



The vertical distribution of volcanic SO2 plumes measured by IASI

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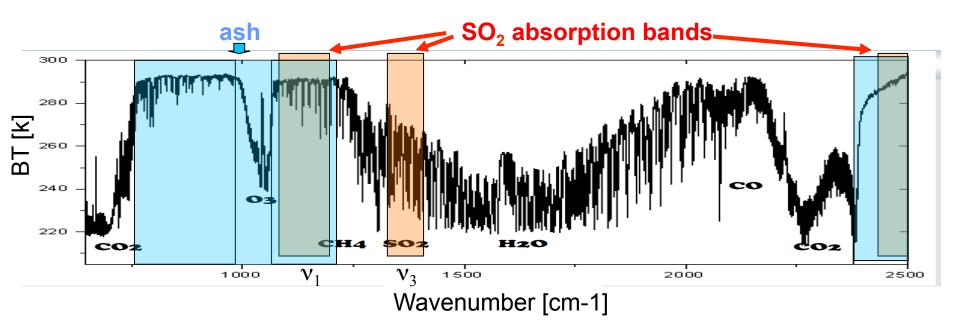
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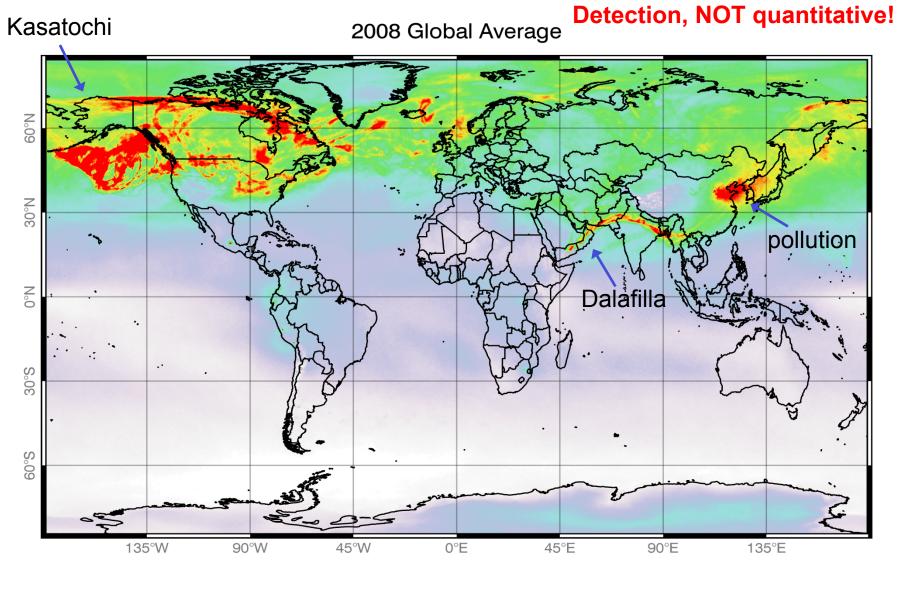
Infrared Atmospheric Sounding Interferometer - IASI

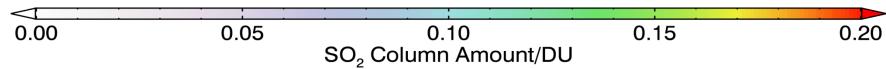
IASI is on board of METeorological OPerational satellite program (METOP-A and METOP-B), a European meteorological satellite that has been operational since 2007.

IASI is a Fourier transform spectrometer, that measures the **spectral range 645 to 2760 cm⁻¹** (3.62–15.5µm) with a spectral sampling of 0.25 cm⁻¹ and an apodised spectral resolution of 0.5 cm⁻¹. Radiometric accuracy is 0.25-0.58K. The IASI field of view (FOV) consists of four circles of 12 km diameter (at nadir) inside a square of 50 x 50 km.

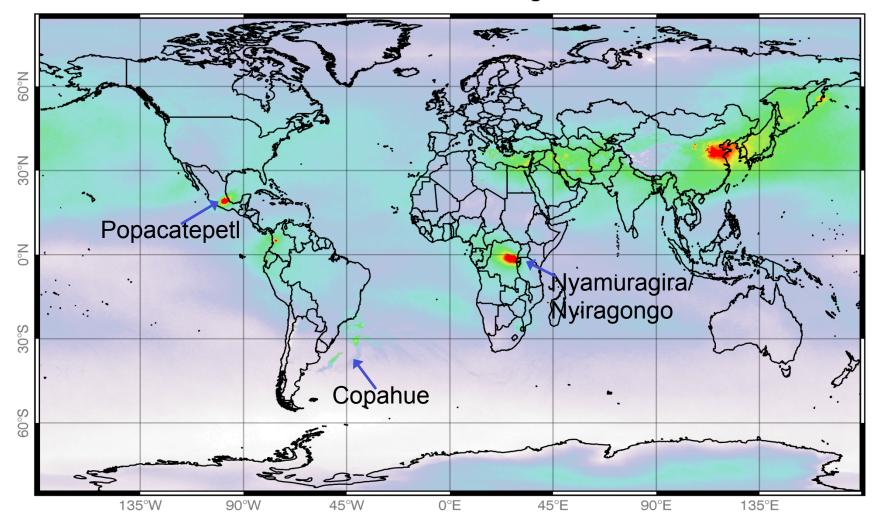
It has a 2000 km swath and nominally can achieved **global coverage in 12 hours** (although there are some gaps between orbits at tropical latitudes). Radiances are collocated with the Advanced Very High Resolution Radiometer (AVHRR) that provides complementary visible/near infrared channel, for cloud and aerosol retrievals.

