Transport, chemistry and aerosols in the Asian monsoon anticyclone

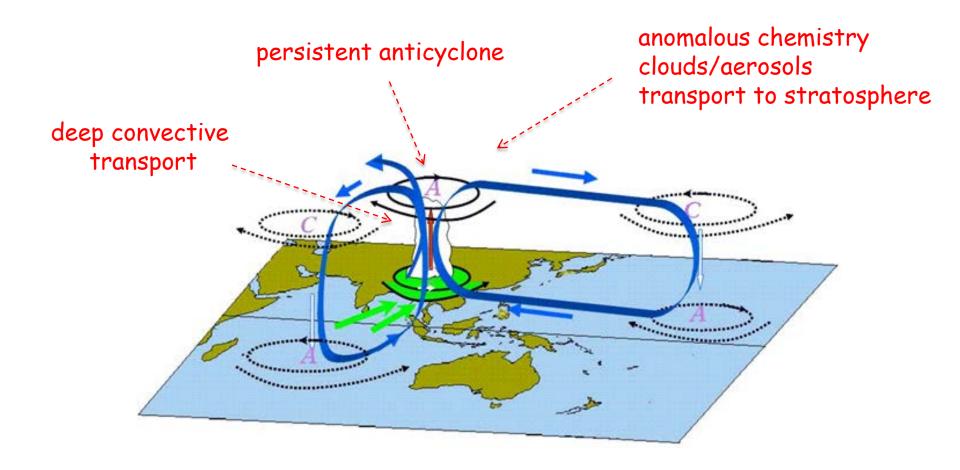
Bill Randel

Atmospheric Chemistry Division NCAR Boulder, Colorado USA

Thanks to: Mijeong Park, Laura Pan, Doug Kinnison, Hella Garney, others.

