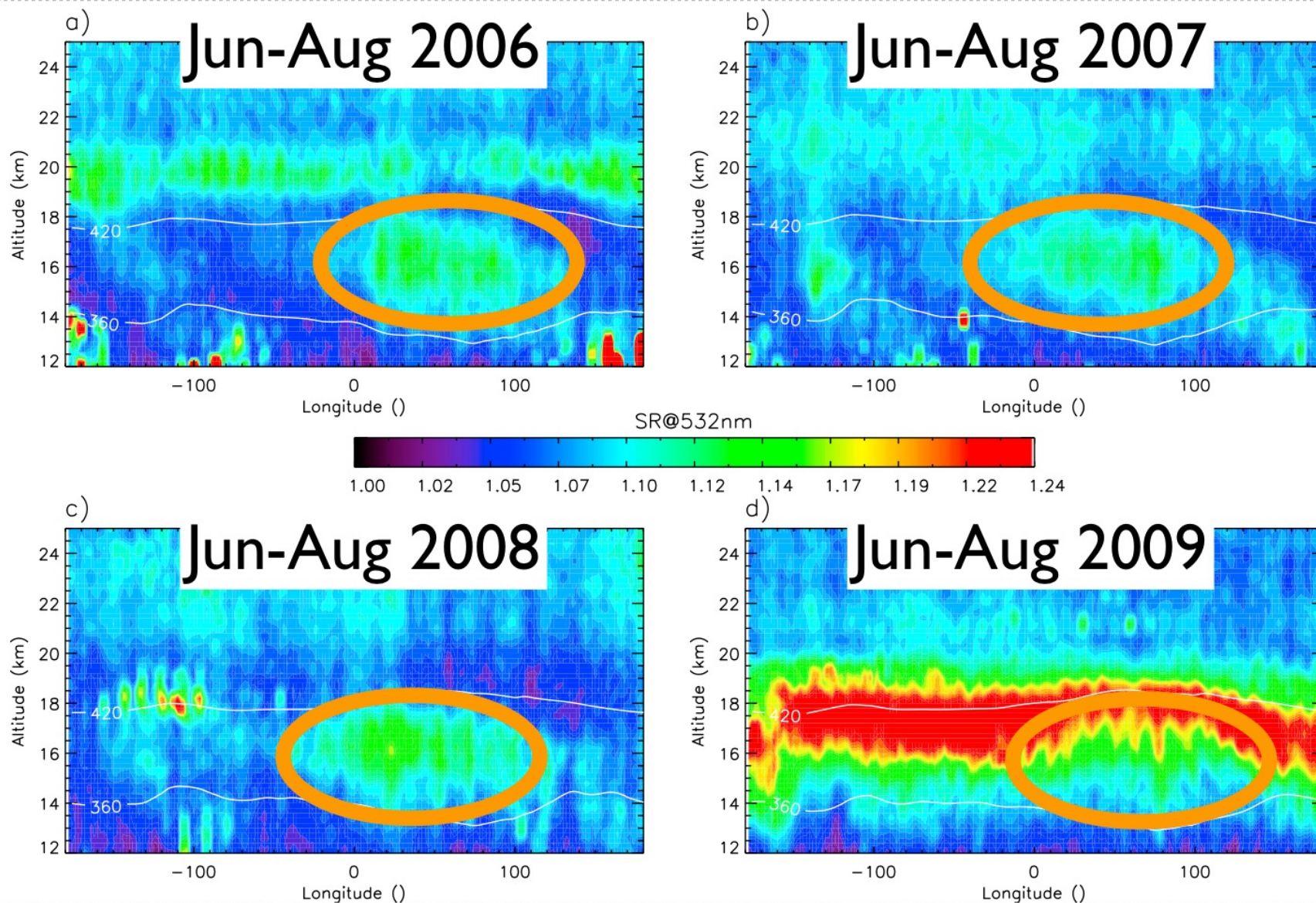
# The Asian Tropopause Aerosol Layer (ATAL)

## Mean Scattering Ratio (SR) from CALIPSO at 532 nm from 15N to 45N





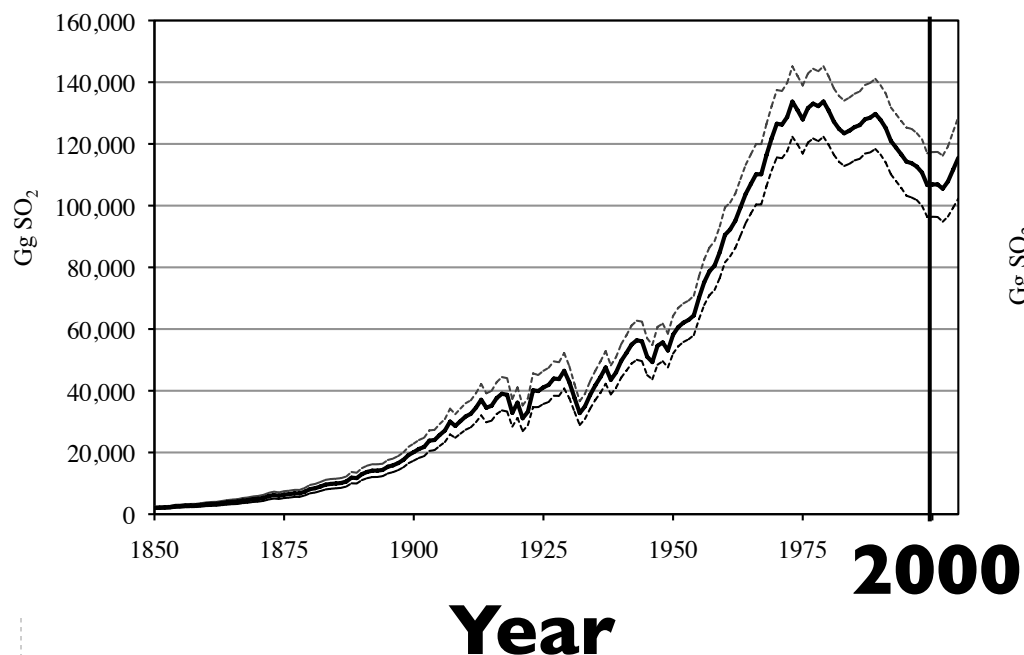
# Current observations cannot attribute the observed variability to sources



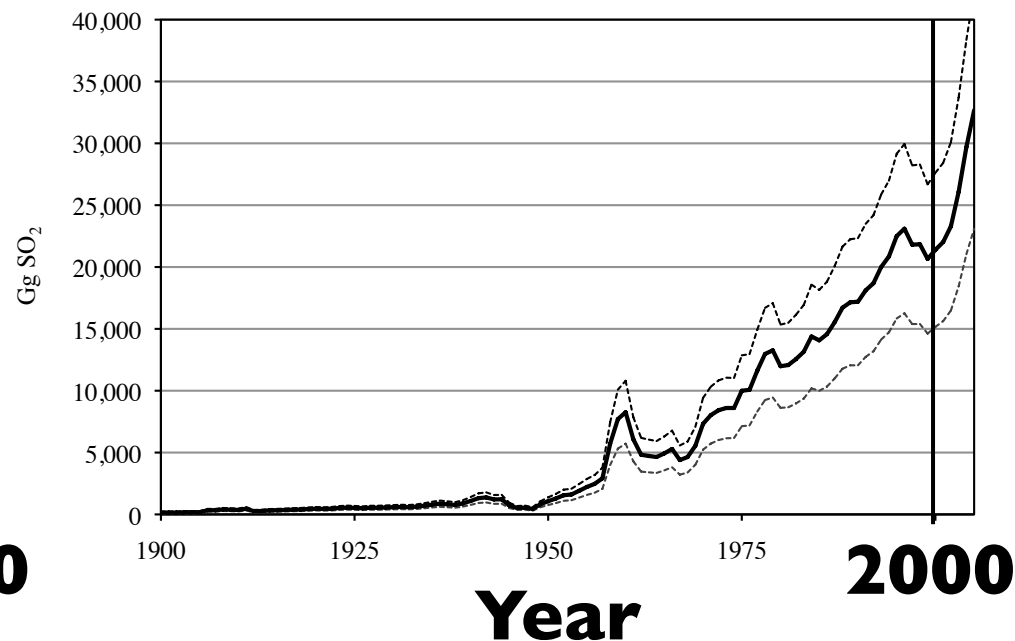


# Hypothesis: Asian Emissions Lead to ATAL

## Global $\text{SO}_2$



## Chinese $\text{SO}_2$

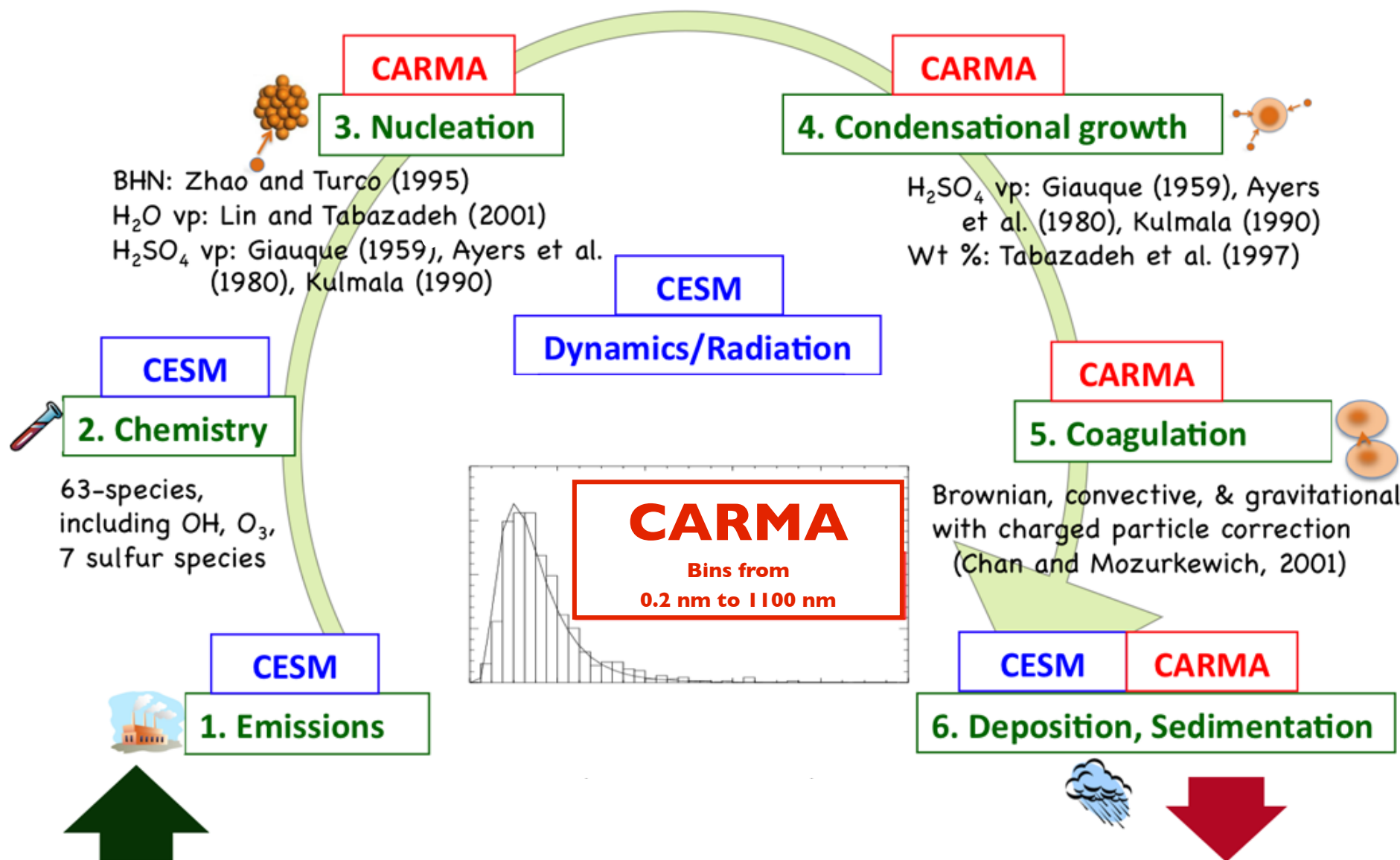


**Increases in other constituents of carbonaceous aerosol (POA, SOA) are also thought to contribute**



# The Model:

**Non-Sulfates:** Primary Organics (POA), Secondary Organics (SOA), Black Carbon (BC), Wind Blown Dust, Sea Salt  
**Sulfates:** Meteoritic Smoke, Pure Sulfates, Mixed Sulfates

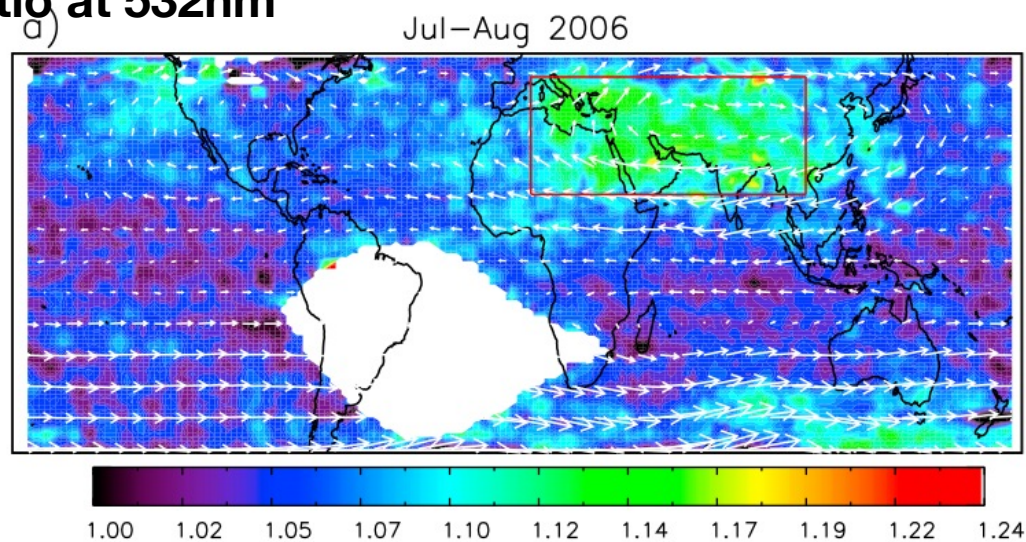
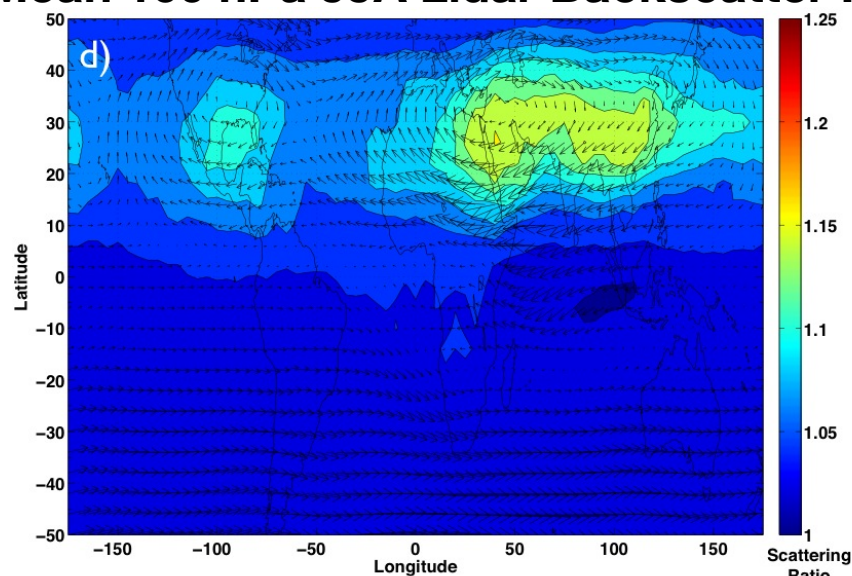