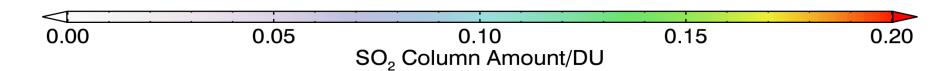






2012 Global Average





SO₂ Retrieval scheme



- Total column amount of SO₂
- Altitude H
- Thickness s
- Surface temperature Ts

+ ECMWF profile (temperature, h2o, p, z)



Forward model: fast radiative transfer (RTTOV + SO2 RAL coefficients)

IASI simulated spectra

y is the measurement vector, x the state vector **F(x)** forward model, Sy error covariance matrix

IASI measurements

Ts

SI



Н

OE retrieval

$$J = (y - F(x))^{T} S_{y}^{-1} (y - F(x)) + (x - x_{a})^{T} S_{a}^{-1} (x - x_{a})$$

best estimate of stare vector:

SO₂ amount, plume altitude, Ts

$$S_y(i,j) = \langle (y_{mi} - y_{si}) - (\overline{y_{mi}} - y_{si}) \rangle \langle (y_{mj} - y_{sj}) - \overline{(y_{mj}} - y_{sj}) \rangle$$

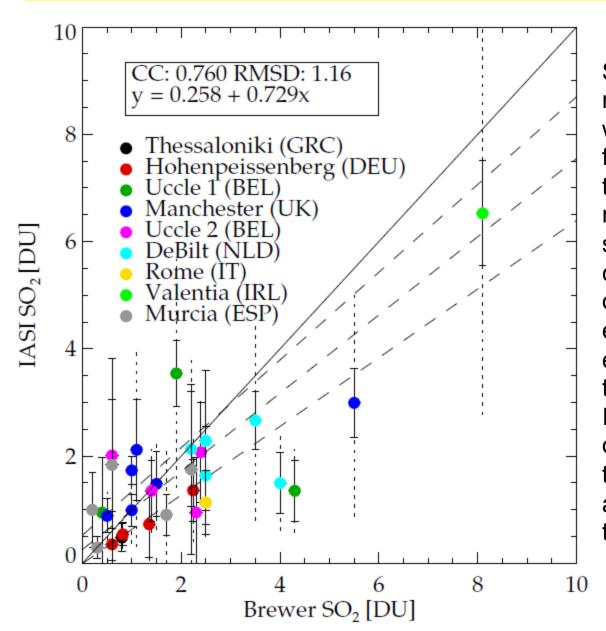
$$y_s = F(SO_2=0)$$

Sy Computed with billions pixels

 S_y , is defined to represent the effects of atmospheric variability not represented in the forward model (FM), as well as instrument noise (cloud and trace-gases...).

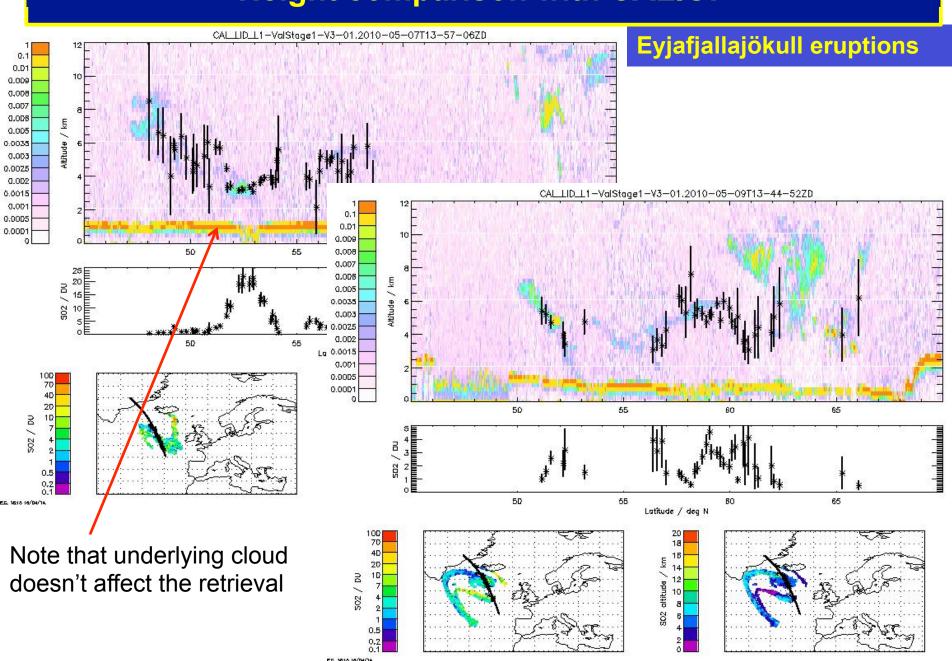
The matrix is constructed from differences between FM calculations (for clear-sky) and actual IASI observations for wide range of conditions, when we are confident that negligible amounts of SO₂ are present.