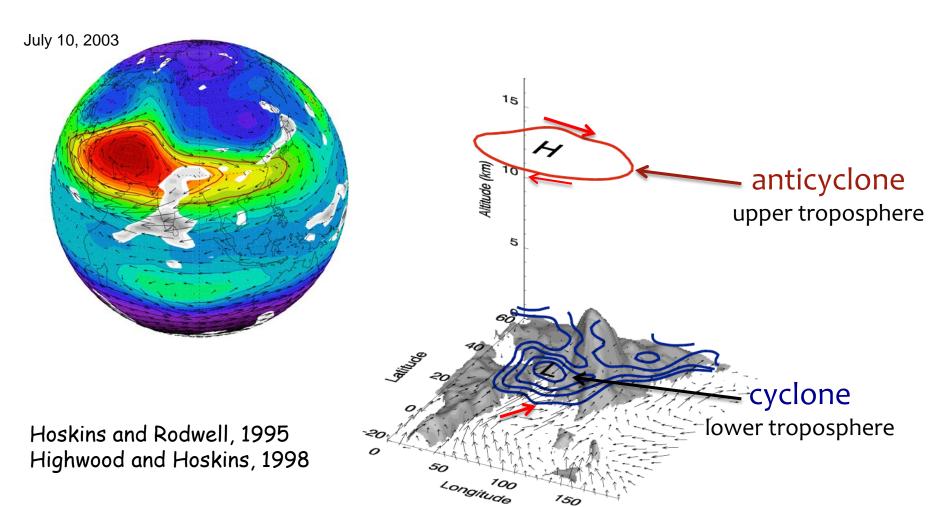


Summer Broad-Scale Circulations

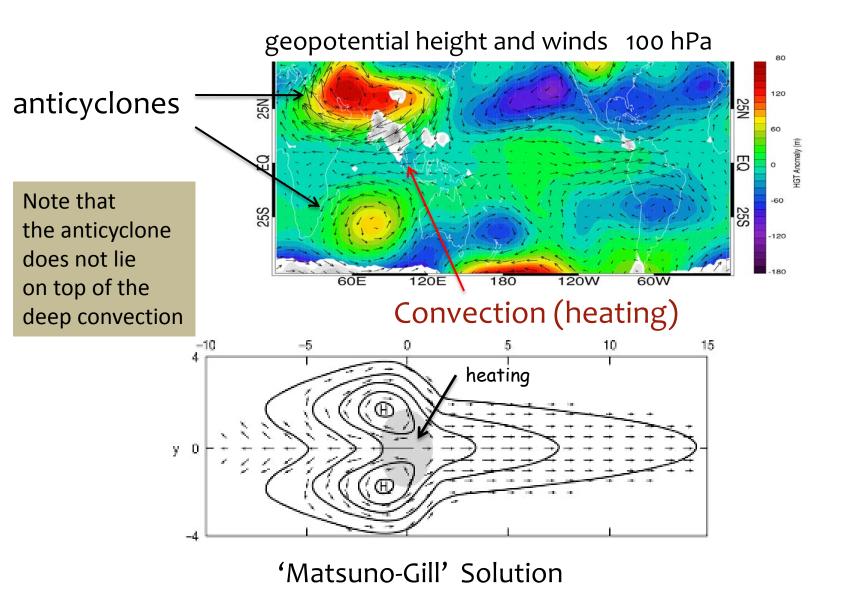


Dynamical Background

Cyclone at the surface, anticyclone in the upper troposphere

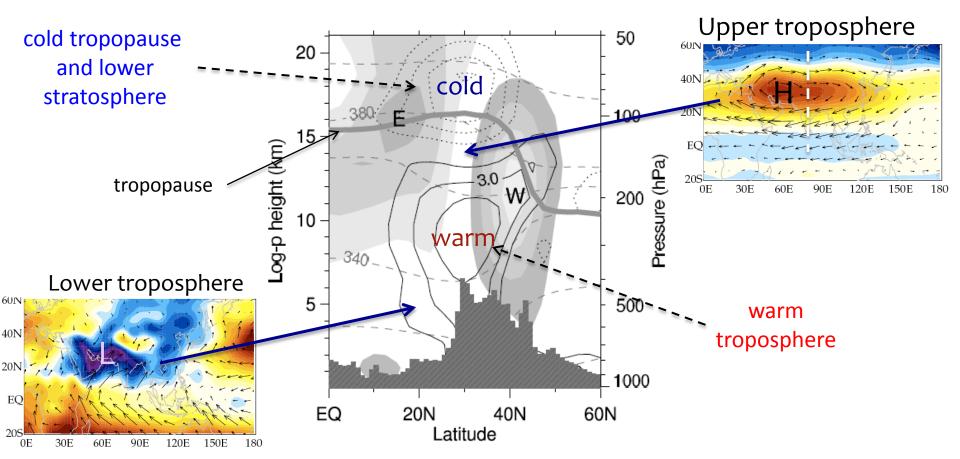


Anticyclones in the Upper Troposphere



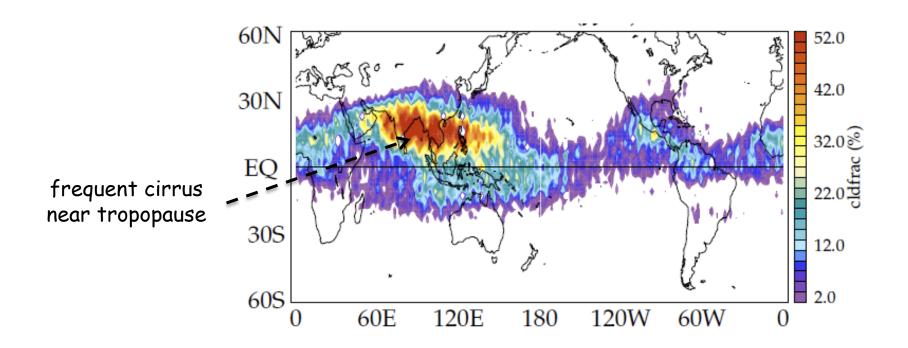
Dynamical Background

Anticyclonic circulation extends into lower stratosphere

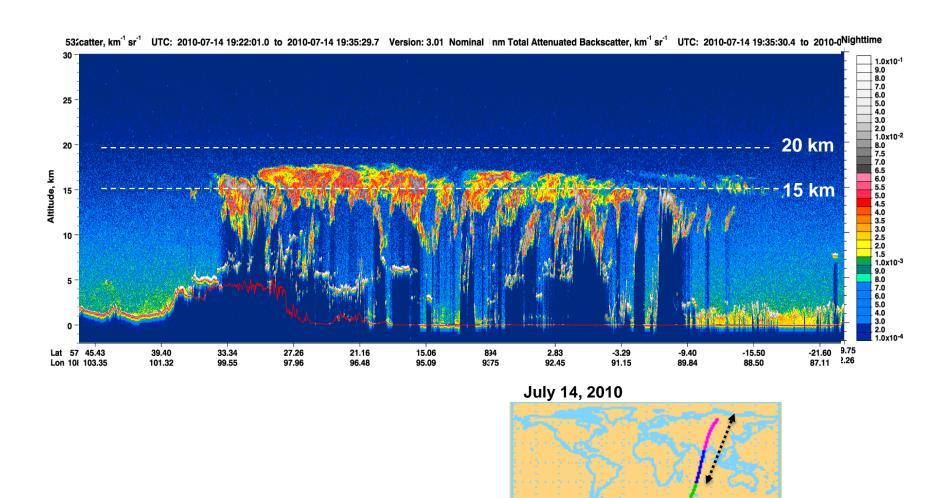


Randel and Park, JGR, 2006

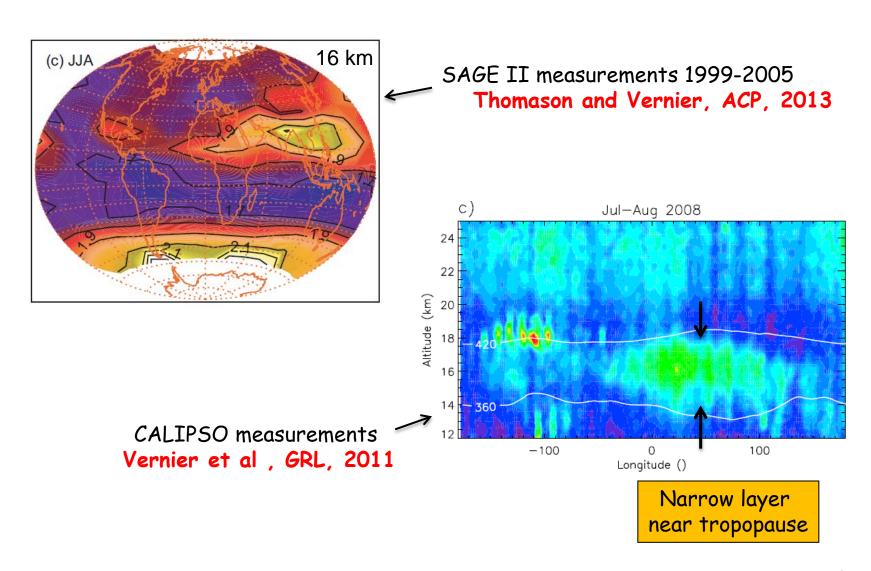
CALIPSO cloud fraction near 16 km



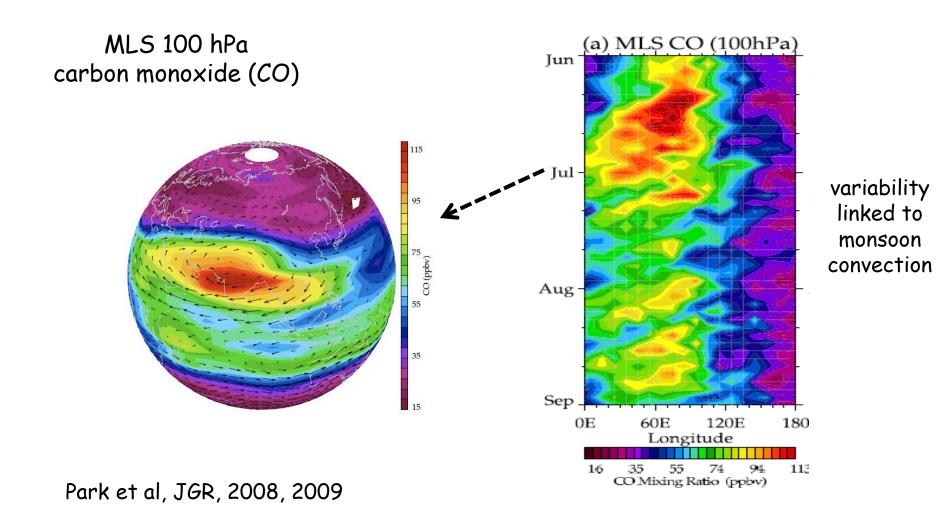
CALIPSO satellite lidar cloud observations



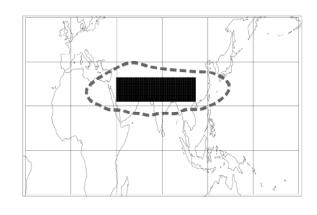
Monsoon <u>aerosol layer</u> near 16 km



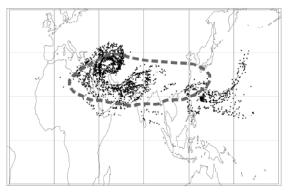
strong chemical influence on summer UTLS



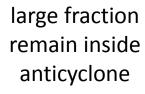
transport simulation at 150 hPa

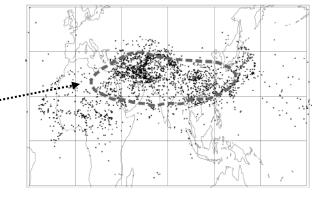


day 0



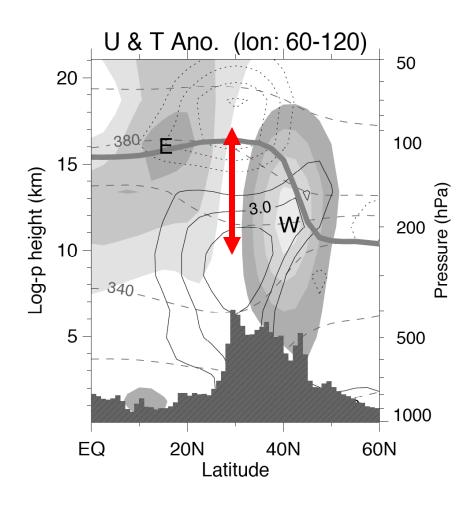
day 10



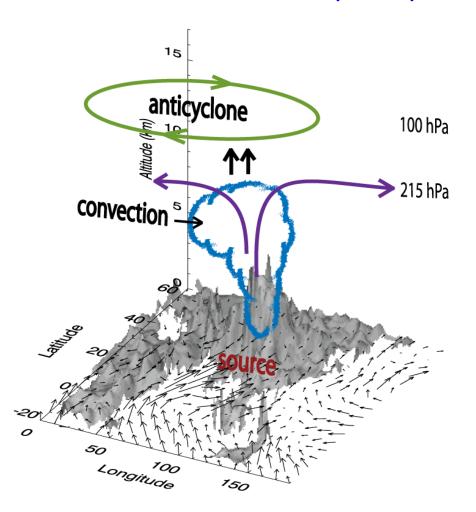


day 20

Confinement within region of strongest winds



Transport pathways



confinement by anticyclonetransport to stratosphere



Transport above 200 hPa by convection / circulation



convective transport

(main outflow near 200 hPa)



surface emission

(India and Southeast Asia)

What happens 3D trajectories initialized at 200 hPa to the outflow in regions of from deep deep convection convection? OLR < 160 K 45°8 -90°E 90°W 90"N 45"N -+ 10 days 45°8 confinement within anticyclone 90°8 + transport to SH onew. SOME 90"N 45°N + 20 days from Hella Garney 45°S

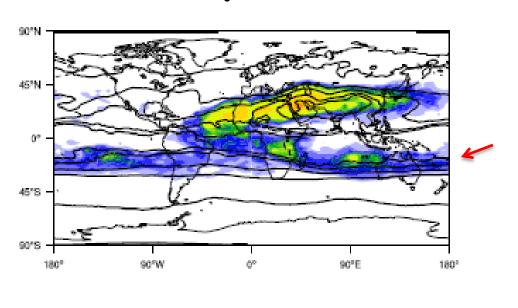
90°E

180*

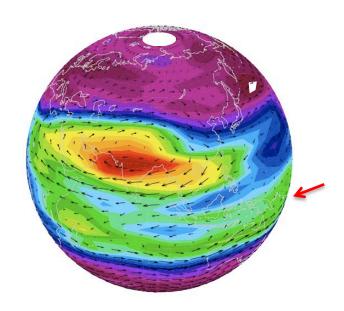
180°

90°W

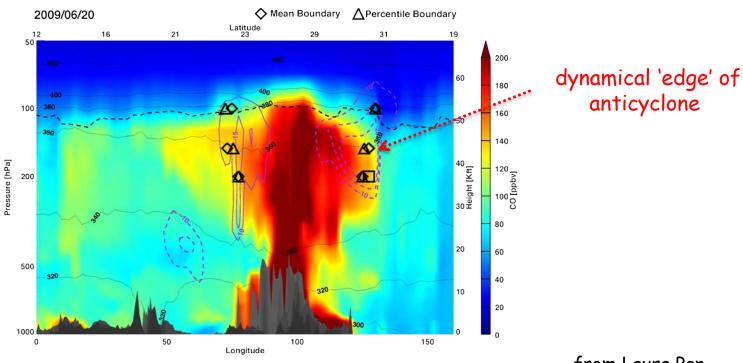
trajectories



MLS 100 hPa carbon monoxide (CO)