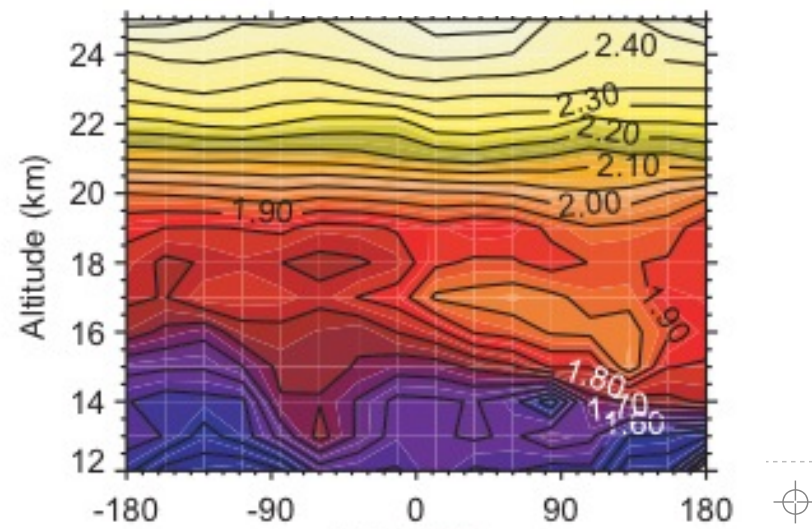
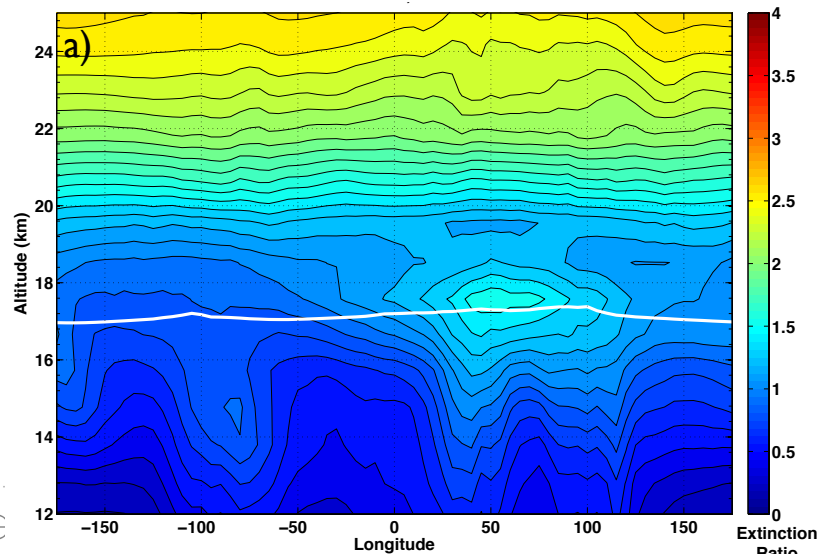


# Model Captures Spatial Pattern and Magnitude of Observations

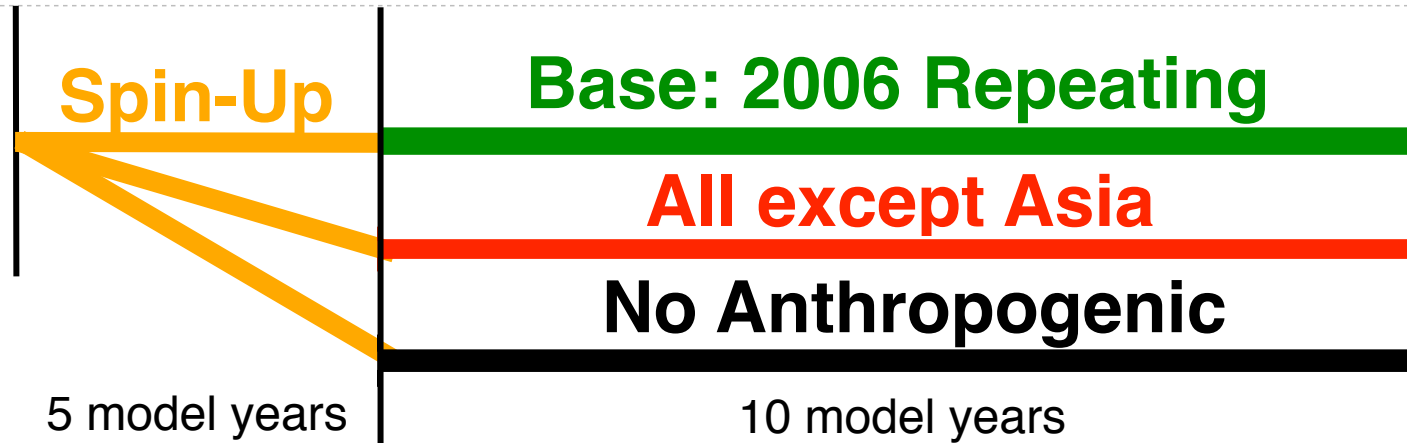
## Mean 100 hPa JJA Lidar Backscatter Ratio at 532nm



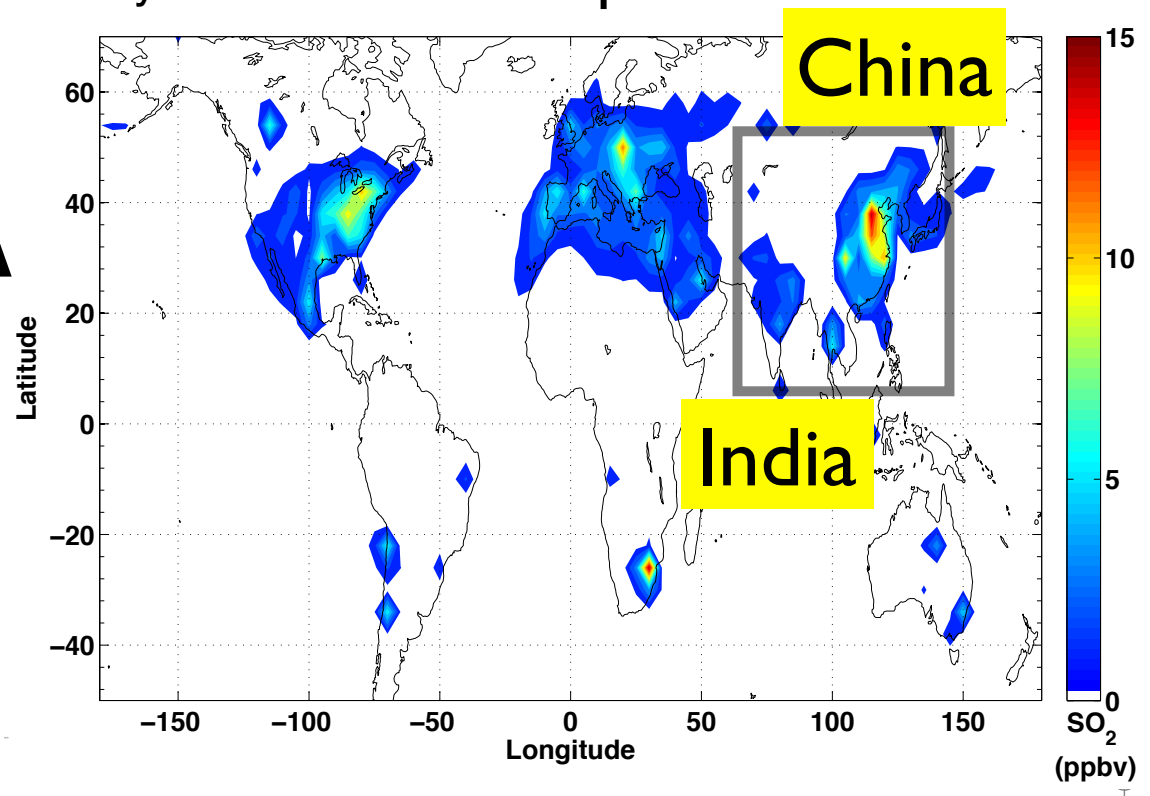
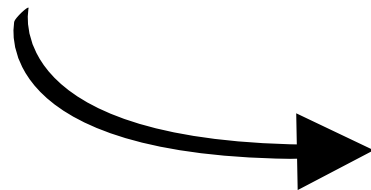
## Mean 1020 nm Extinction Ratio from 14N to 46N, June-August **SAGE II**



# Model Experiment: SO<sub>2</sub> Schemes



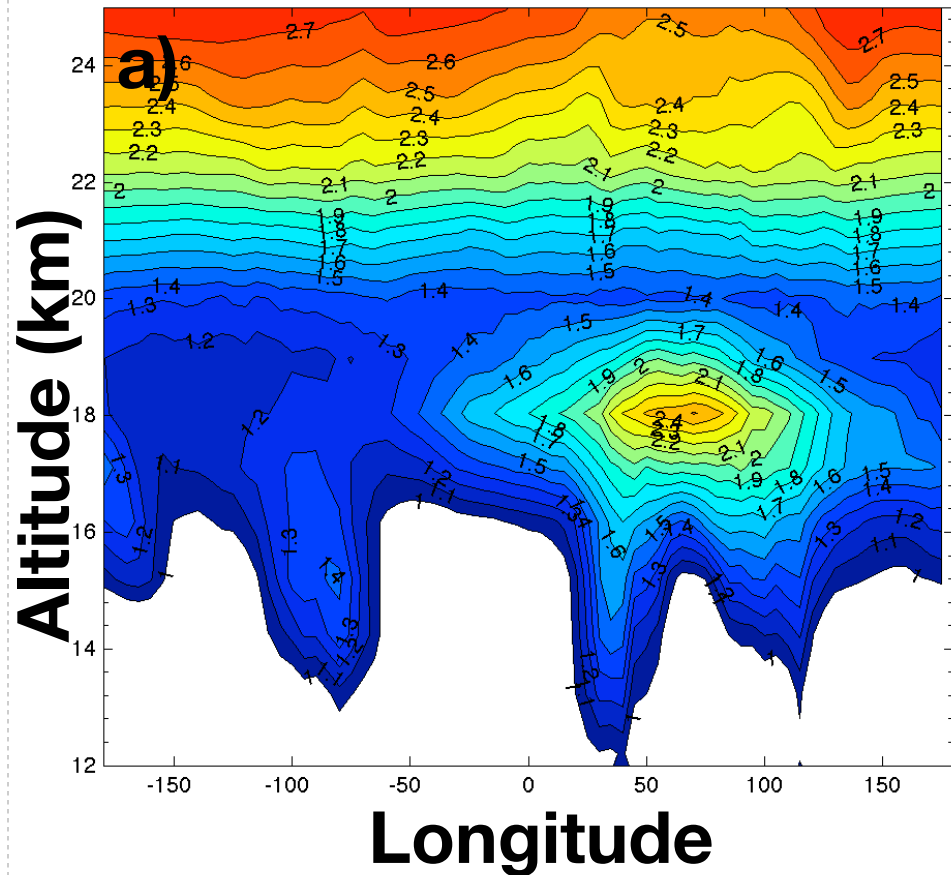
**Baseline Surface JJA  
SO<sub>2</sub> Concentrations**



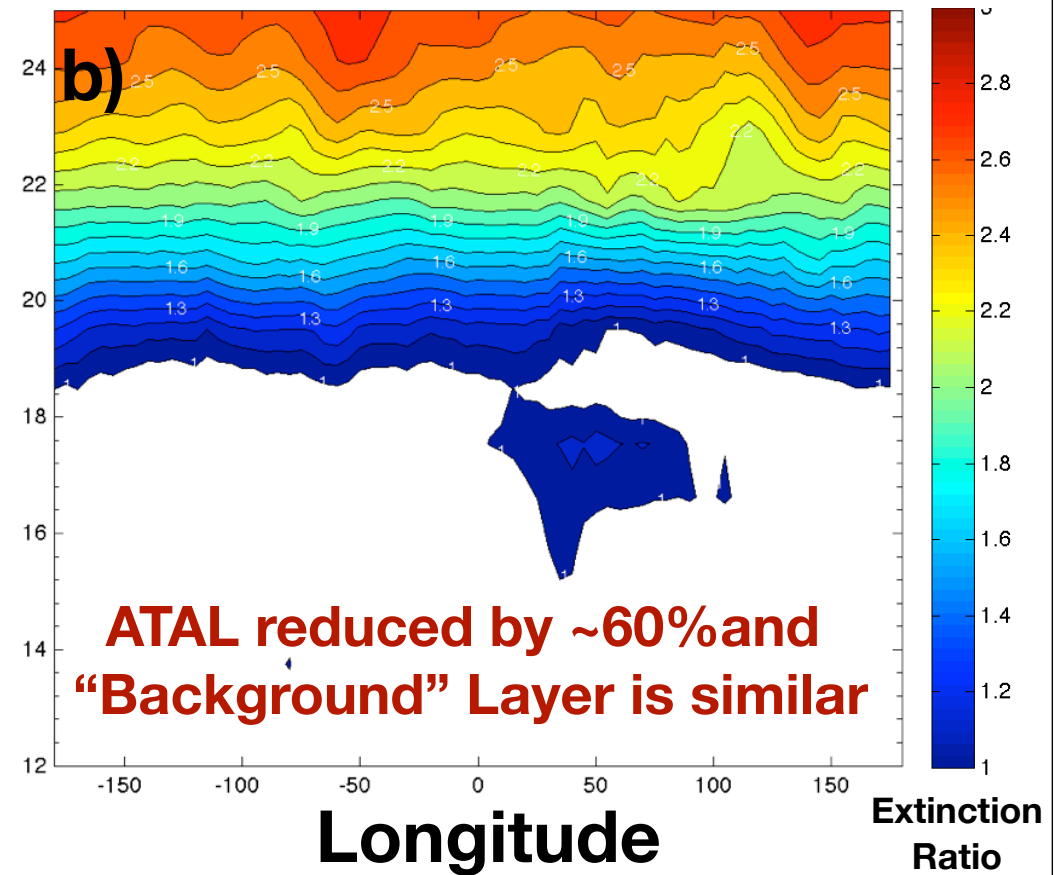
# Asian Anthropogenic Influence on the ATAL