Comparison with Brewer ground data

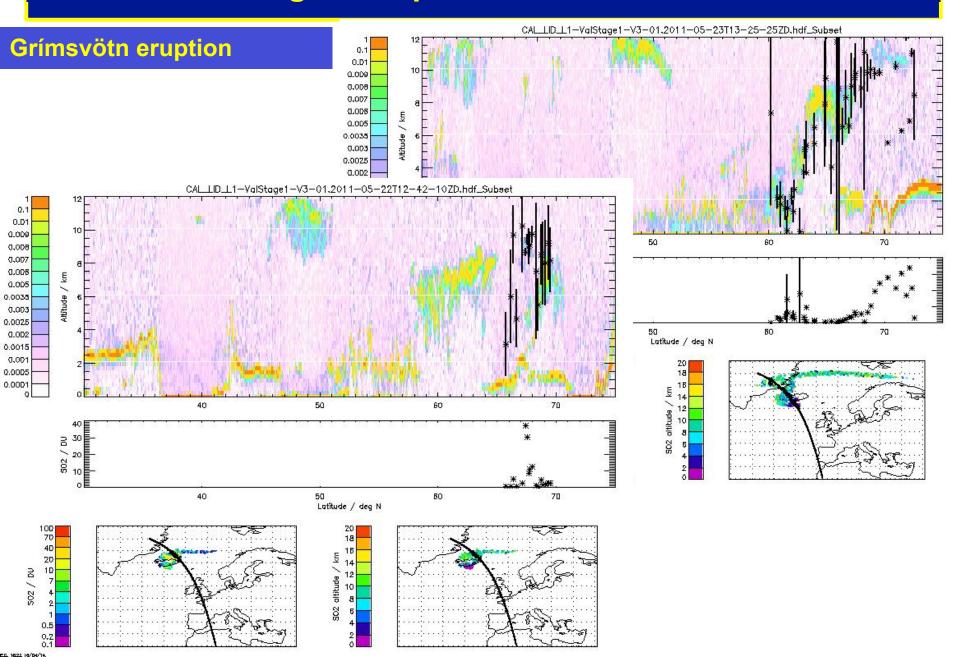


Scatter plot of IASI SO2 measurements, averaged within a distance of 200 km from the ground station, versus the daily SO2 column amount, measured from Brewer spectrometers. Different colours correspond to a different ground station. Black error-bars are the IASI average errors; dotted error-bars are the standard deviation of the IASI data within the selected distance. Black lines represent the ideal line y=x; dotted lines are the best fits with error in the best fit