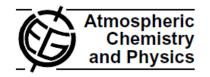


Vertical structure of CO from model simulation



from Laura Pan

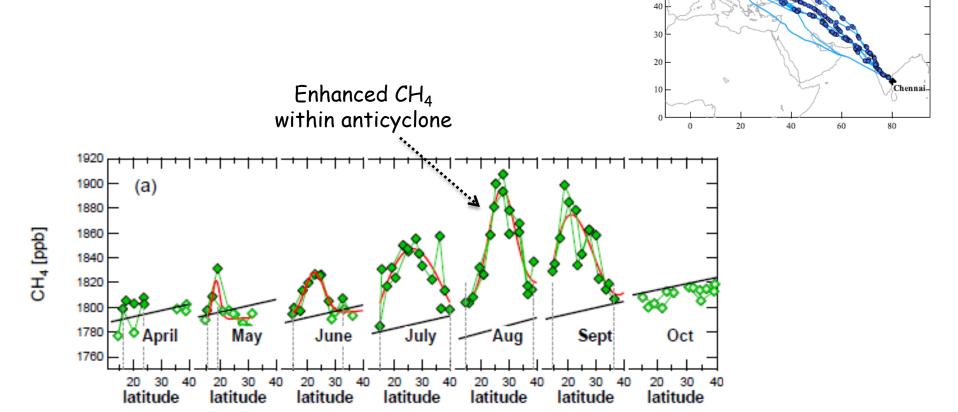
- How sharp is the 'chemical edge'?
- When and where does air 'escape' the anticyclone?



Frankfurt

Greenhouse gas relationships in the Indian summer monsoon plume measured by the CARIBIC passenger aircraft

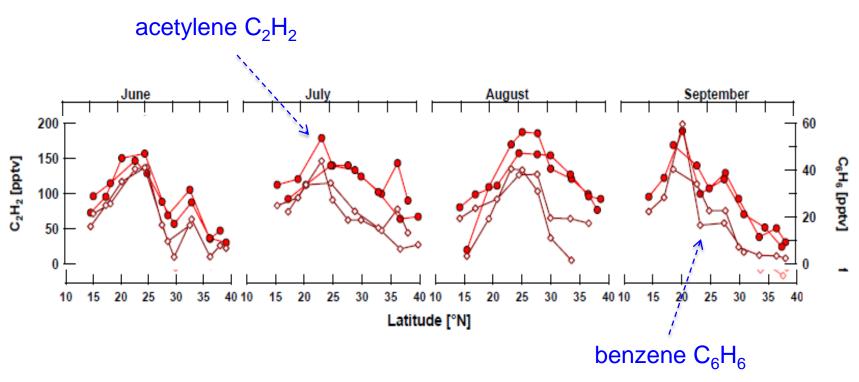
T. J. Schuck¹, C. A. M. Brenninkmeijer¹, A. K. Baker¹, F. Slemr¹, P. F. J. von Velthoven², and A. Zahn³



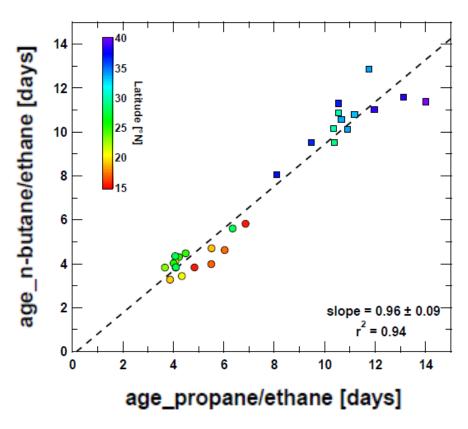
Characterization of non-methane hydrocarbons in Asian summer monsoon outflow observed by the CARIBIC aircraft

A. K. Baker¹, T. J. Schuck¹, F. Slemr¹, P. van Velthoven², A. Zahn³, and C. A. M. Brenninkmeijer¹

ACP, 2011



Age of air estimated from short-lived hydrocarbons

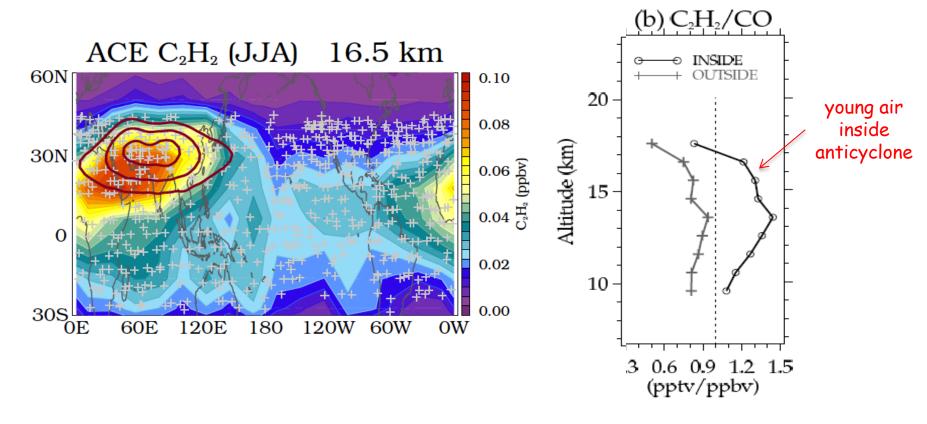


Baker et al, ACP, 2011

Result: air is relatively young: ~5-12 days

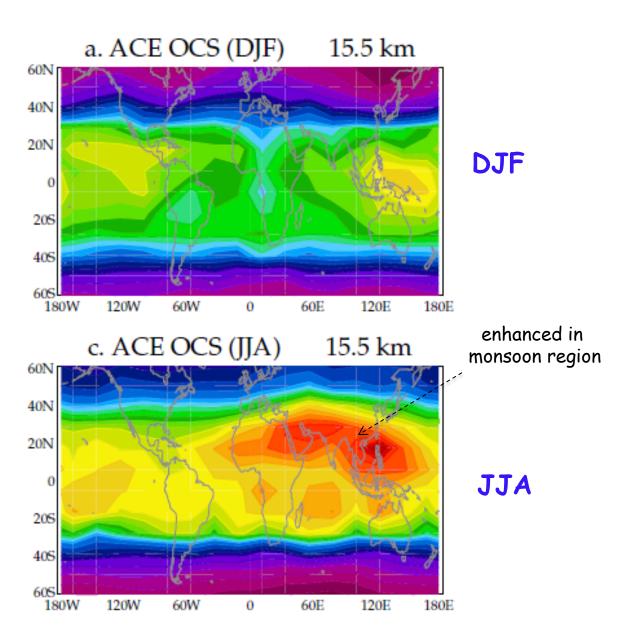
C₂H₂ measurements from ACE-FTS satellite

photochemical lifetime ~ 2 weeks

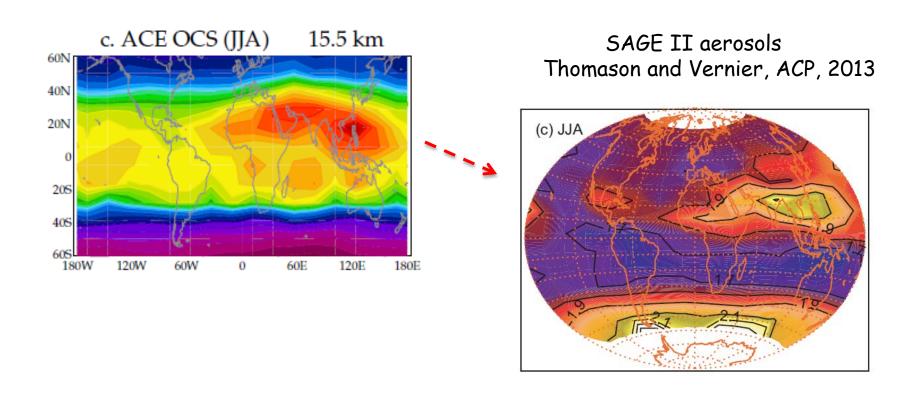


ACE measurements of OCS

Note the large annual cycle



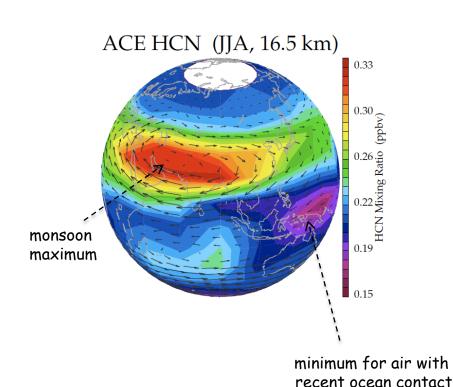
Simple mechanism for ATAL?