## Modeled Mean 1020 nm Extinction Ratio from 14N to 46N, June-August

**With** Global Anthropogenic Sulfur Emissions



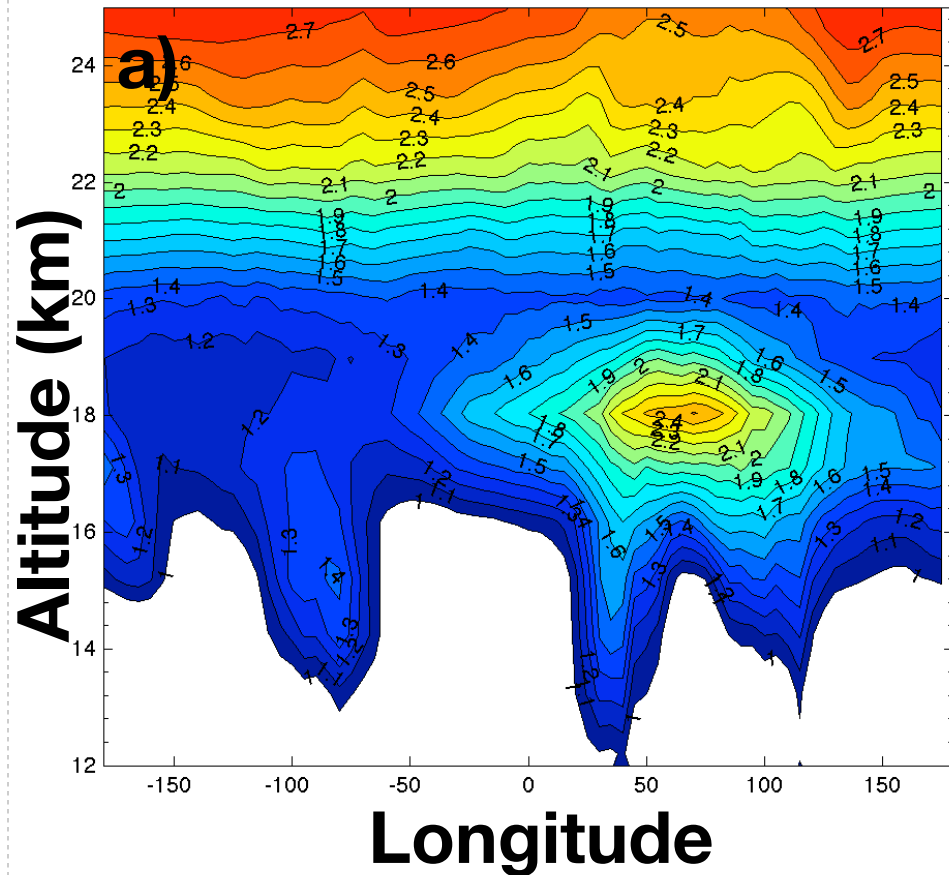
**Without** Asian Sulfur Emissions



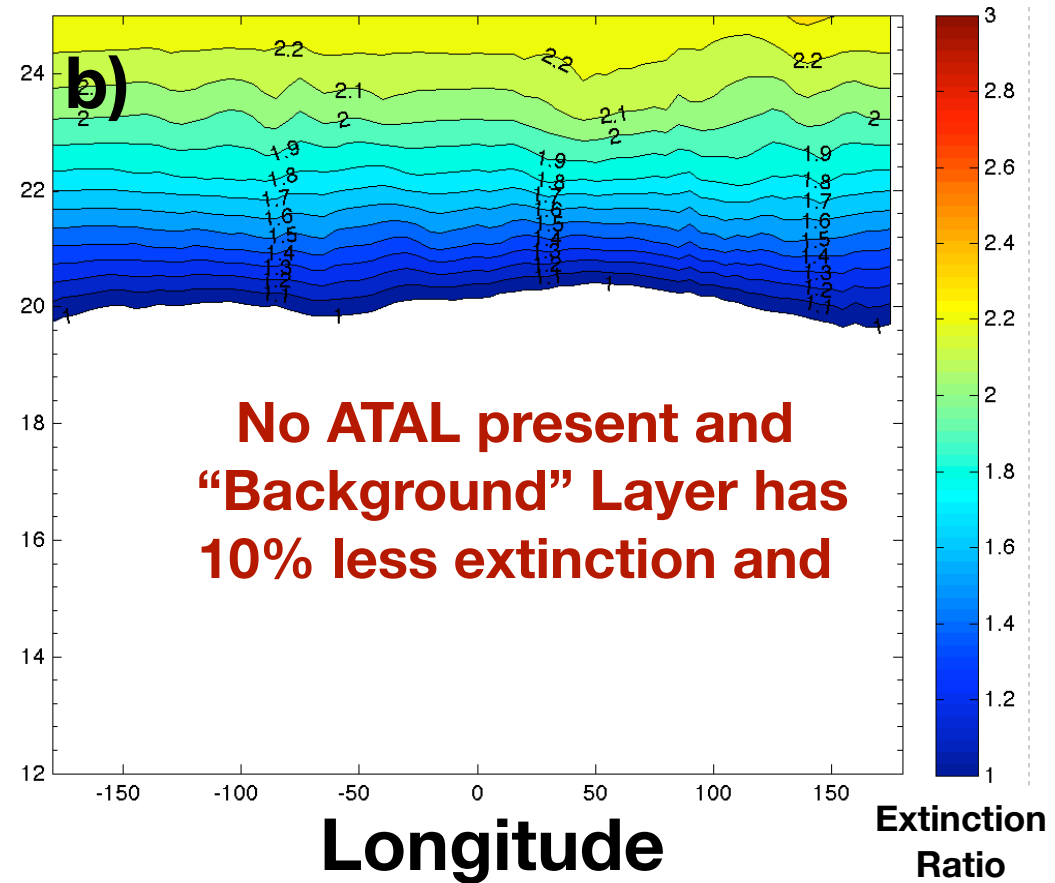
# Total Anthropogenic Influence

## Modeled Mean 1020 nm Extinction Ratio from 14N to 46N, June thru August

**With** Global Anthropogenic Sulfur Emissions



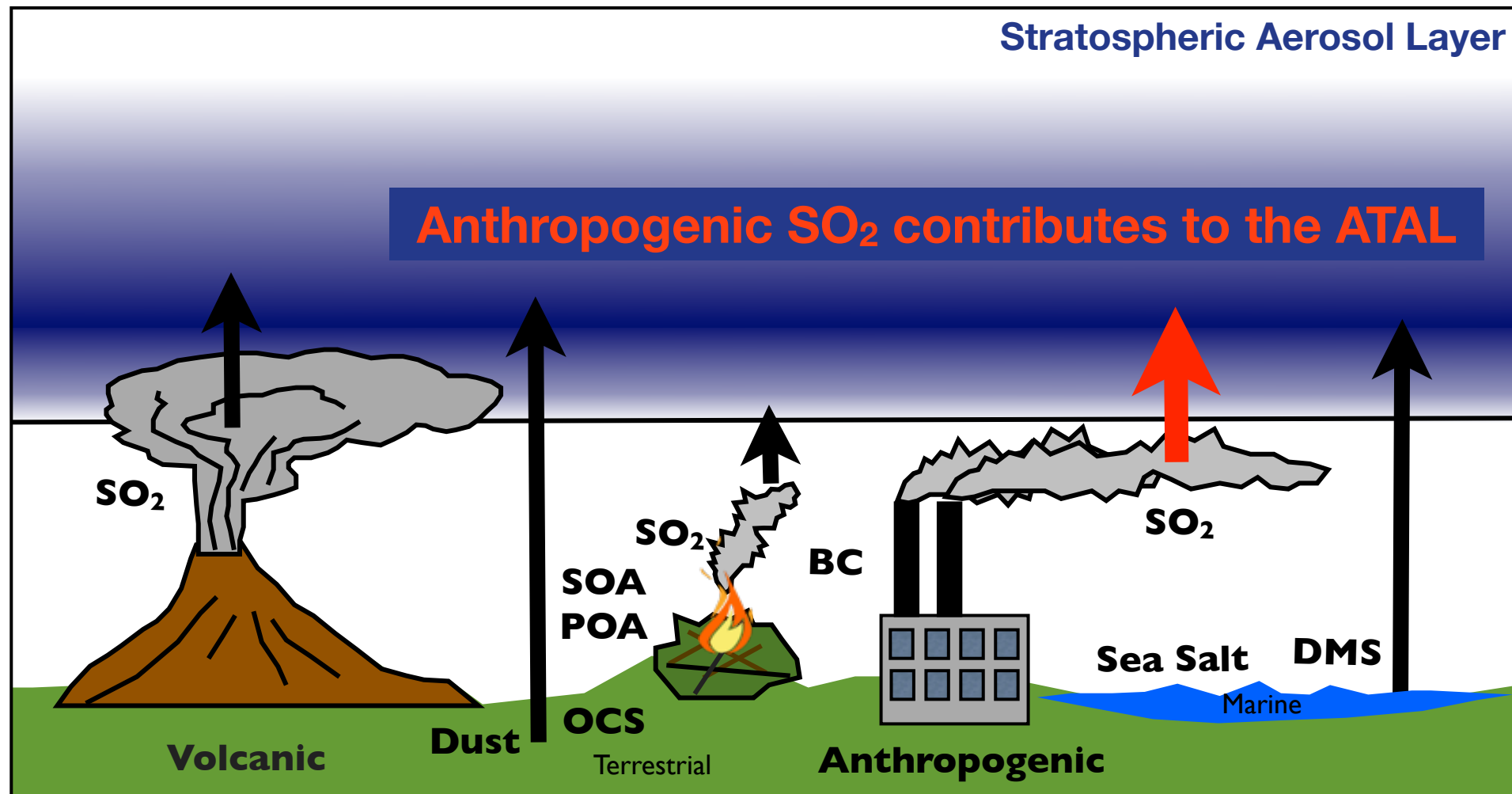
**Without** Global Anthropogenic Sulfur Emissions