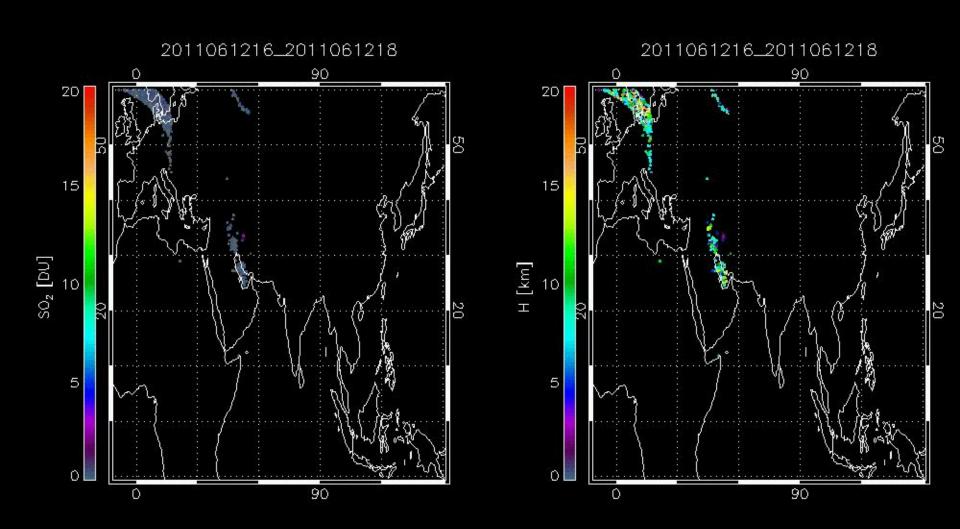
Height comparison with CALIOP

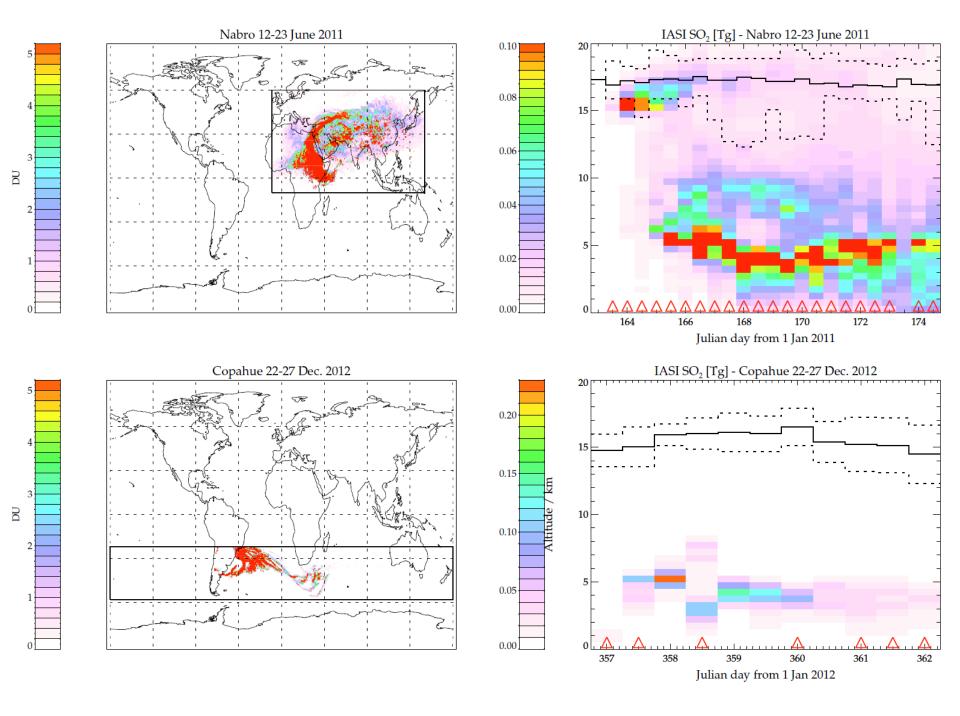


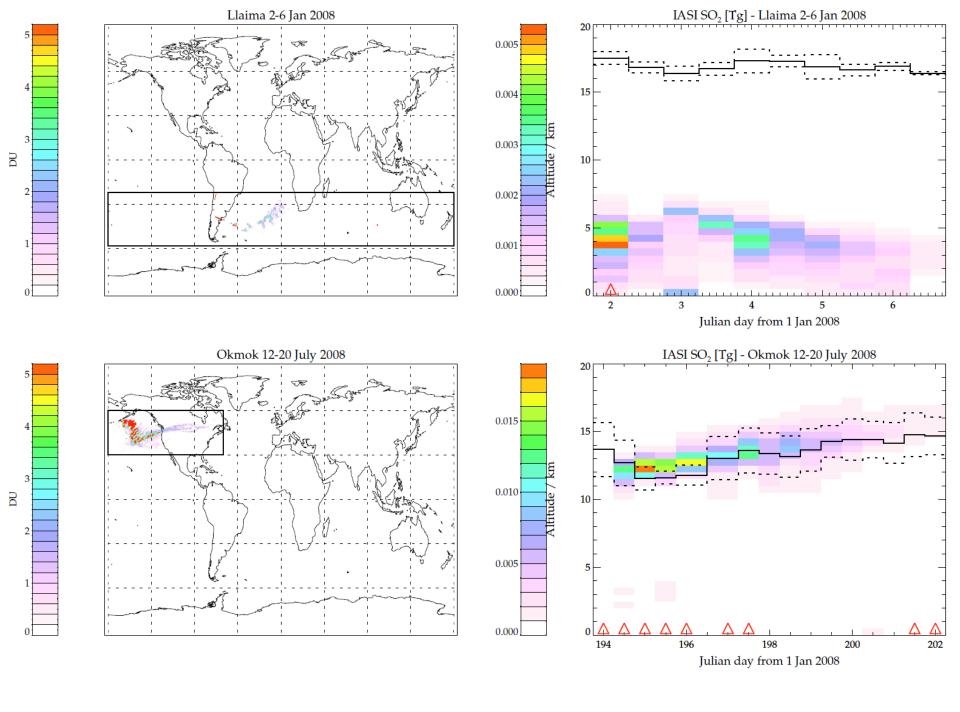
Height comparison with CALIOP

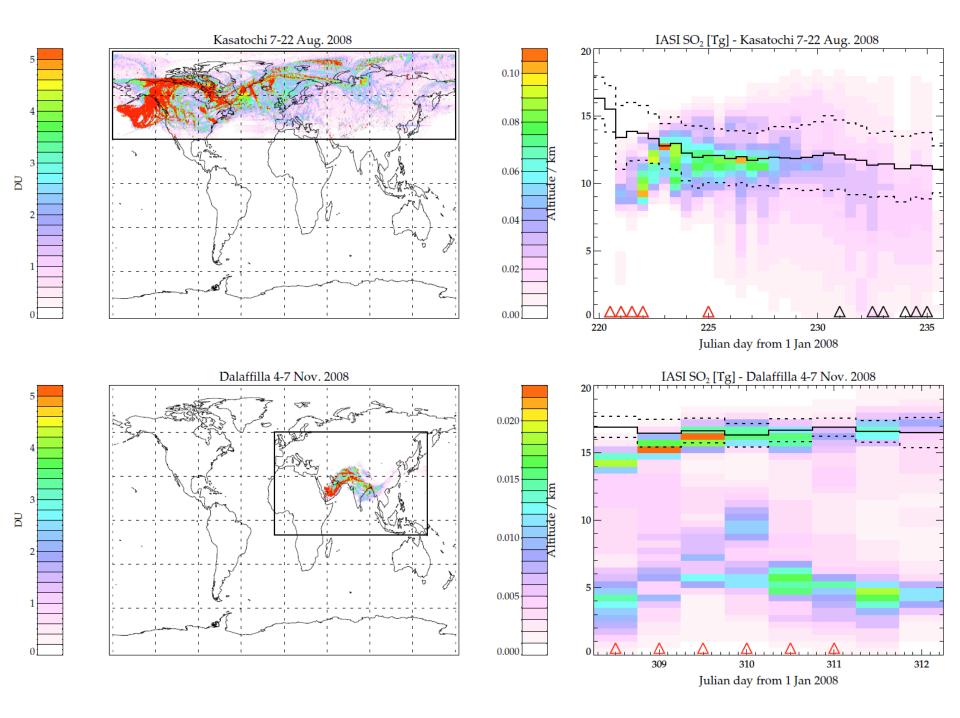


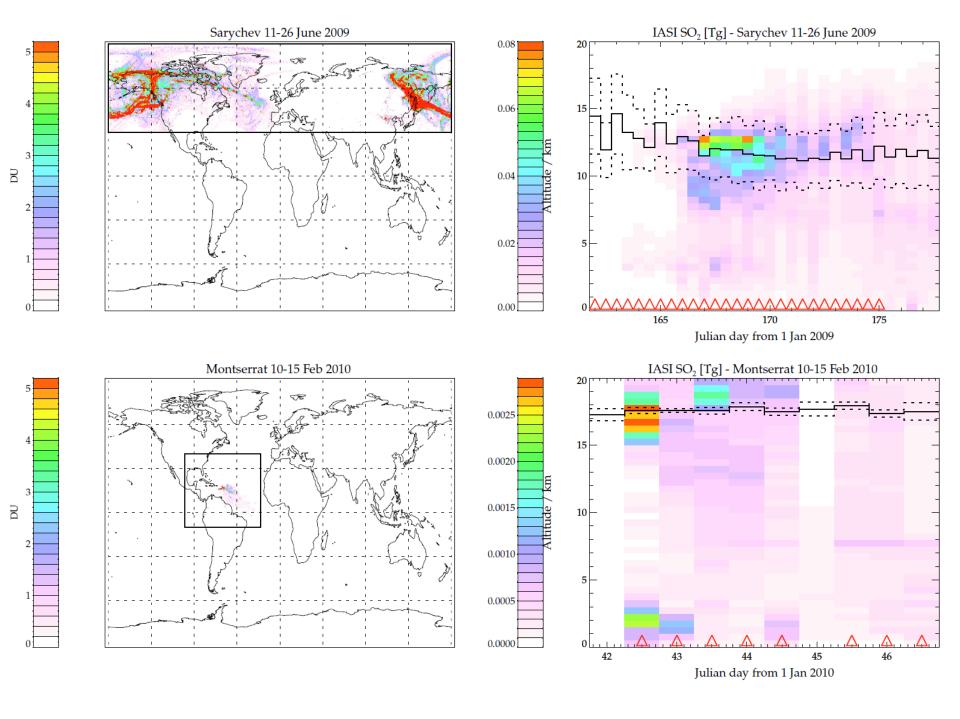