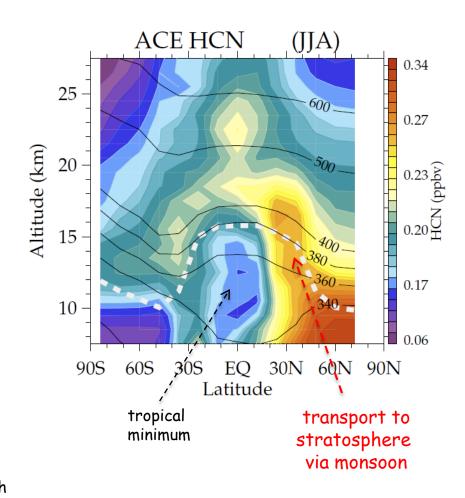


Transport to the stratosphere observed from ACE hydrogen cyanide (HCN)

HCN - biomass burning tracer

- Minimum in tropics (ocean sink)
- Long lived in free atmosphere

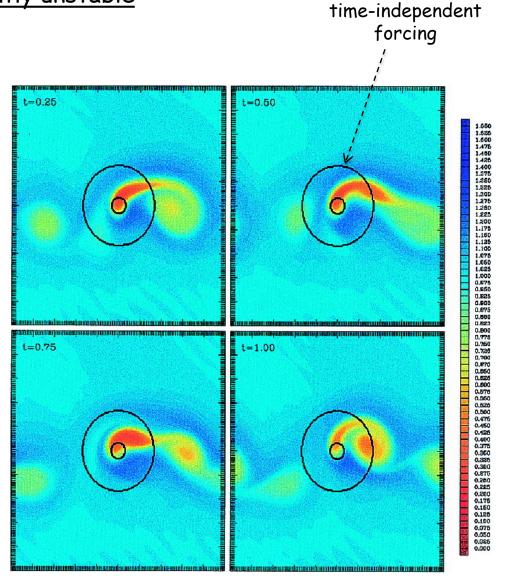




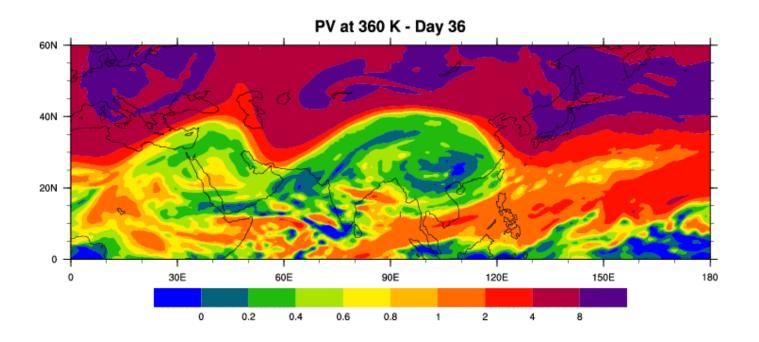
Monsoon circulation is inherently unstable

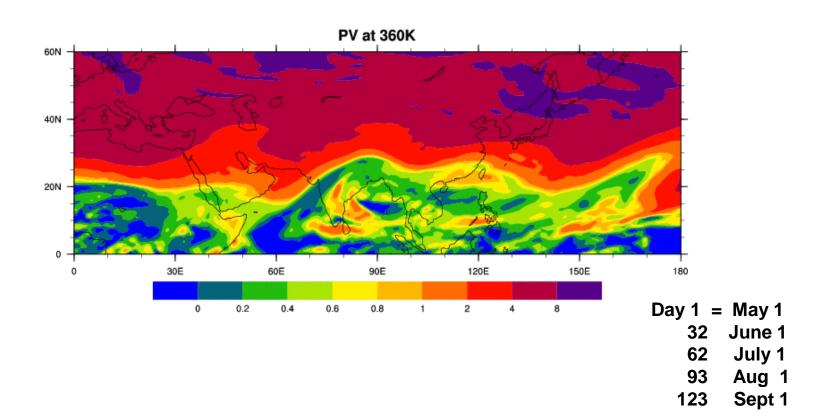
Hsu and Plumb 2000

eddy shedding from monsoon circulation



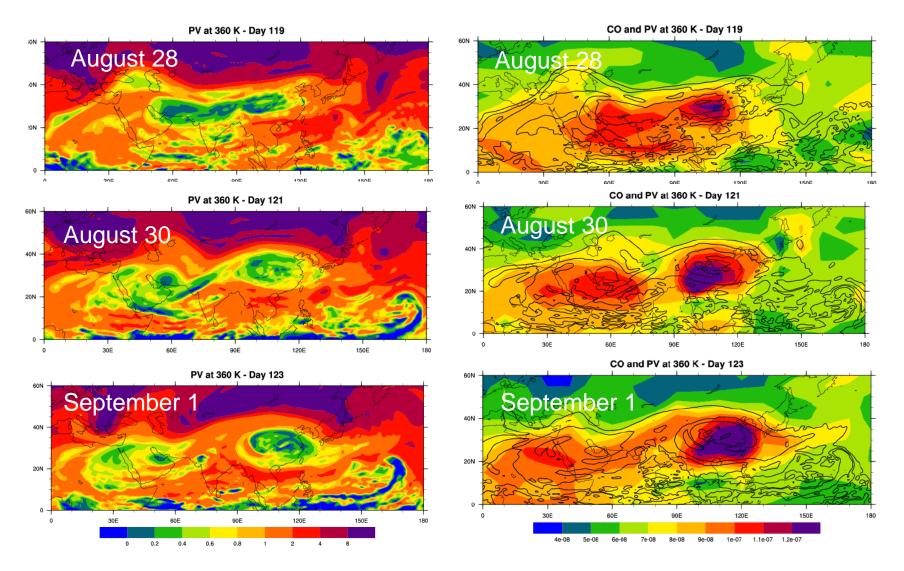
Anticyclone viewed in potential vorticity





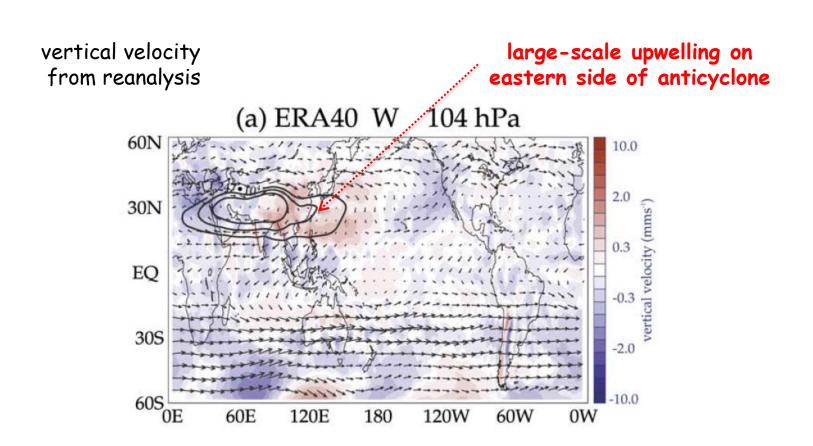
PV maps at 360 K

CO from Aura MLS

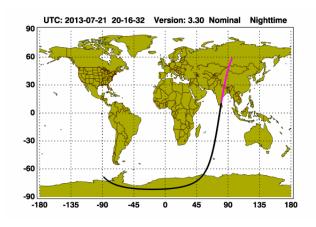


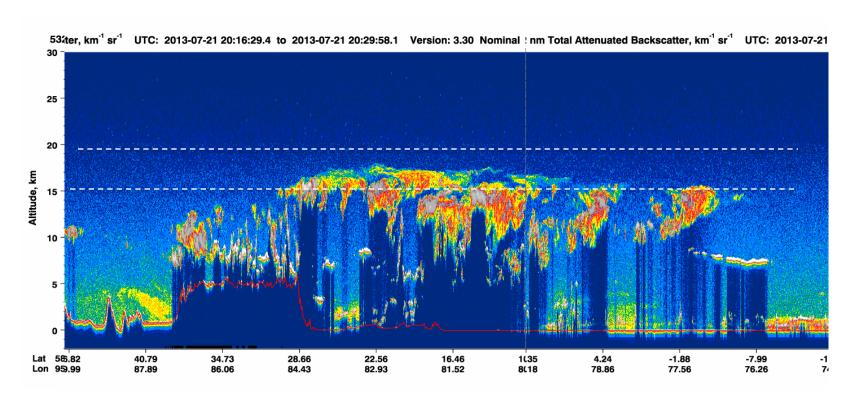
What controls transport to stratosphere?