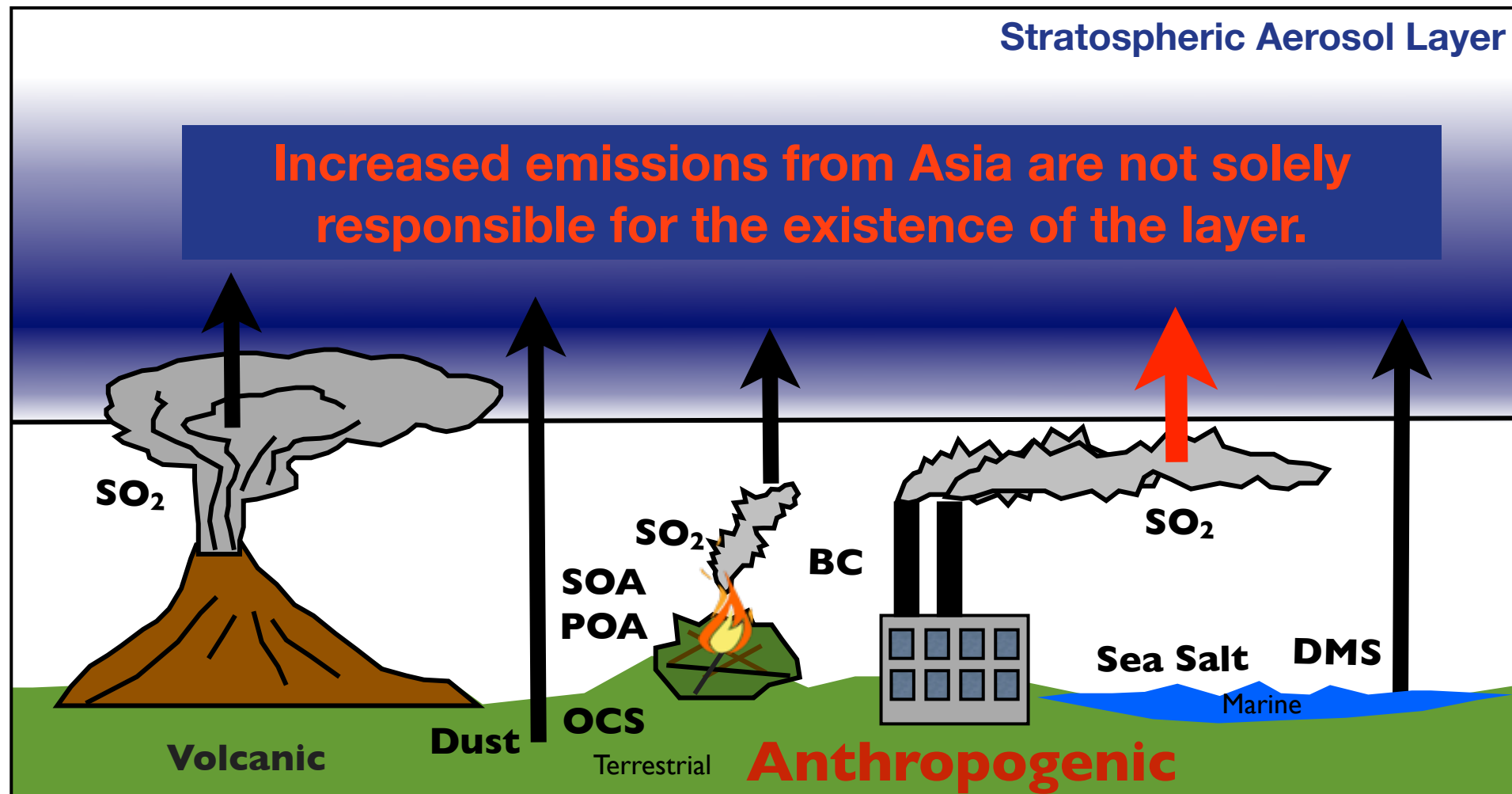


# Conclusions





# Conclusions





# Back Up Slides

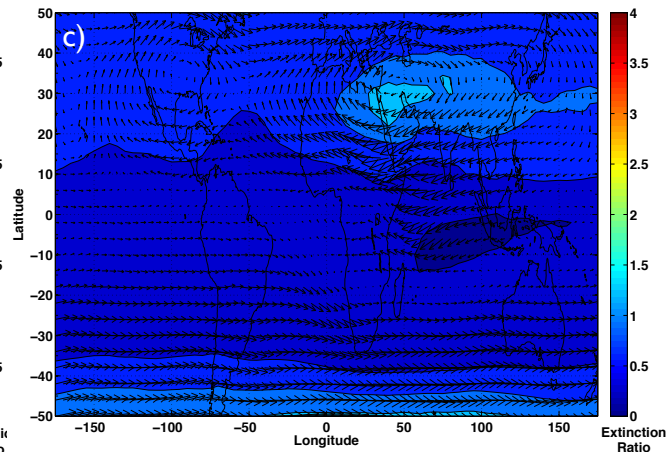
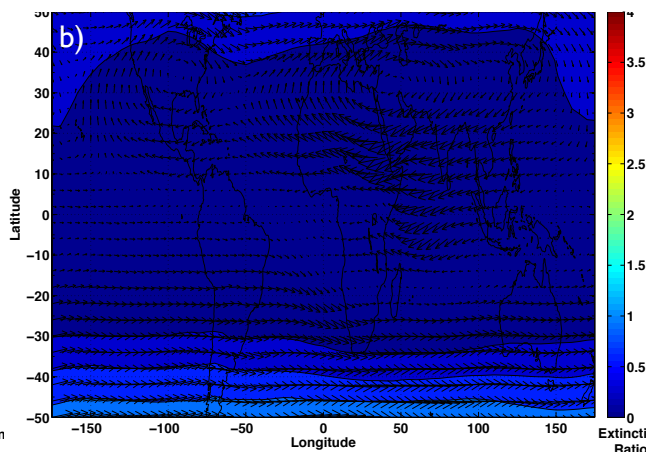
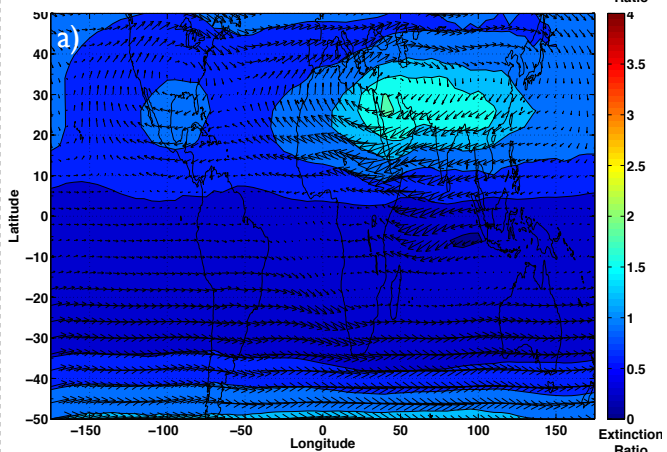
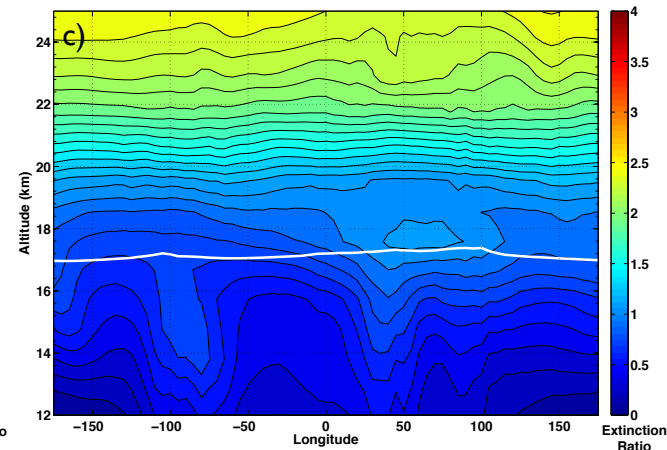
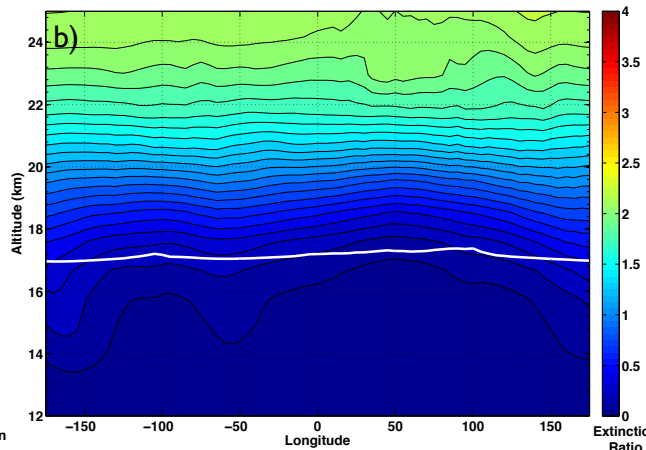
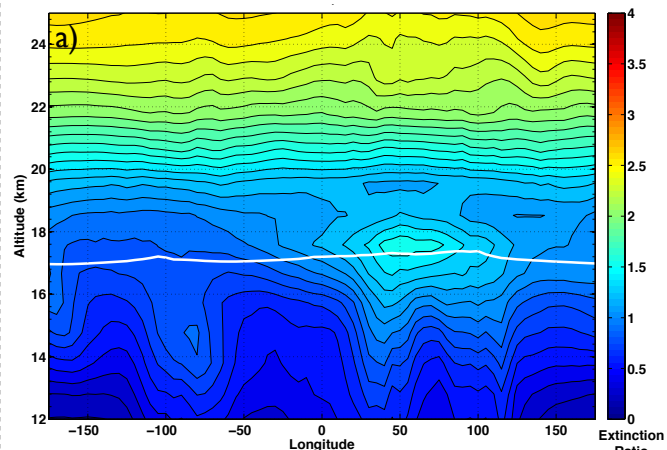