Nabro eruption 13-24 June 2011

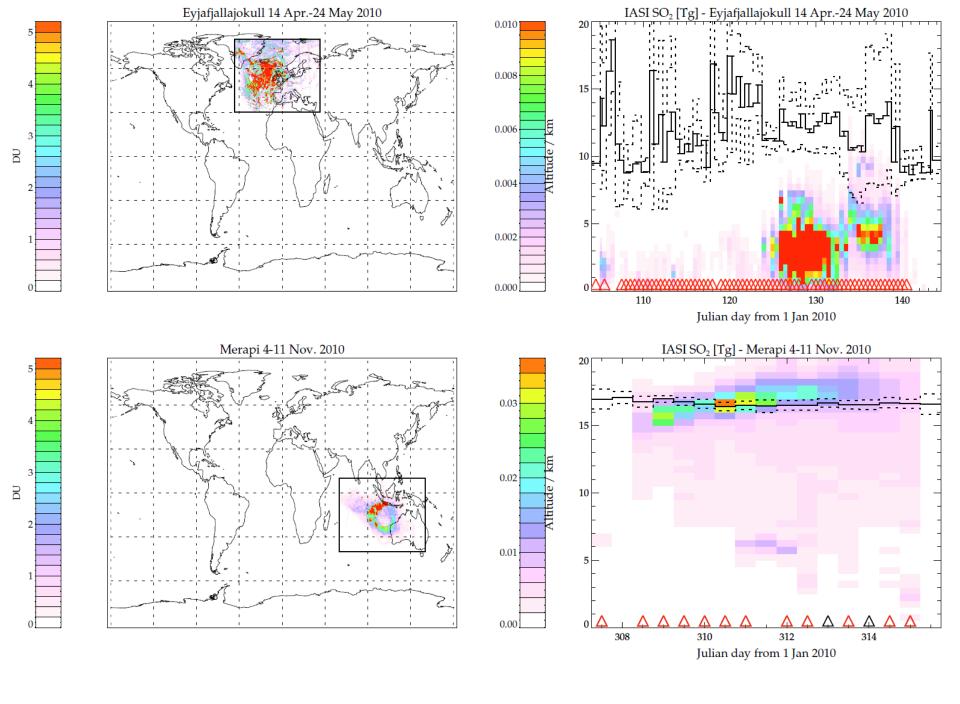


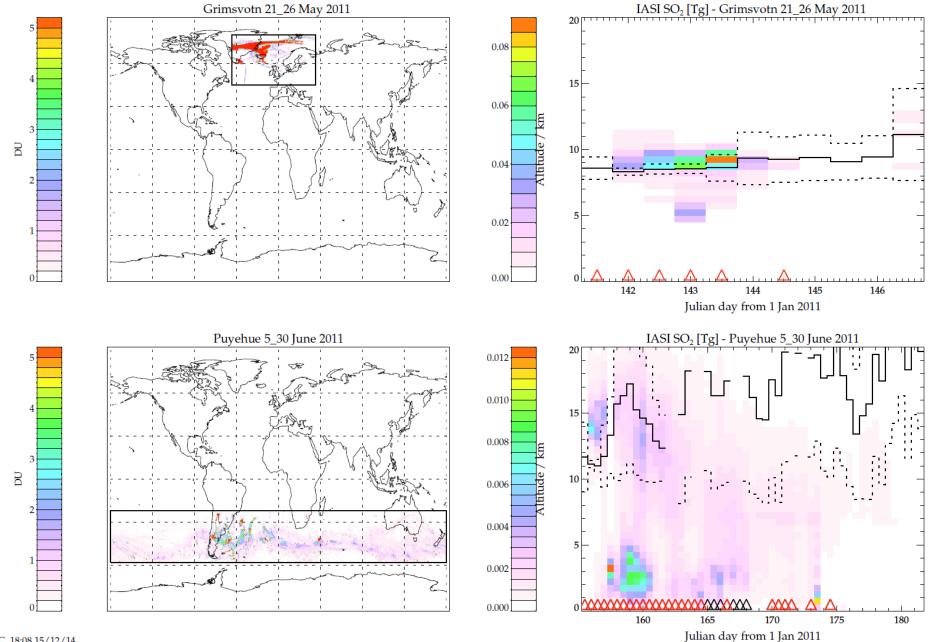




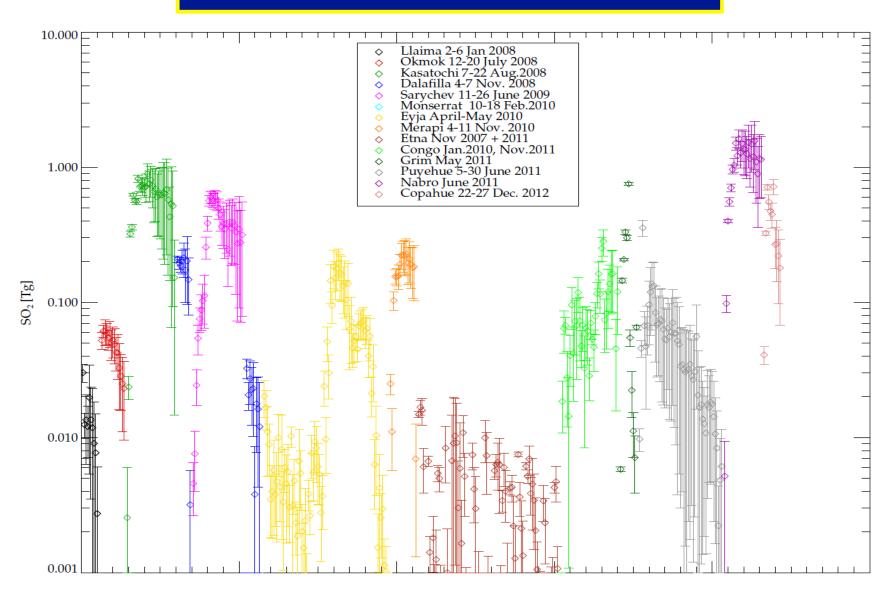




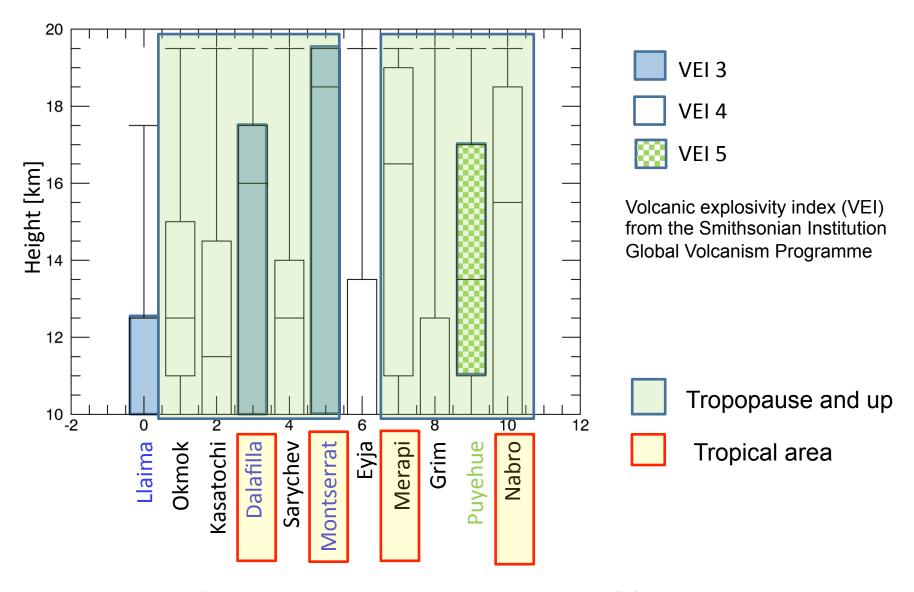




SO2 total mass

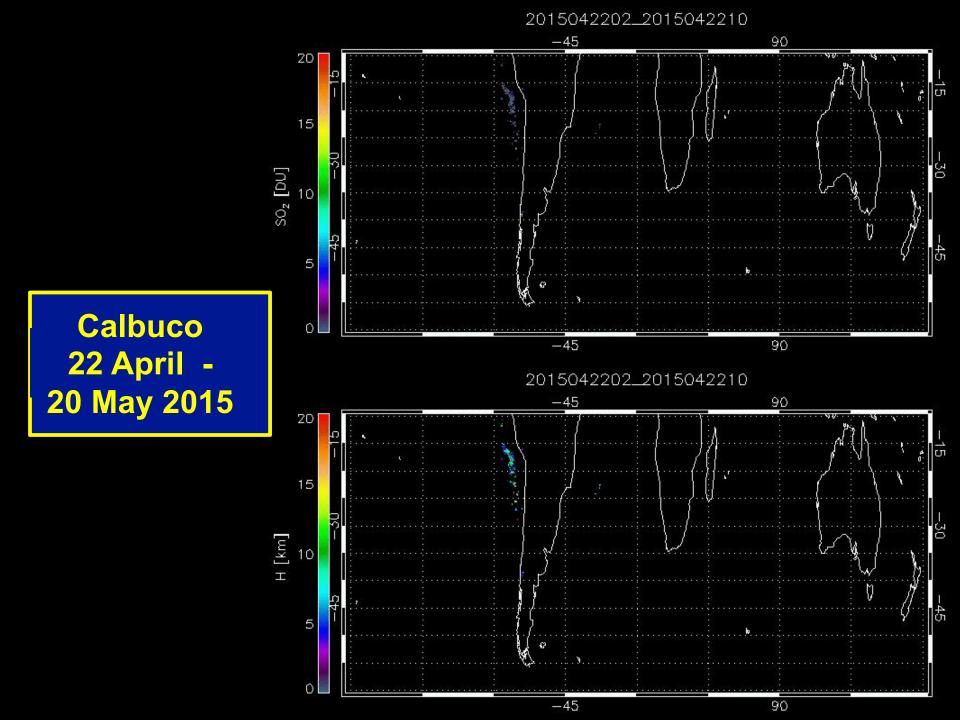


 SO_2 retrieved from IASI data. The values are the measured amount on a particular day and vary with volcanic emission, gas removal and satellite sampling. Points are separated by ~12 hours.



VEI is a poor index of the potential height to which volcanic SO2 is injected.