Deep convection and/or large-scale circulation?



Deep convective transport near/above tropopause



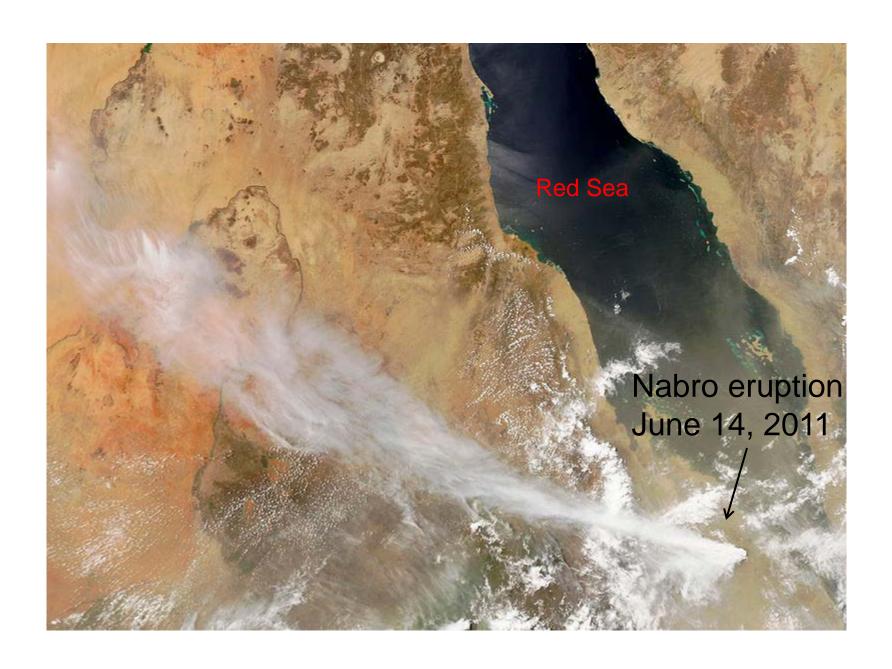


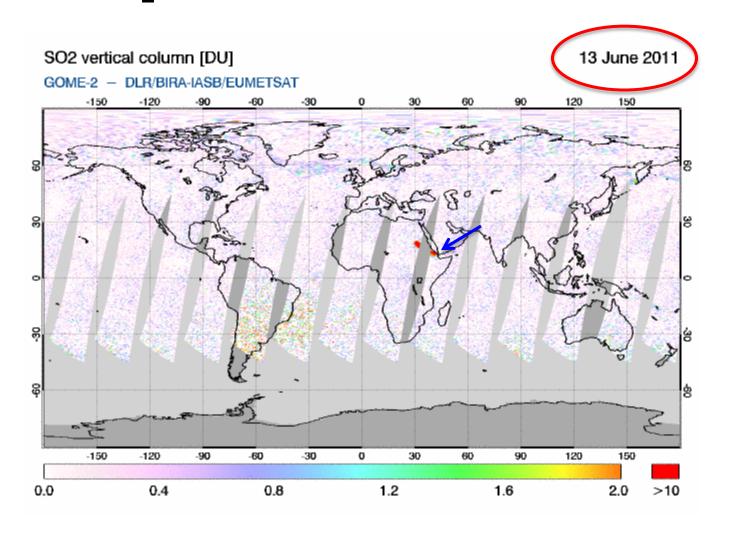
Eruption of Mt. Nabro

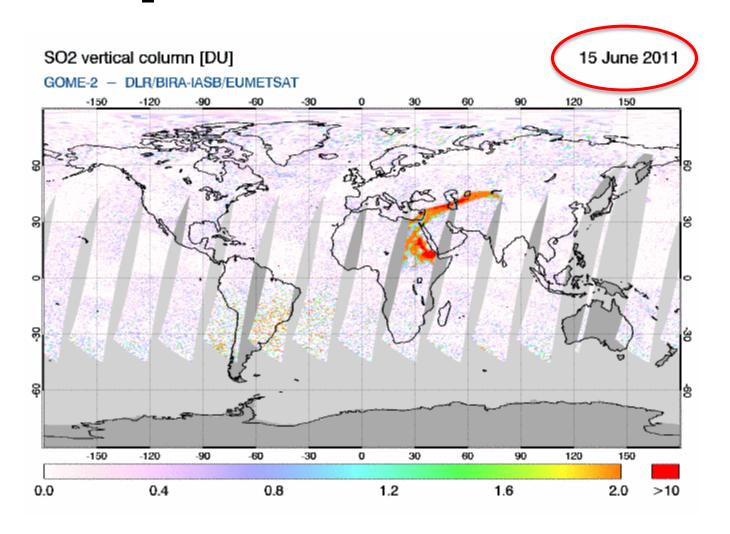
June 13, 2011

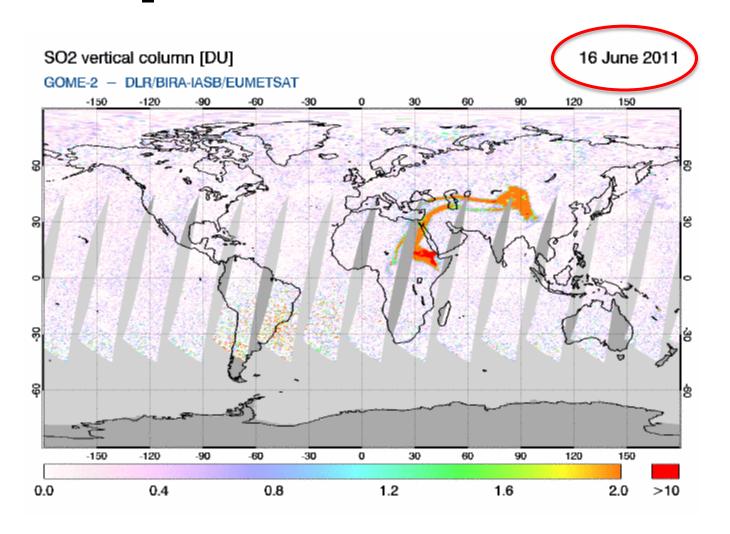
Eritria, Africa

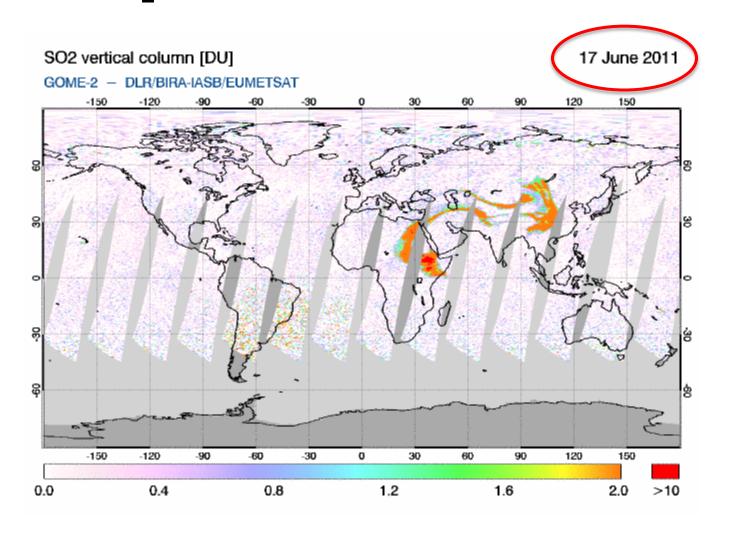




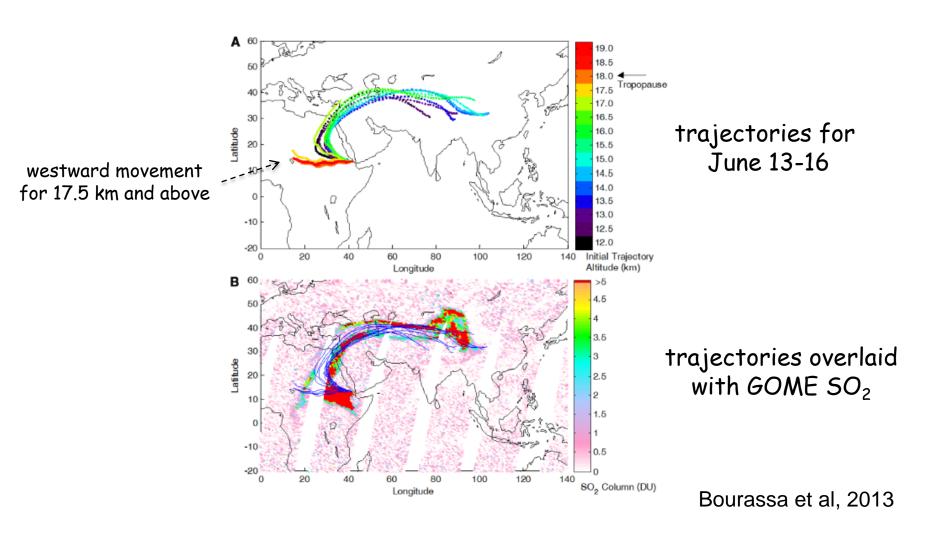






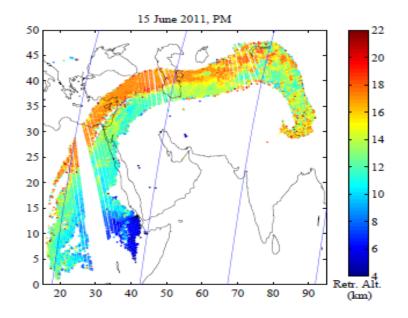