# Anthropogenic Influence Summary

2000

No Anthropogenic SO<sub>2</sub>

No China and India



No ATAL present

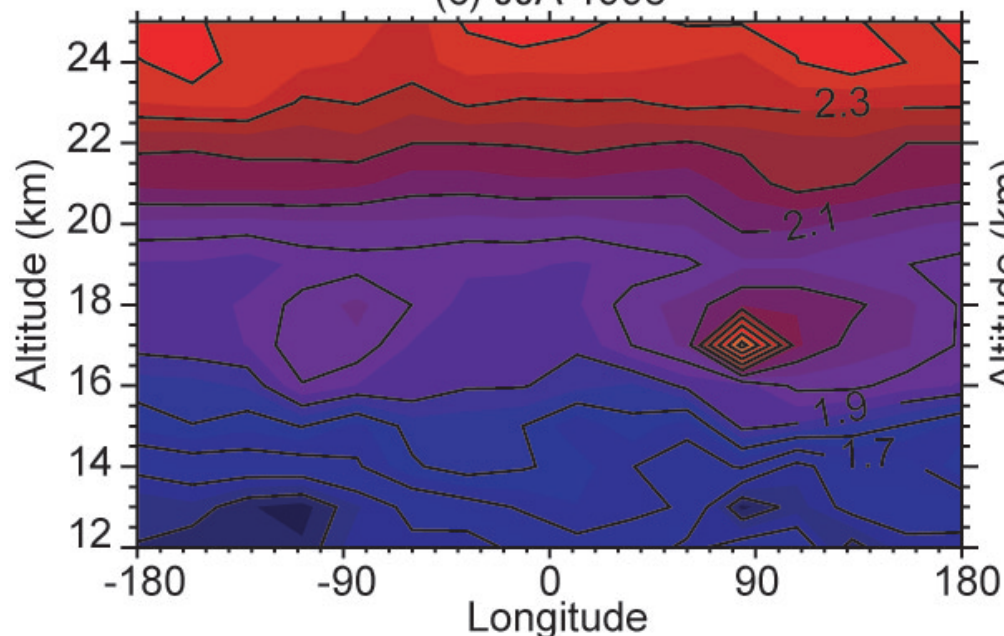
ATAL reduced by ~60%



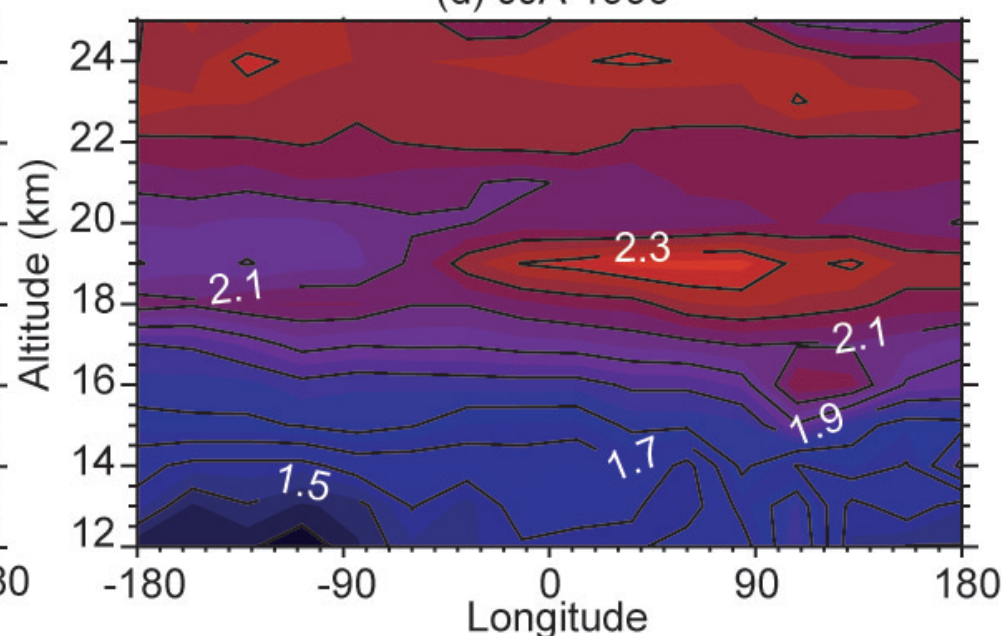
# First Observed in 1998

## Median 1020 nm Extinction Ratio Observed by SAGE II from 15N to 45N, June thru August

(c) JJA 1998

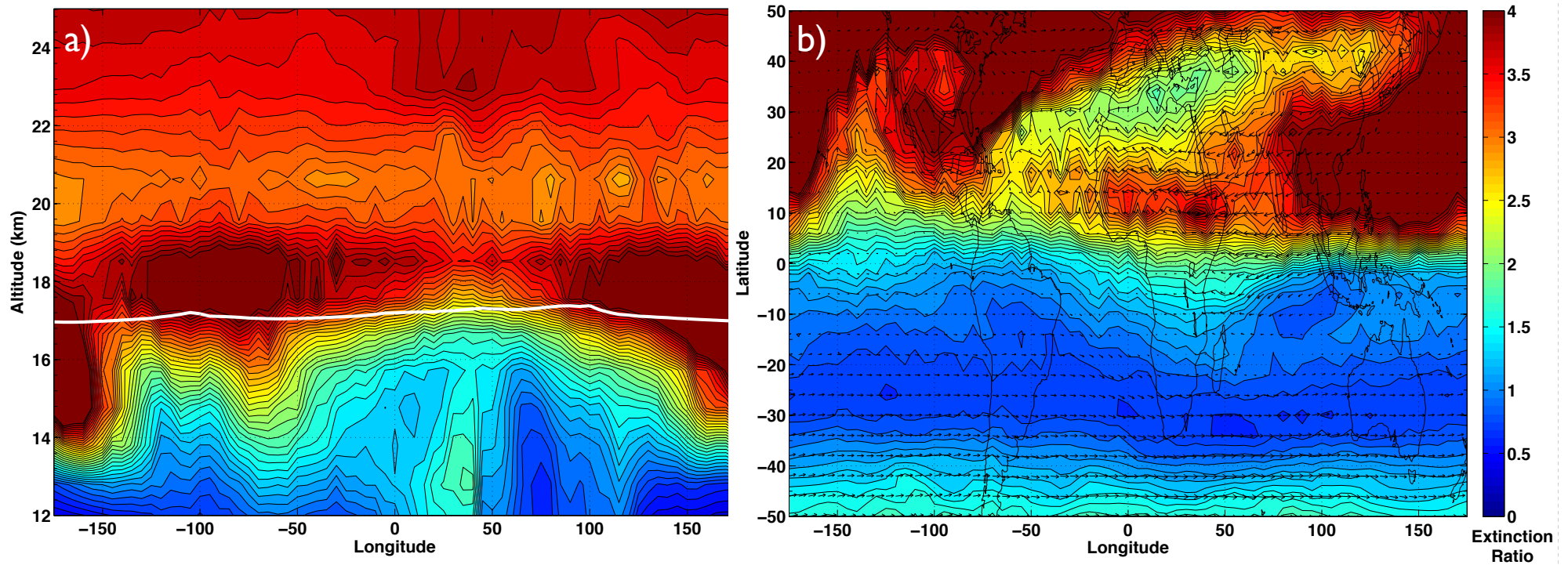


(d) JJA 1999



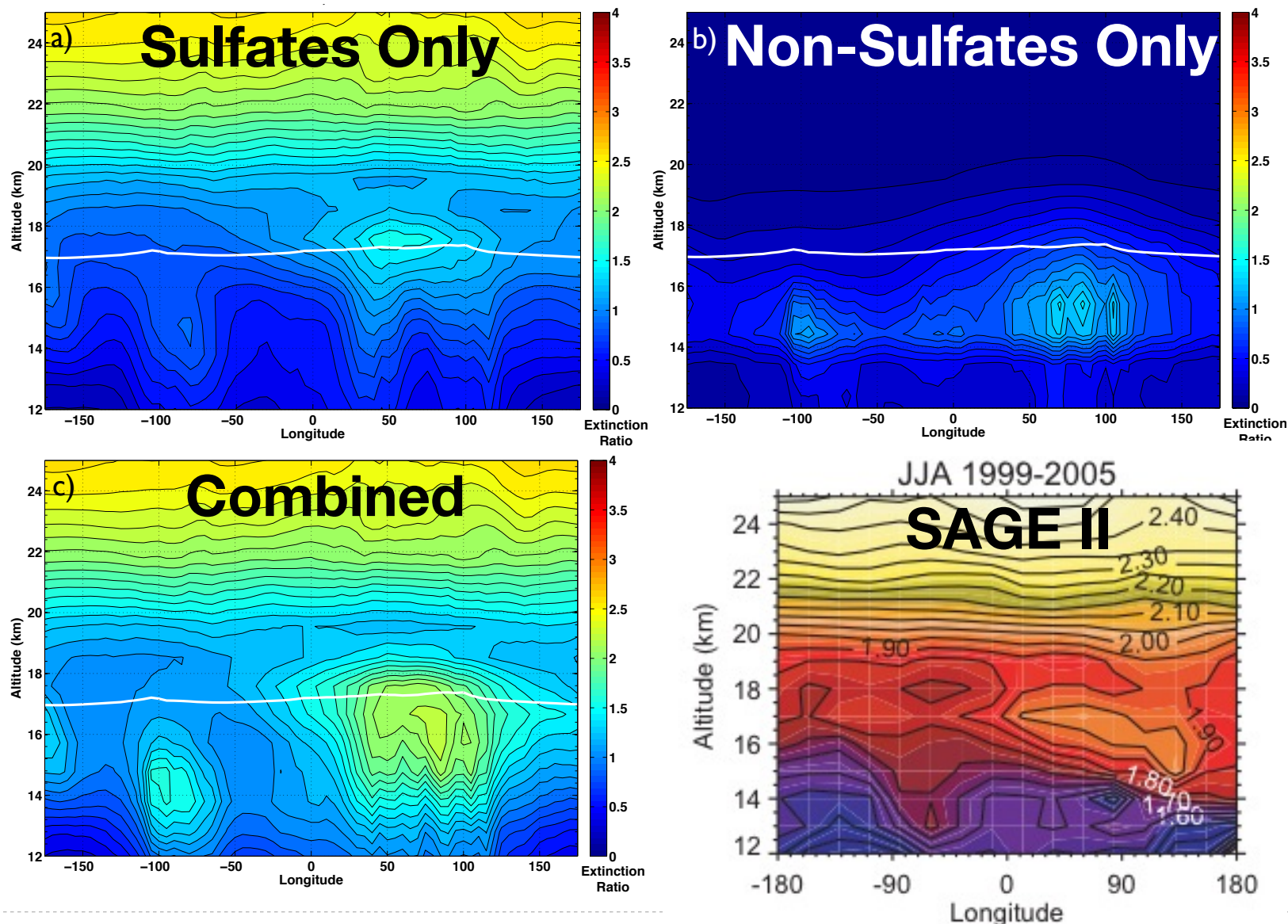
## Hypothesis: Asian Emissions Lead to ATAL

# Why was the ATAL Unobserved Until 1998?

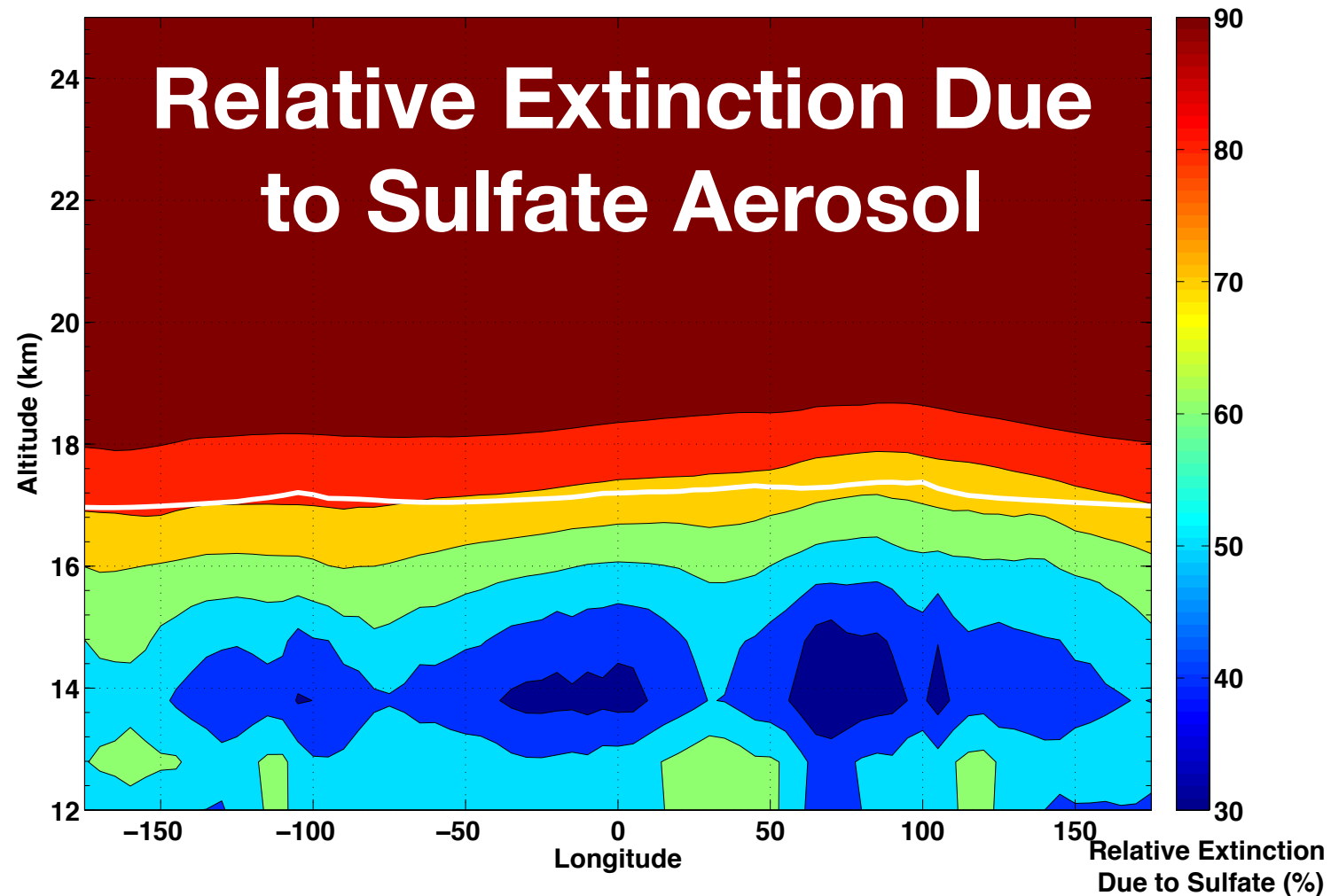




# Vertical Comparison to SAGE II (14N to 46N)



# ATAL Composition: Sulfate Splits With Non-Sulfate at Tropopause

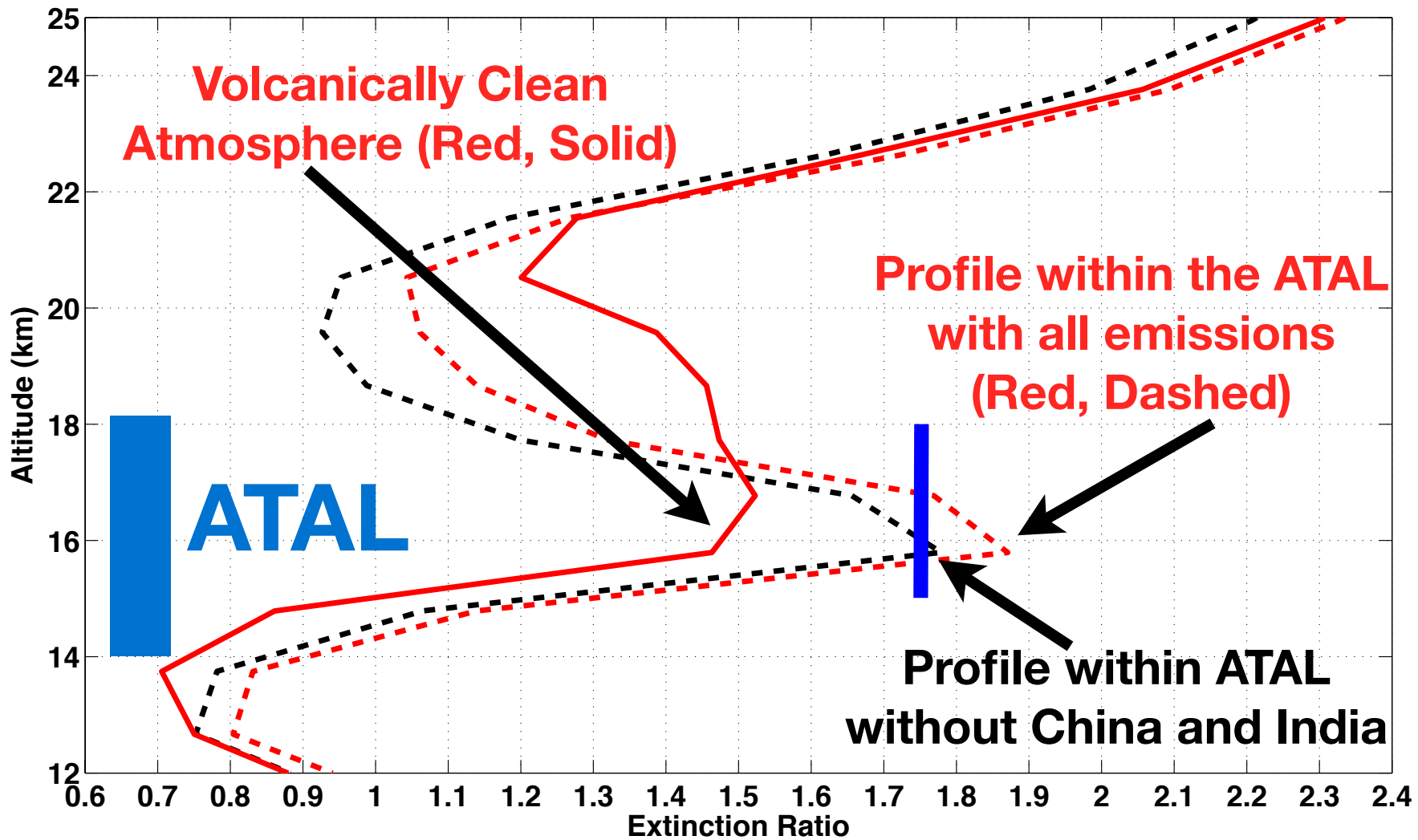


Agrees with near tropopause *in situ* observations of CARIBIC and PALMS (Papaspriopoulos et al., 2002 and Murphy et al. 1998; 2007)





# Why was the ATAL Unobserved Until 1998?



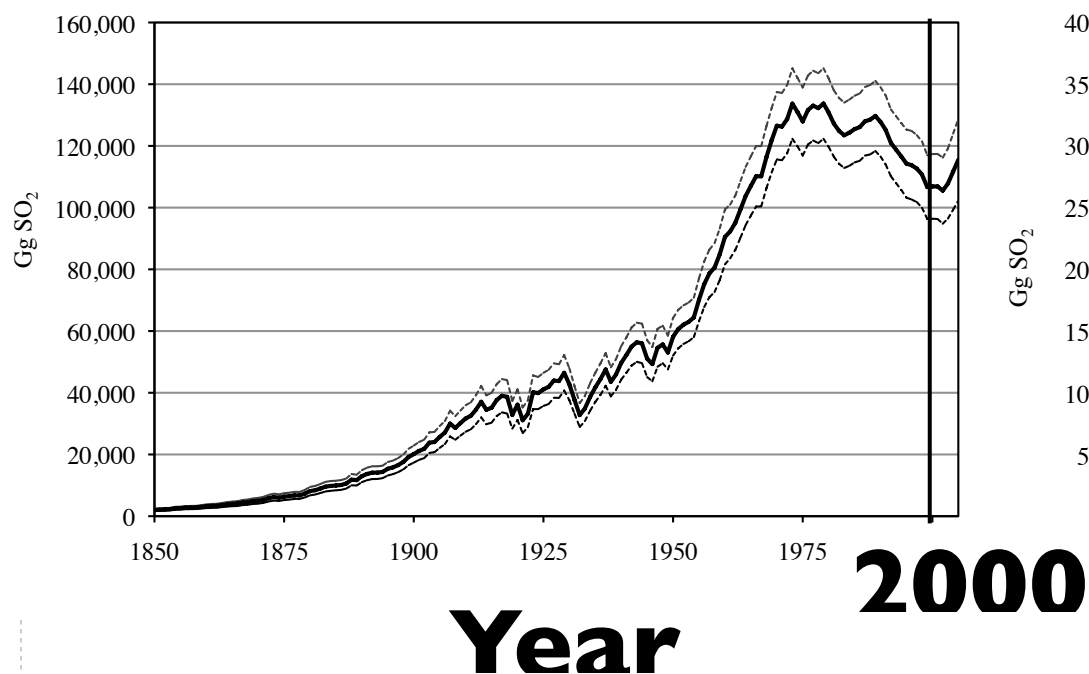
**Lower Bound of Observed 1989 Extinction Ratio from 16 to 18 km**  
**(Thomason et al, 1997)**