All of the eruptions in the tropics (except Nyamuragira, VEI 1,2), reached the tropopause.



Summary

SO2 detection: (AMT Walker et al 2011, JRL Walker et al 2012)

Very fast => global survey tool

- ⇒ show emission from volcanic eruptions, anthropogenic source and degassing.
 - IASI archive 2007-2014
 - NRT processing

SO2 iterative: (ACP Carboni et al 2012, ACP Carboni et al. 2016)

We use simultaneously cannel between 1000-1200 cm-1 and 1300-1410 cm-1 (v1 and v3 SO2 absorption band)

- retrieve both column amount and altitude for volcanic plume.
- IASI retrieved values are consistent with CALIPSO altitudes and Brewers column amount
- we can study the plume vertical distributions, and evolution in time.
- There is a tendency for volcanic SO2 plumes to reach the tropopause. (All of the eruptions in the tropics (except Nyamuragira) reached the tropopause).

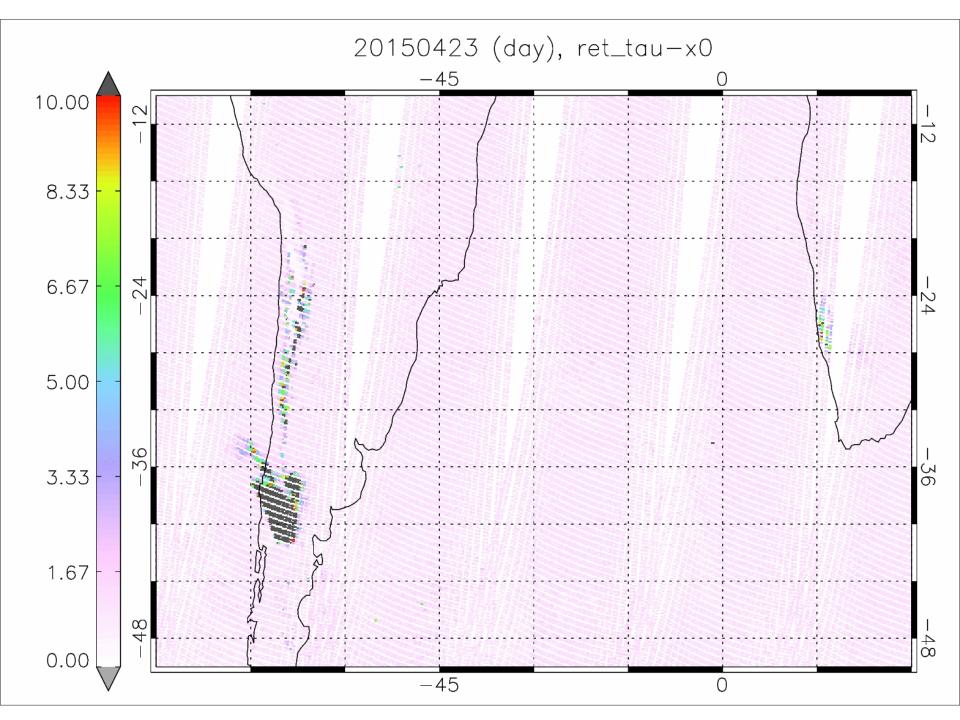
More papers using this IASI SO2: Schmidt et al. 2015 - Bardabunga

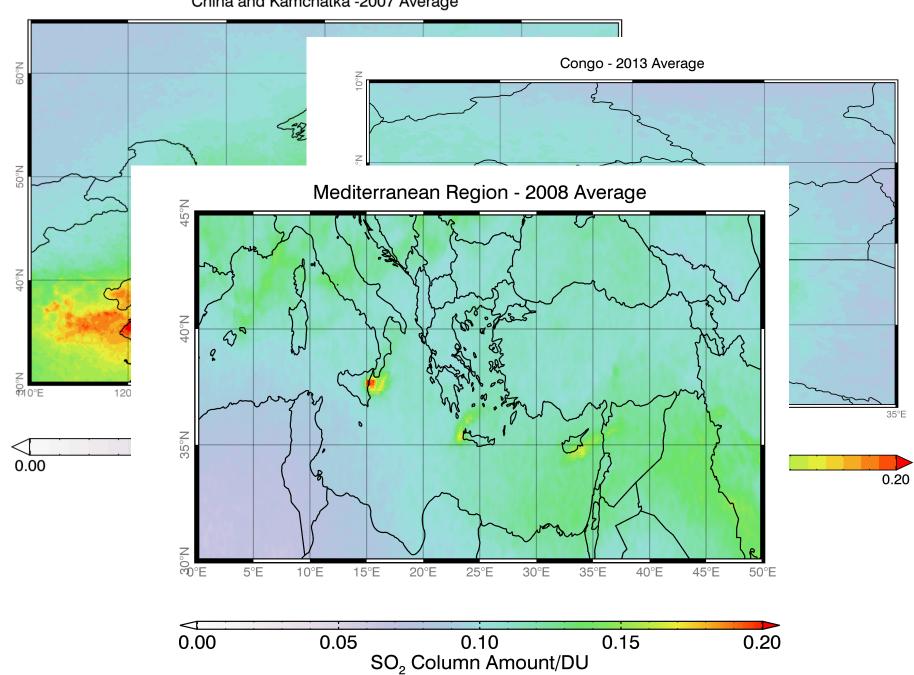
Dataset is available (elisa@atm.ox.ac.uk)