How high did the eruption go?

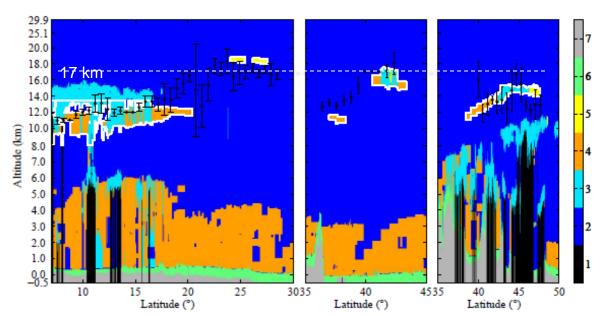


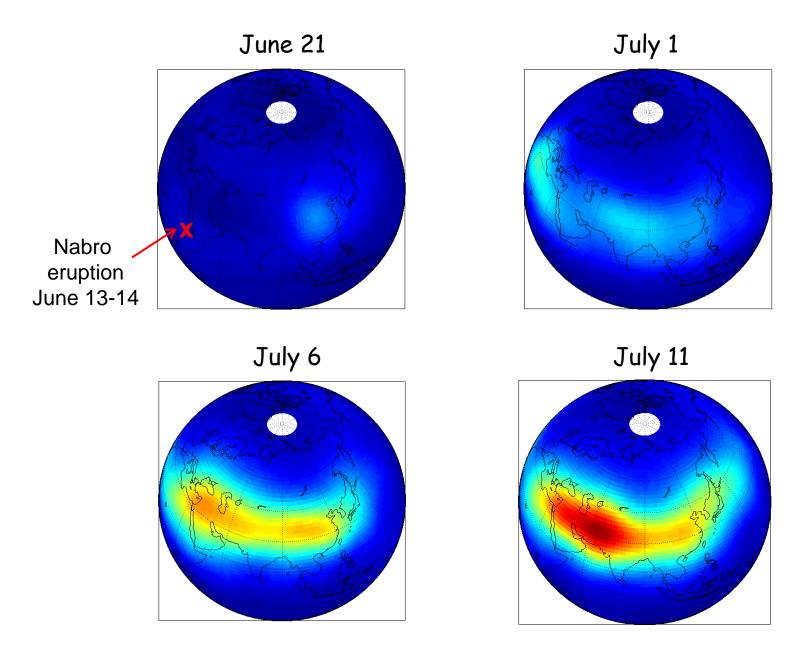
Nabro SO₂ plume height derived from IASI

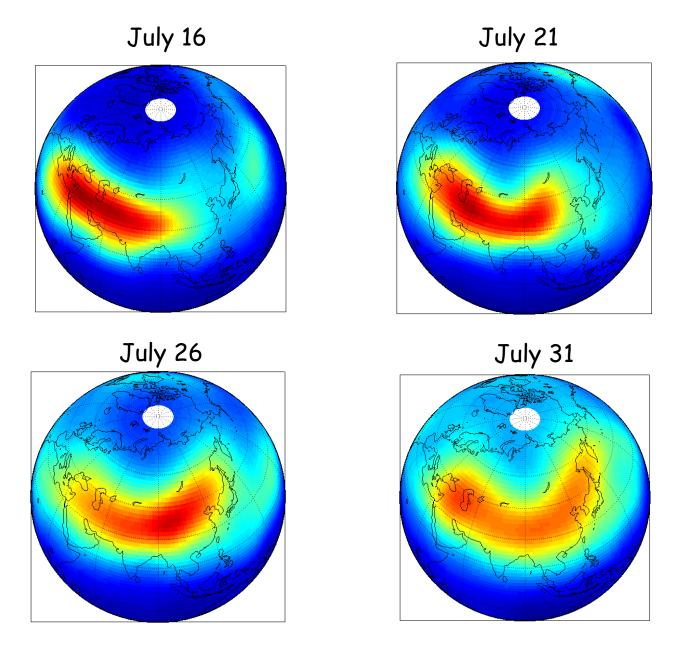
Clarisse et al, submitted



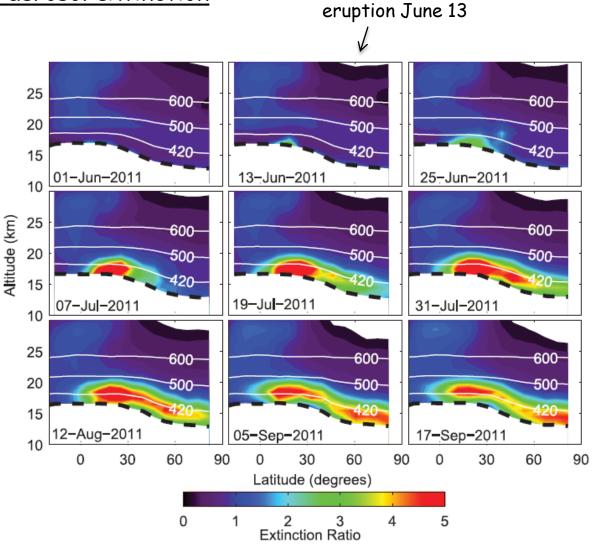
primary injection to ~12-17 km





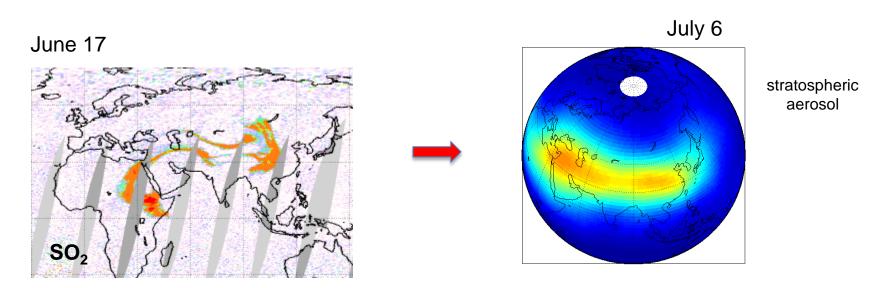


OSIRIS aerosol extinction

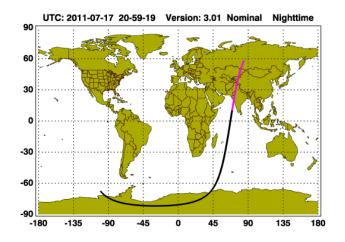


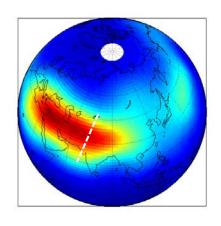
Interpretation:

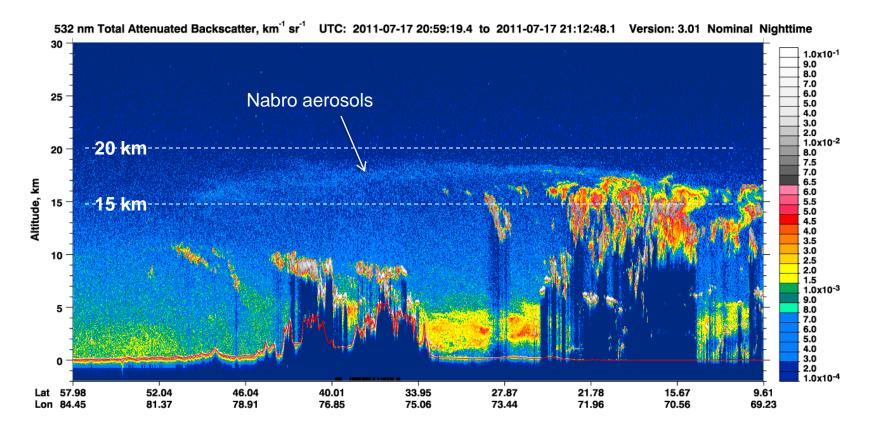
- Nabro SO₂ plume in upper troposphere, transported around monsoon circulation to eastern side.
- Transport to stratosphere through circulation / convection (?)
- Confined to anticyclone, converted to stratospheric sulfate aerosol ~ 1 month
- Further evidence of transport to lower stratosphere via monsoon (Nabro in right place at right time)



July 17 34 days after eruption







Ongoing questions:

- What are the contributions of different chemical source regions to the upper troposphere? Is reactive chemistry important? How much reactive nitrogen is in the anticyclone?
- When and where does air escape the anticyclone? Are there sharp gradients across edges?
- What is the role of deep convection vs. large-scale upward circulation to the stratosphere? How important are diurnal variations in convection?
- What is the nature of the tropopause aerosol layer? Does it influence UTLS clouds?

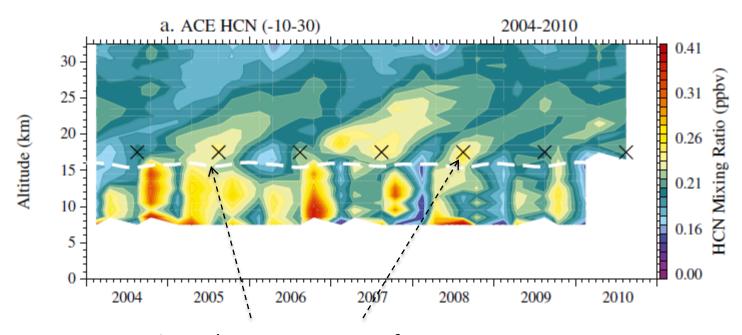


Extra slides

HCN 'tape recorder' from ACE-FTS measurements

PARK ET AL.: HYDROCARBONS FROM ACE-FTS AND WACCM4

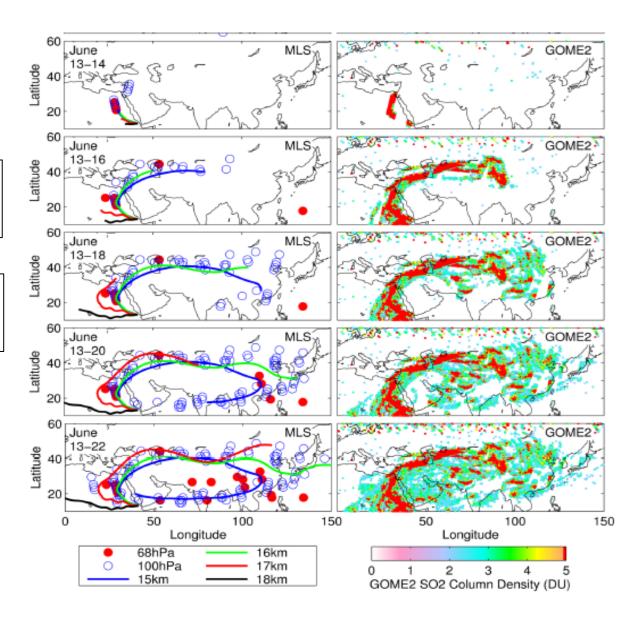
JGR, 2013



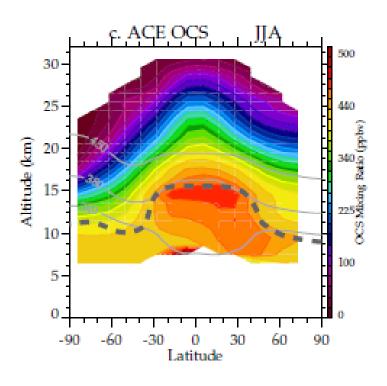
boreal summer maxima from Asian monsoon circulation