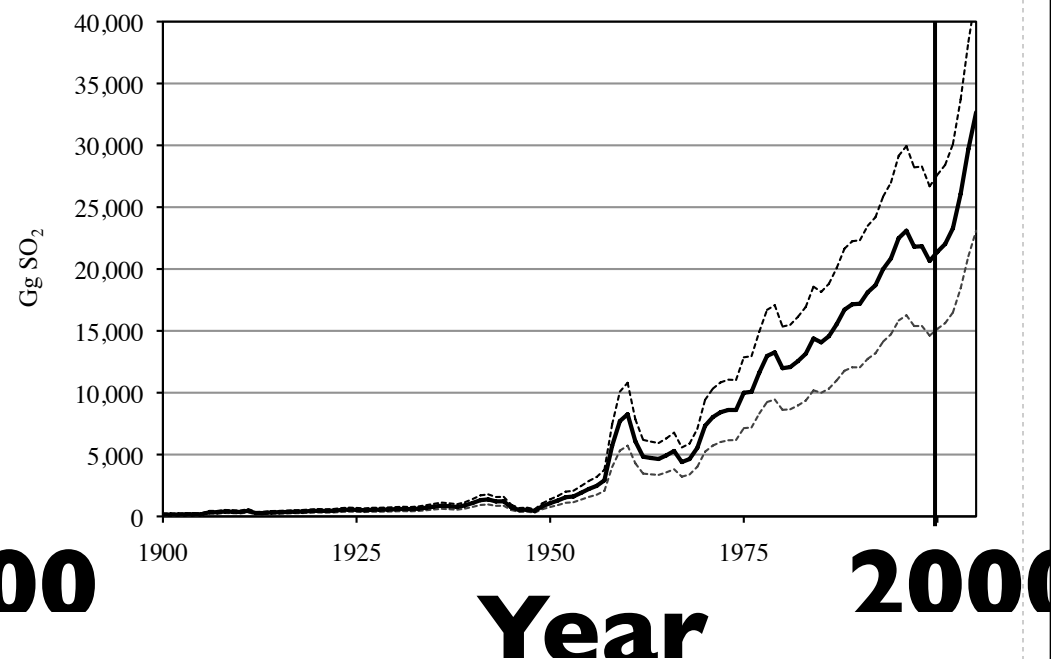


# Hypothesis: Asian Emissions Lead to ATAL

## Global $\text{SO}_2$



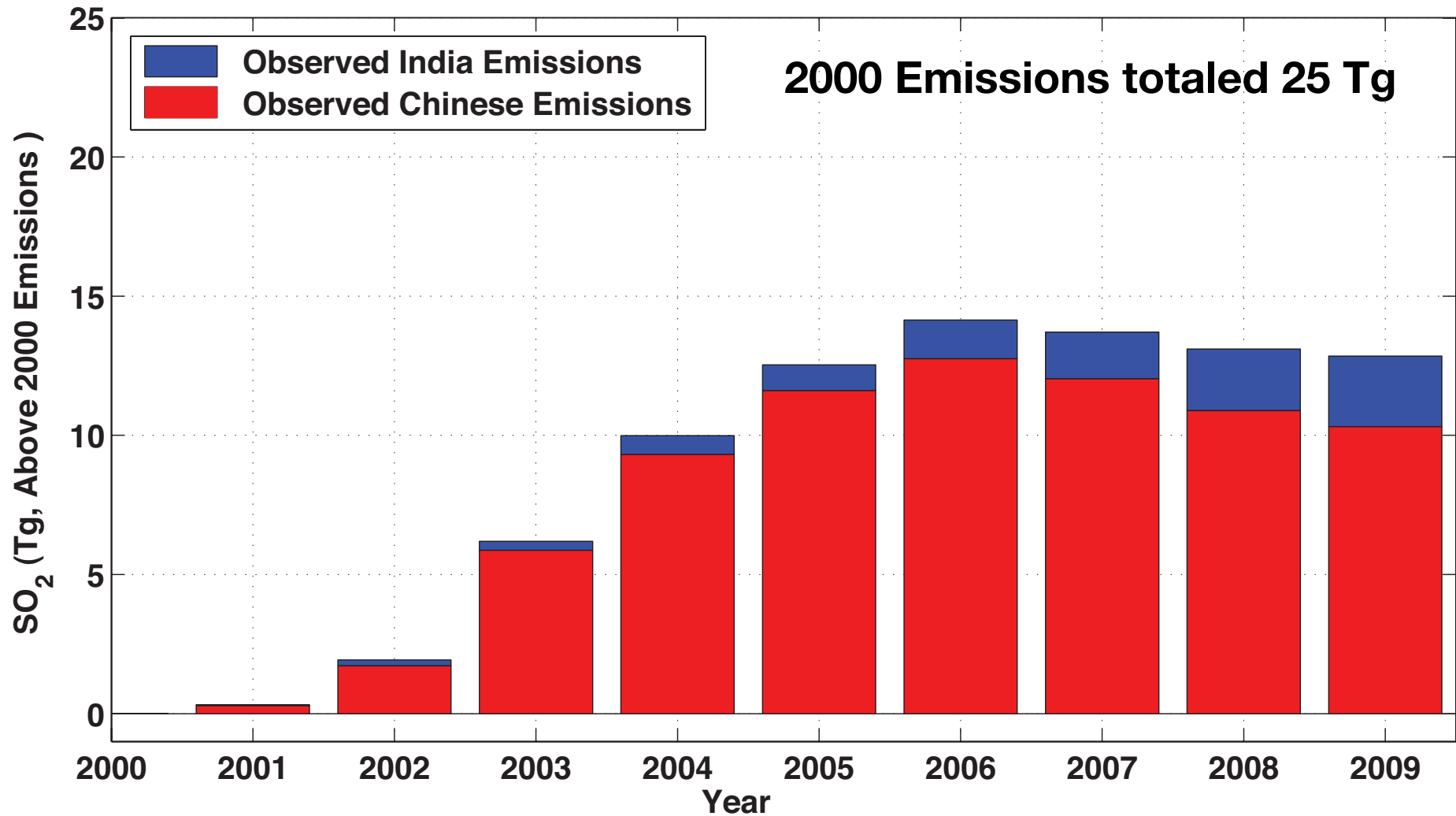
## Chinese $\text{SO}_2$



**Increases in other constituents  
of carbonaceous aerosol (POA,**

Adapted from Smith et al. ACP 2011

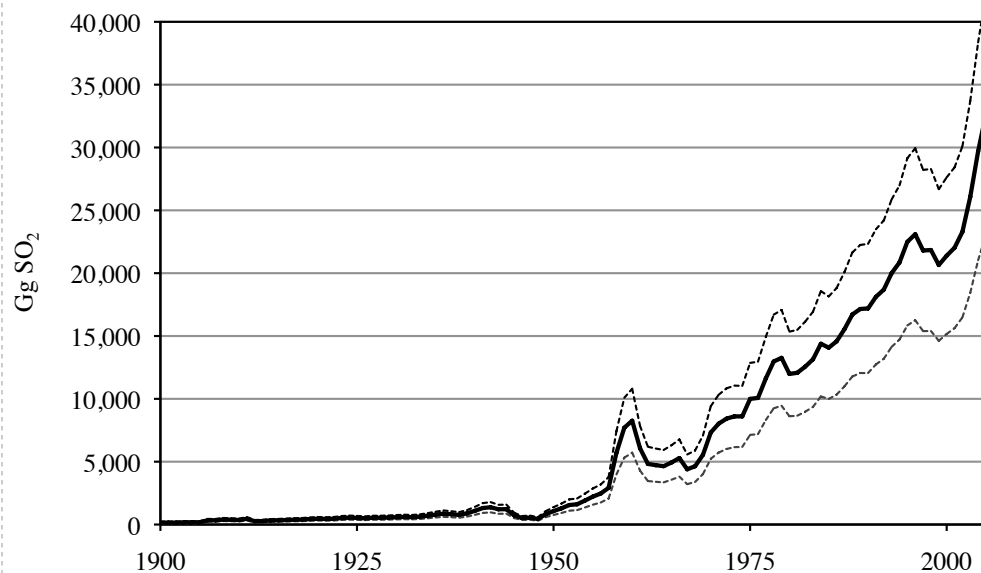
# Increases in Asian Anthropogenic Emissions since 2000



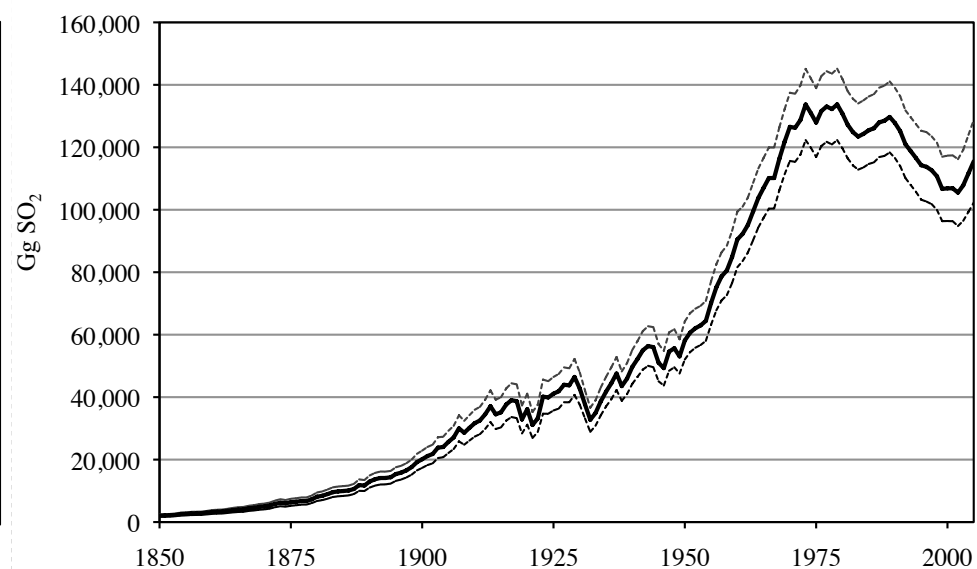
Data source: Lu, Z., Zhang, Q. & Streets, D. G. Sulfur dioxide and primary carbonaceous aerosol emissions in China and India. 1996–2010, Atmospheric Chemistry and Physics 11, 9839–9864 (2011).

# Asian Emissions

China SO<sub>2</sub> Emissions



Global SO<sub>2</sub> Emissions

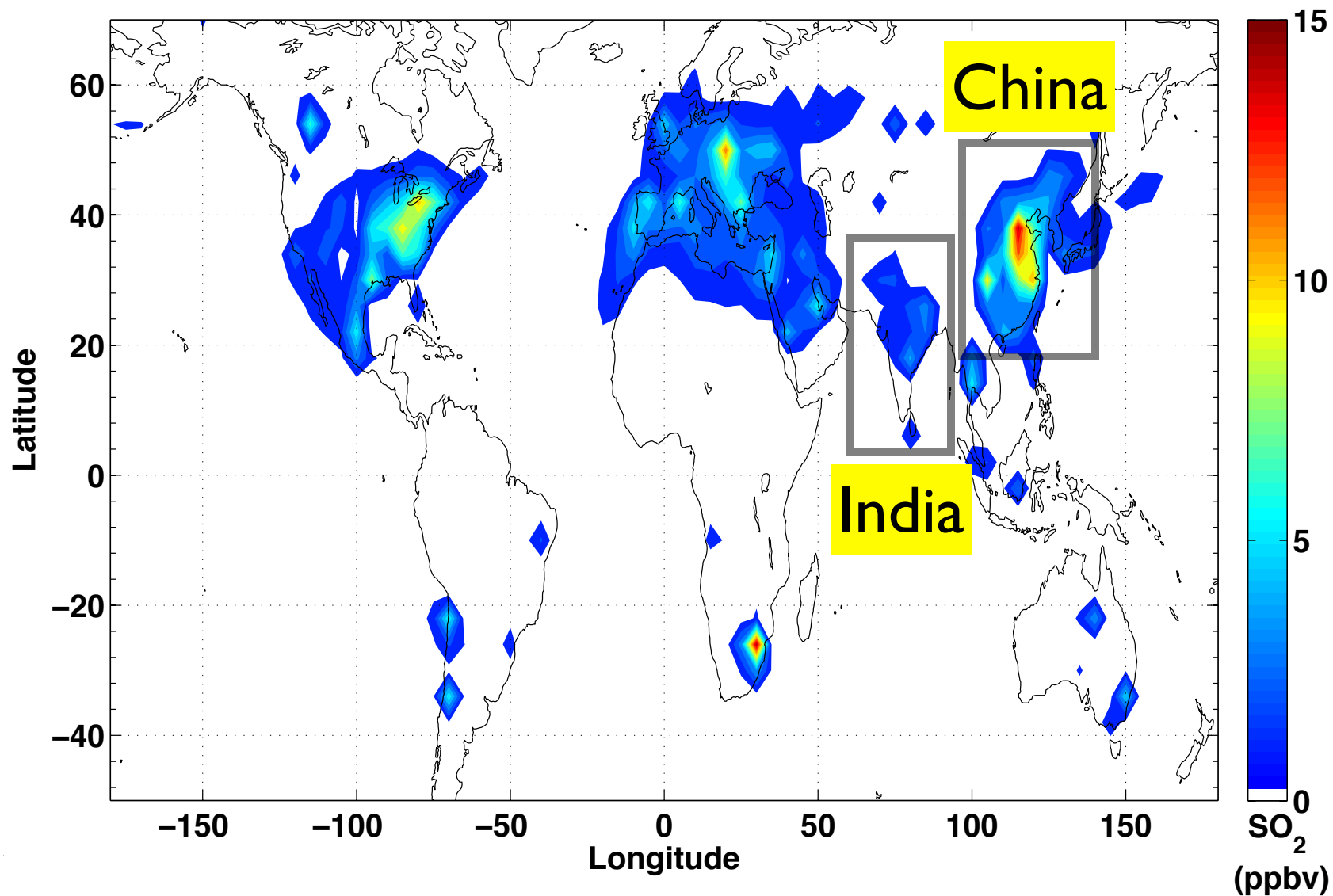


Adapted from Smith, S. J., J. van Aardenne, Z. Klimont, R. J. Andres, A. Volke, and S. Delgado Arias (2011), Anthropogenic sulfur dioxide emissions: 1850–2005, *Atmos. Chem. Phys.*, 11(3), 1101–1116, doi:10.5194/acp-11-1101-2011.

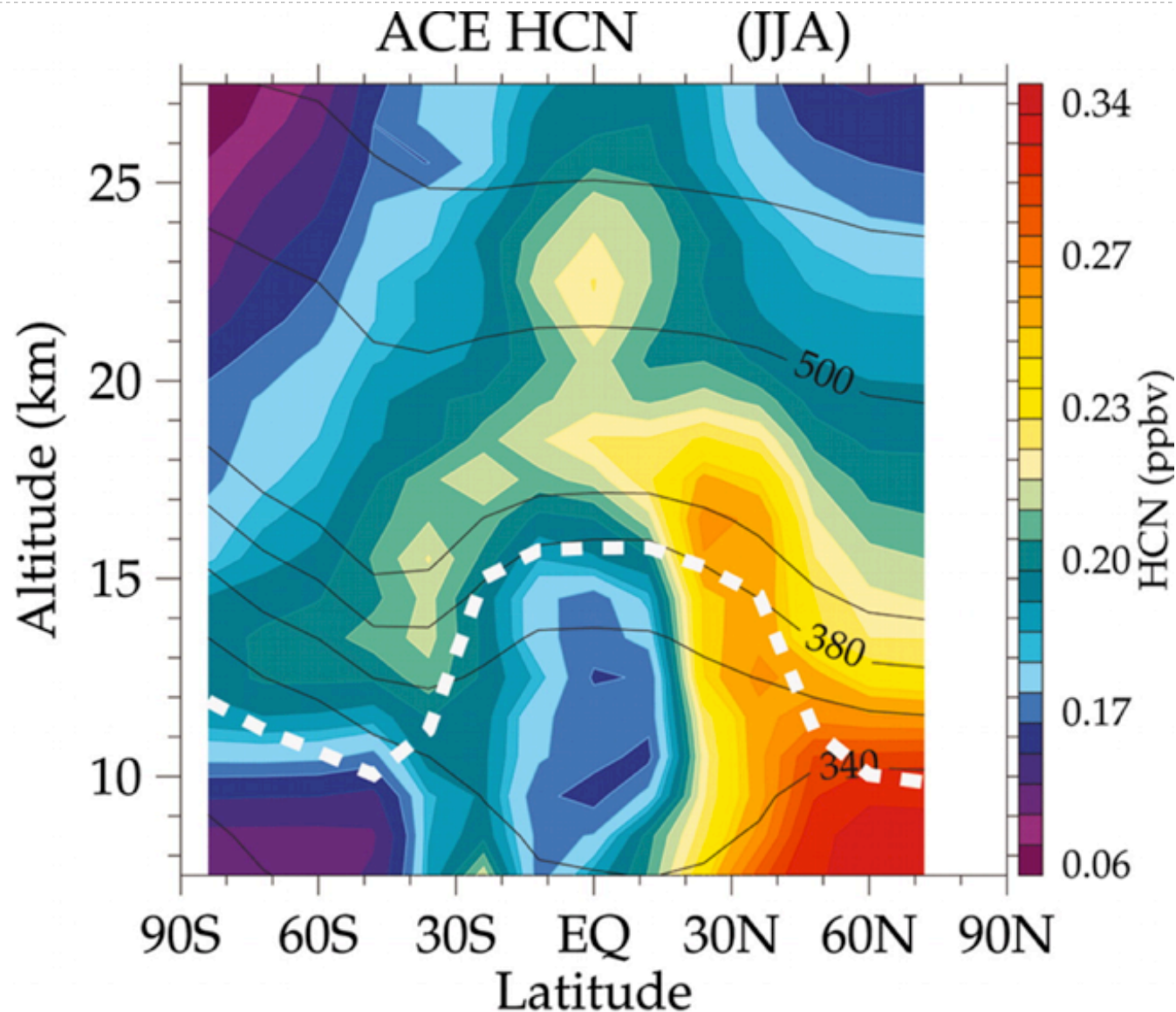


# Experiment: Modulate Global Anthropogenic Emissions

## Global JJA Model Surface SO<sub>2</sub> Concentrations

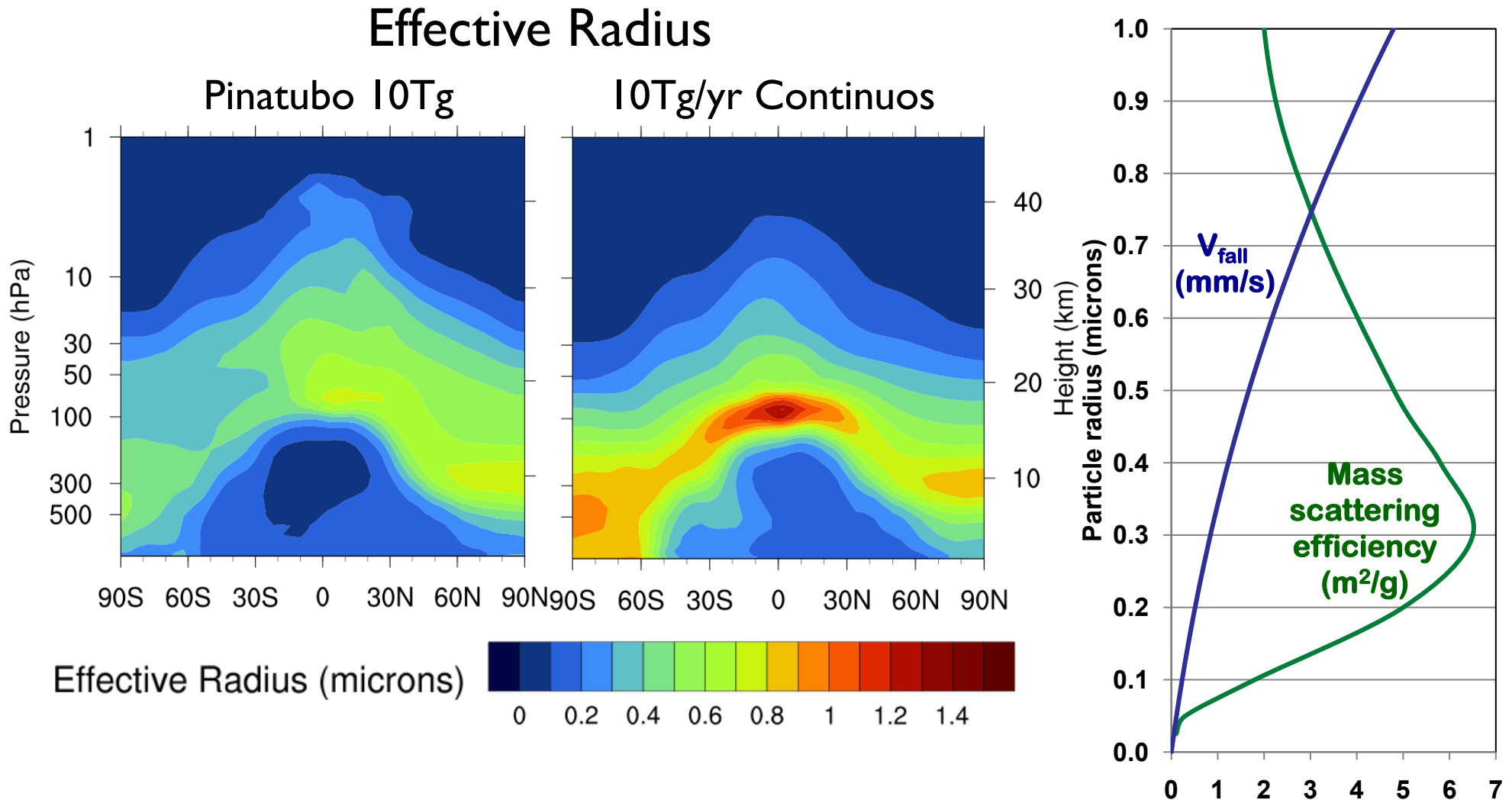


# Anthropogenic emissions transported to the stratosphere via the Asian Monsoon



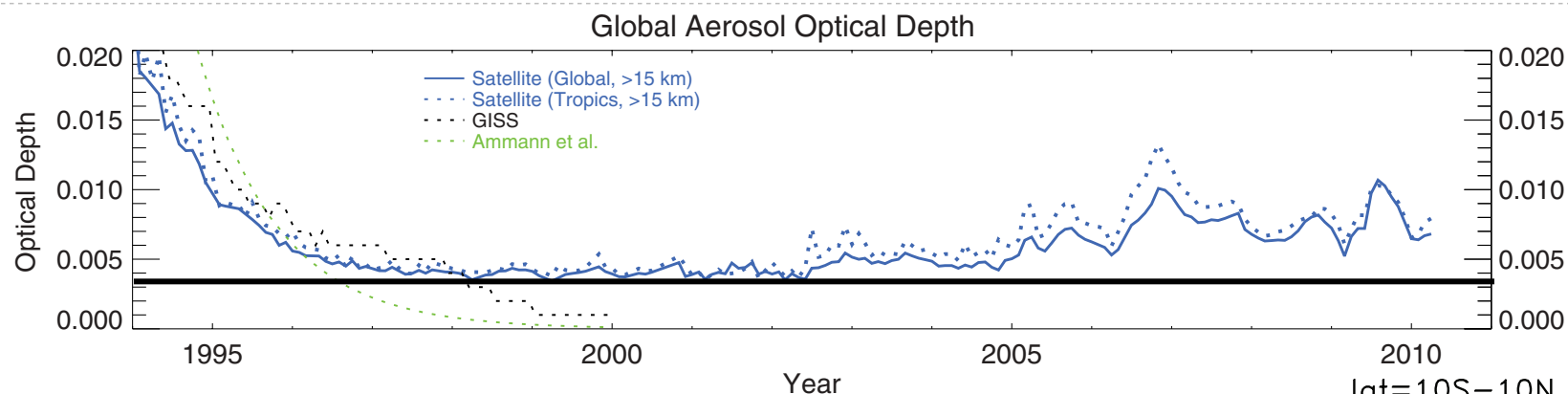
# Why Are Volcanoes More Efficient at Making Aerosol?

27



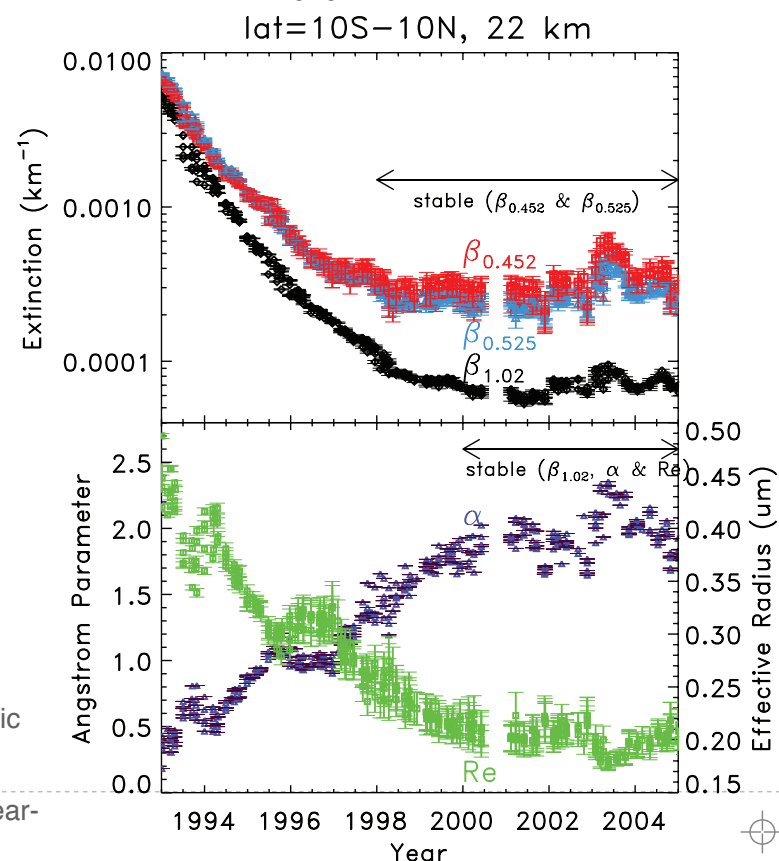


# 2000 to 2010 is an unprecedented “background” period



**Layer became stable only in 2000.**

**Previous observations will be influenced  
by 1991 Pinatubo eruption.**



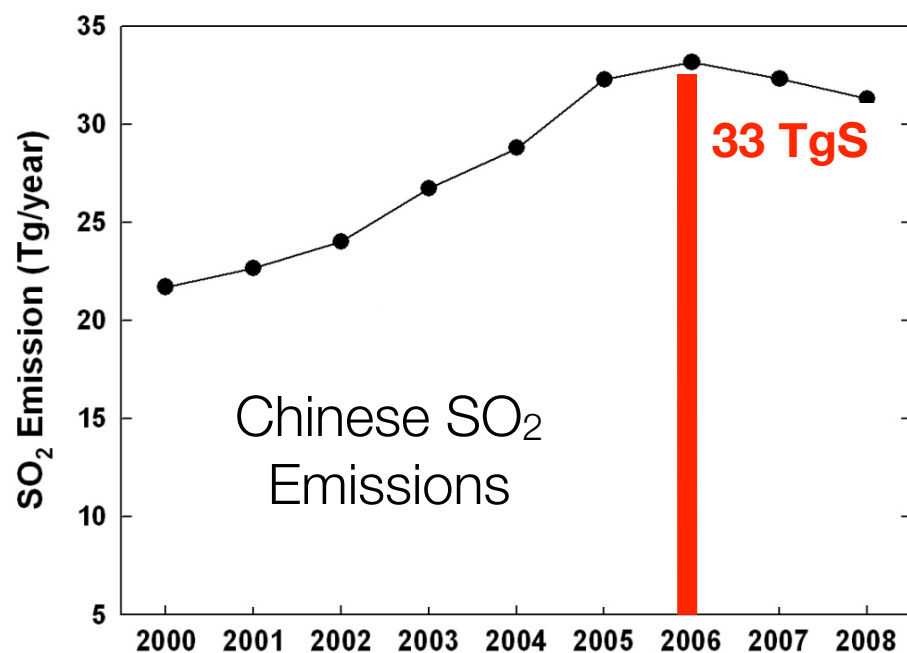
Adapted from Solomon et al. (2011), The Persistently Variable “Background” Stratospheric Aerosol Layer and Global Climate Change, *Science*

Niwano et al. (2009), Seasonal cycles of Stratospheric Aerosol and Gas Experiment II near-background aerosol in the lower stratosphere, *J. Geophys. Res.*, 114(D14), D14306.