blue dots: enhanced at 100 hPa

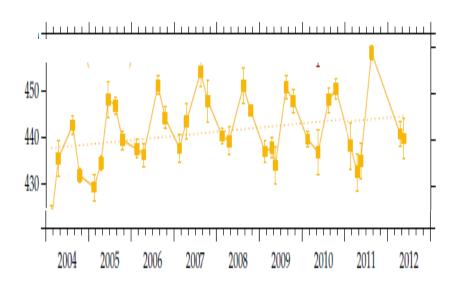
red dots: enhanced at 68 hPa



OCS climatology from ACE-FTS

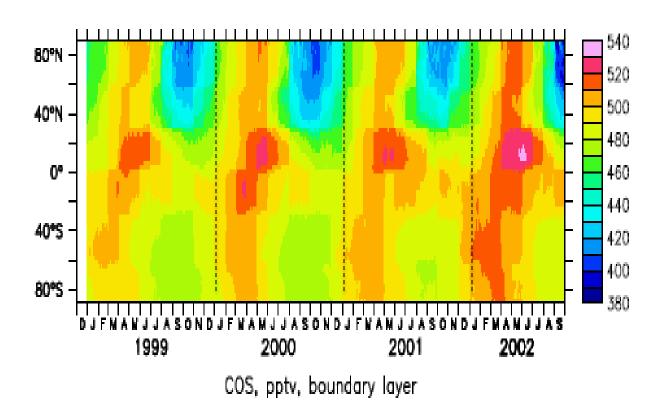


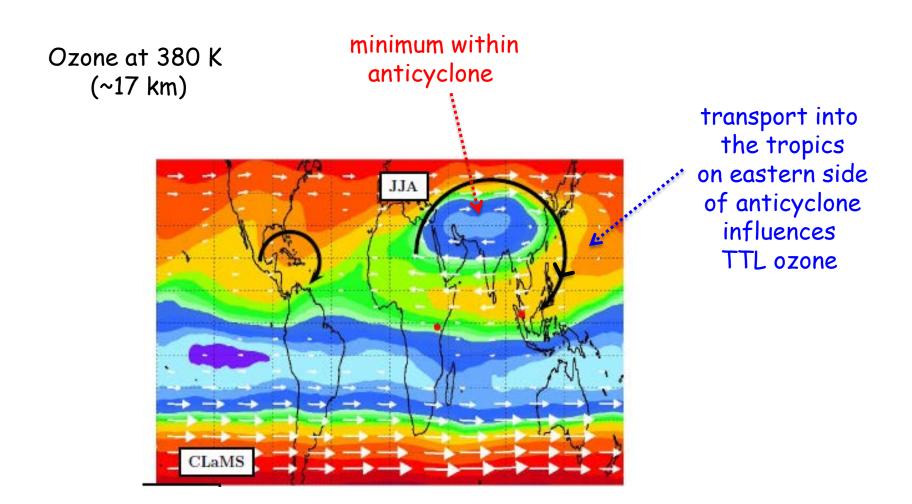
OCS at 16.5 km 30° N-S



increase of ~2% / decade over 2004-2012

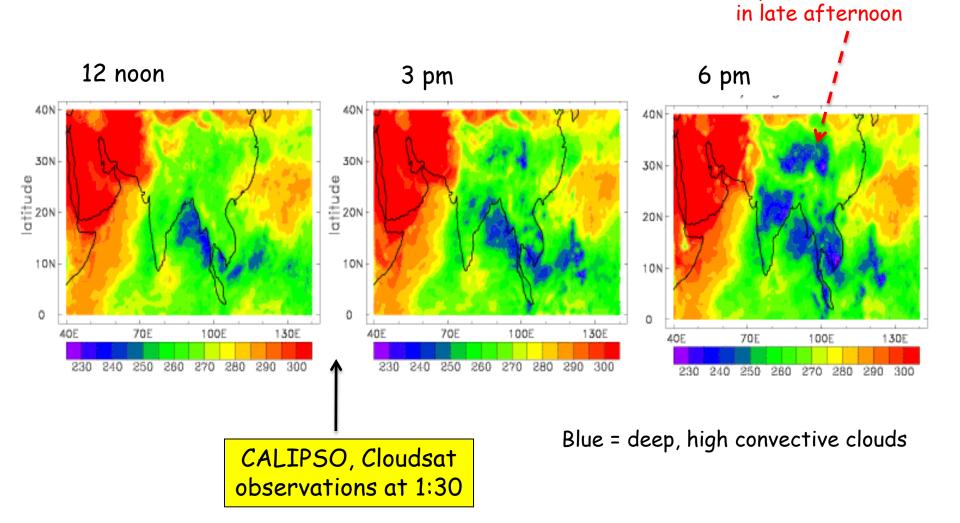
Boundary layer OCS Bruhl et al, 2012,





Konopka et al, 2009, 2010; Ploeger et al, 2012

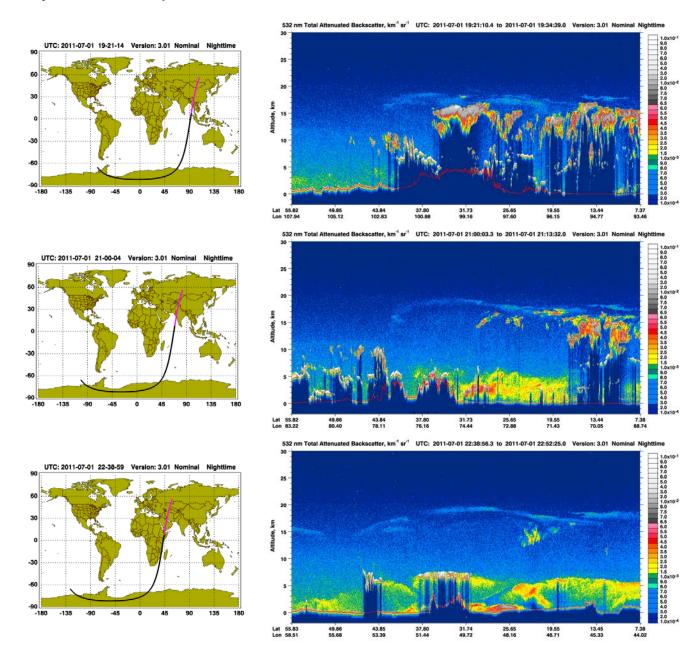
convective cloud statistics from 3-hour geostationary (CLAUS) data



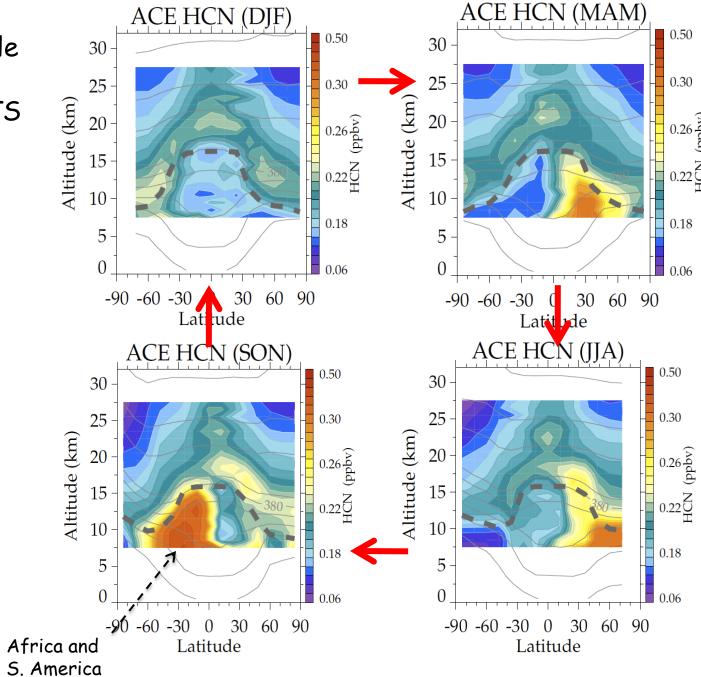
(Motivated from Jonathan Wright, Rong Fu)

Tibet plateau convection

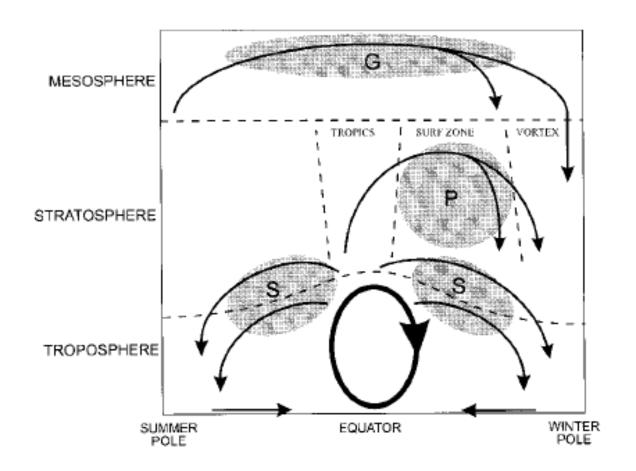
July 1: 18 days after eruption

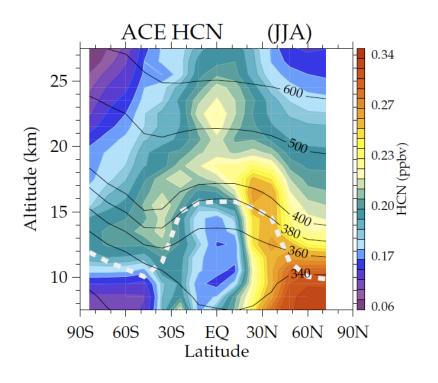


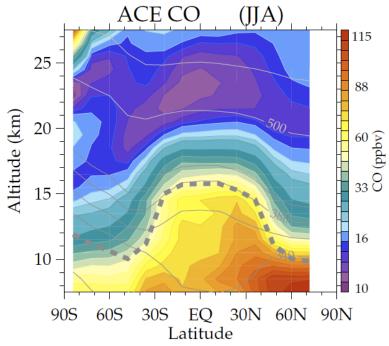
Seasonal cycle of HCN from ACE-FTS



Plumb (2002) 'Stratospheric Transport'



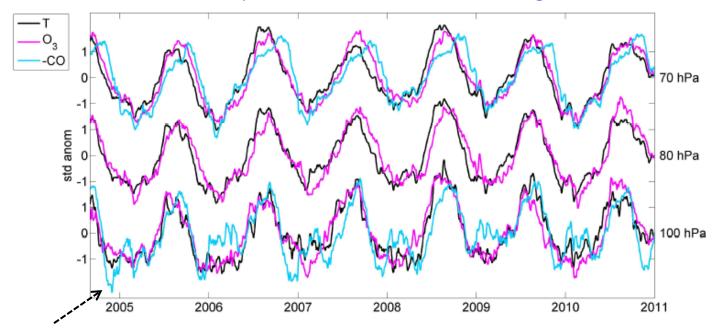




Variability in tropical upwelling drives temperatures and ozone and other tracers in tropical lower stratosphere

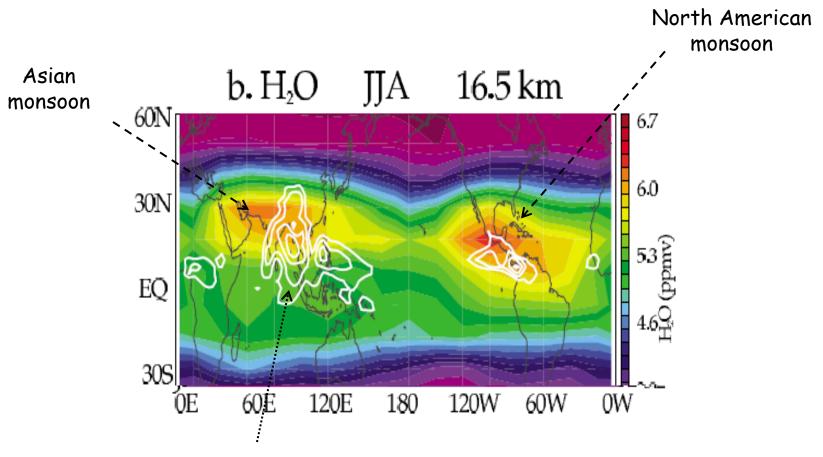
- * large annual cycle
- * sub-seasonal (week-to-week) variability

Zonal mean temperature, ozone and CO averaged 18° N-S



CO inverted scale

Also enhanced water vapor in monsoon regions



white contours: deep convection

PV maps at 360 K

CO from Aura MLS