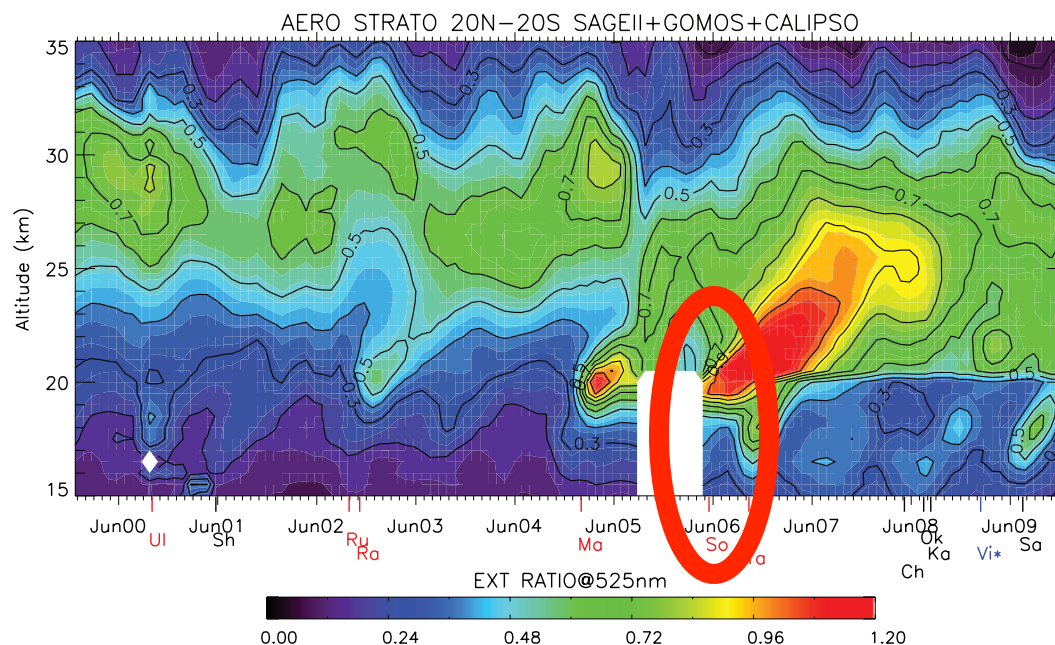




# Anthropogenic vs Volcanic emissions



Lu et al. (2010), Sulfur dioxide emissions in China and sulfur trends in East Asia since 2000, *Atmos. Chem. Phys.*, 10(13)



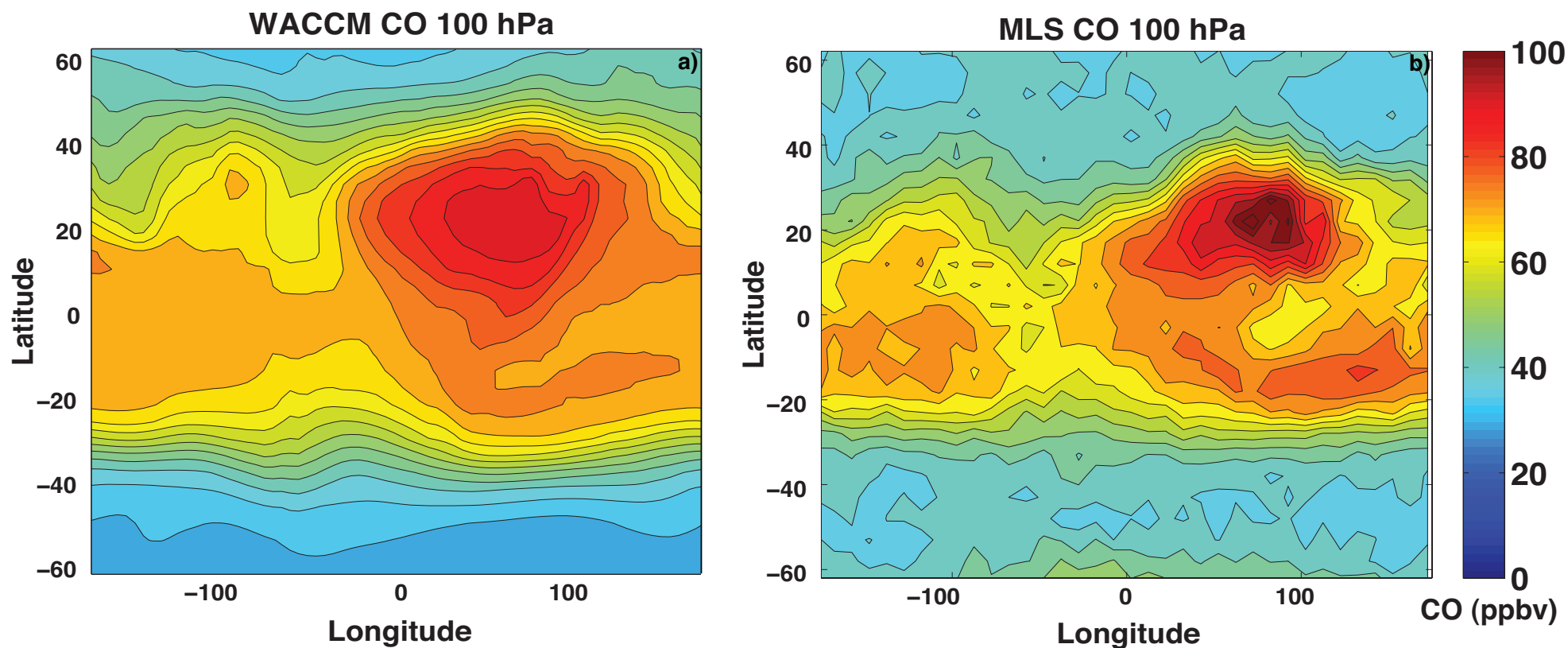
**Estimated Emission to Stratosphere**  
(0.6% of Global Emissions must make it to stratosphere to maintain sulfur burden (Hofmann et al. 2009))

Year	China	Volcano
2006	0.2 TgS	Soufrière Hills 0.17 TgS





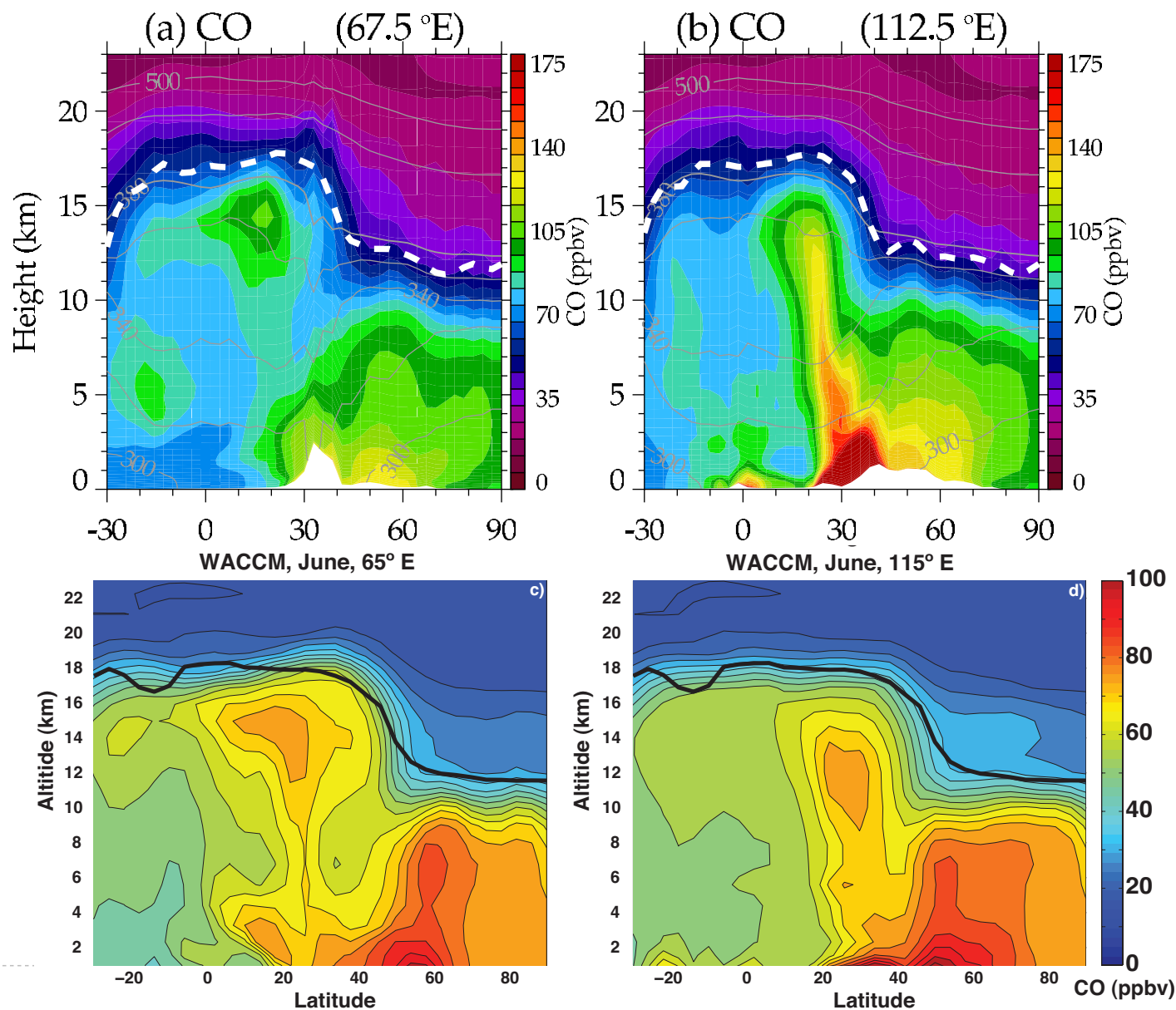
# Transport in WACCM



# Transport pathways of carbon monoxide in the Asian summer monsoon diagnosed from Model of Ozone and Related Tracers (MOZART)

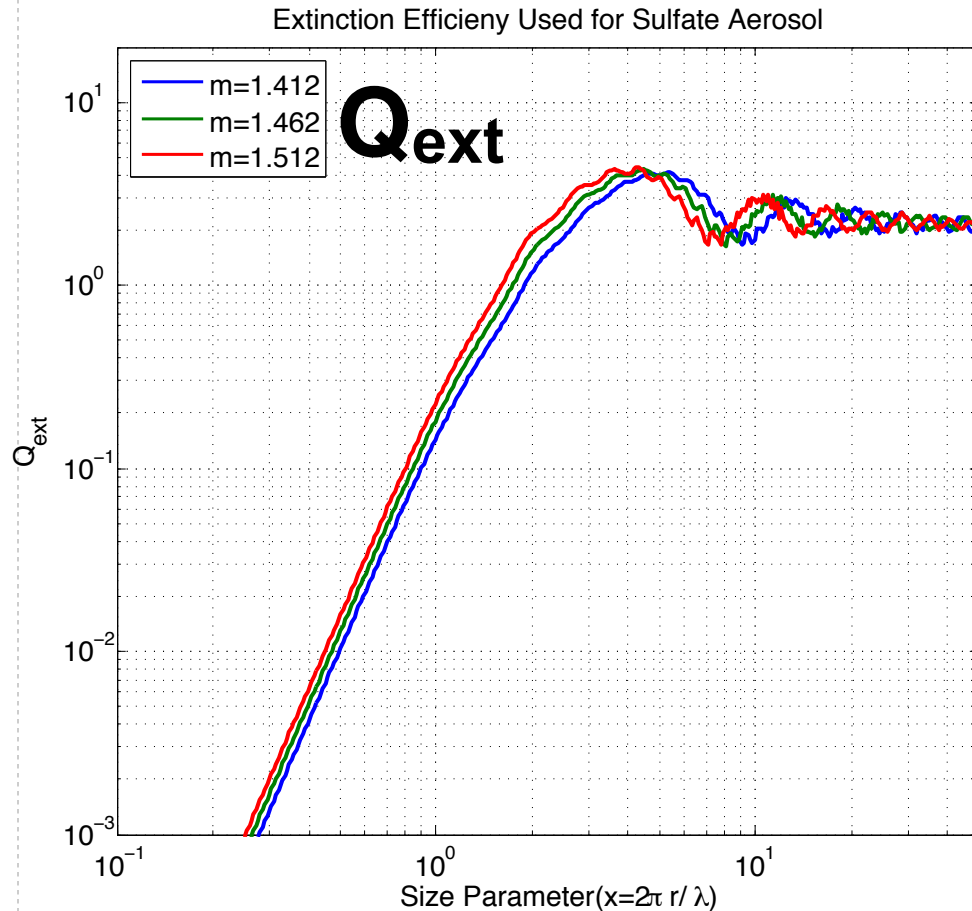
Mijeong Park,<sup>1</sup> William J. Randel,<sup>1</sup> Louisa K. Emmons,<sup>1</sup> and Nathaniel J. Livesey<sup>2</sup>

**Figure 7.** Latitude-altitude cross-sections of monthly mean MOZART-4 CO at the (a) western (67.5°E) and (b) eastern (112.5°E) sides of the monsoon maximum in June 2005. Thermal tropopause derived from the model temperature profile is denoted as thick dashed lines. Thin solid lines are isentropes (320, 340, 360, 380, 450, and 500 K).

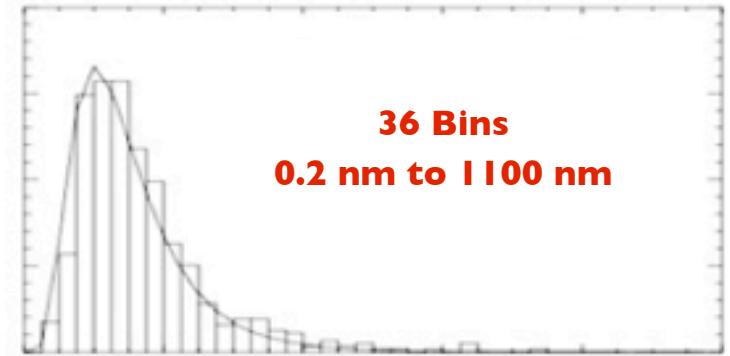




# Scattering Calculations



$$n(r, z) =$$



$$\beta_a(z, t) = \pi \int_0^\infty r^2 Q_\pi(\tilde{m}, x) n(r, z) dr$$

Bohren and Huffman (1983)