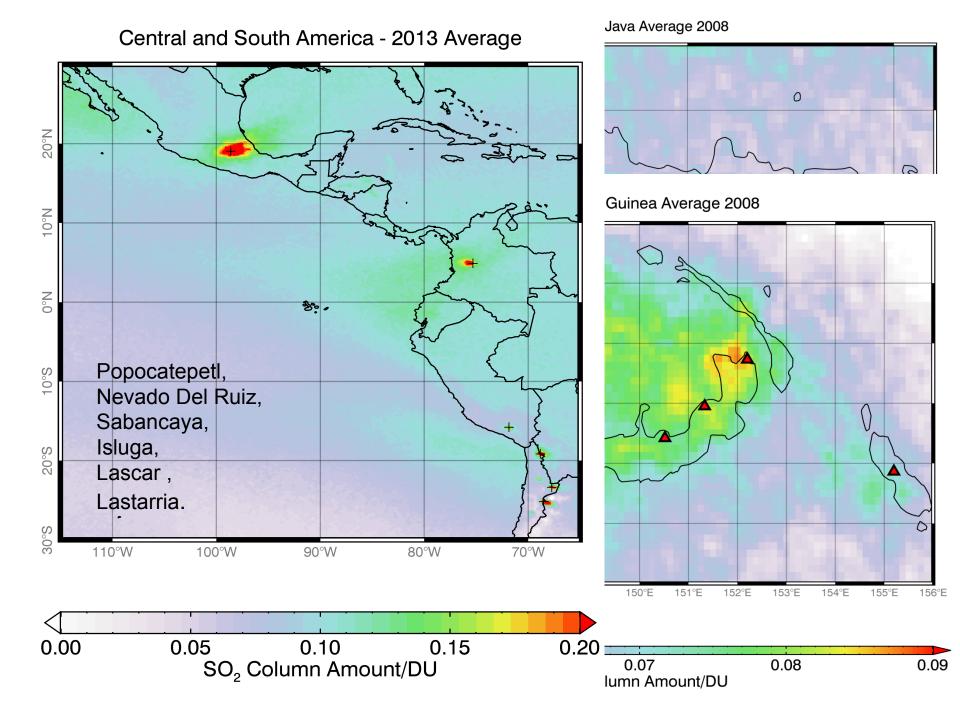
Fromm et al 2014 – Nabro

Koukouli et al 2015, Spinetti et al 2015, Sears et al. 2013

Thank you!



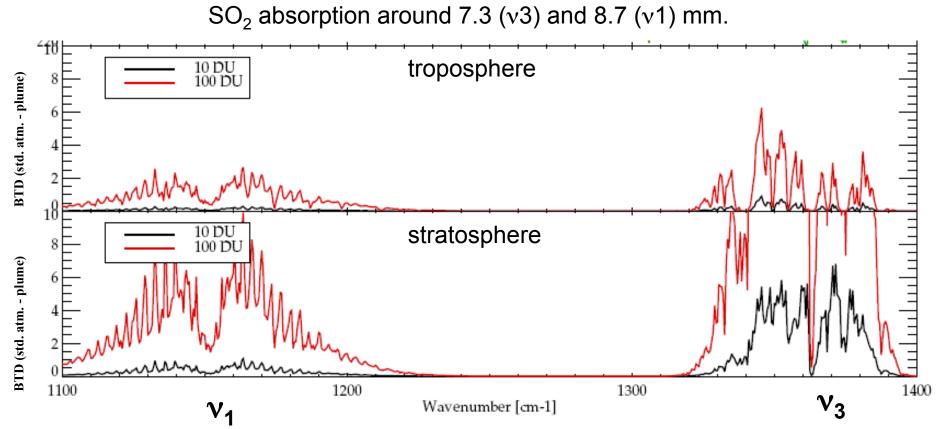




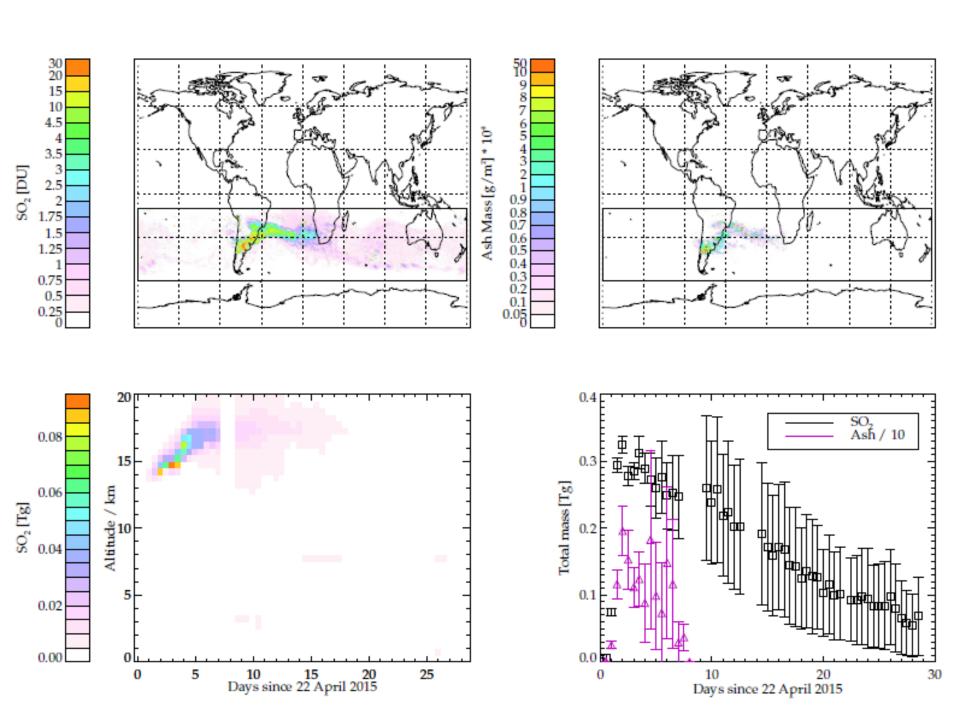


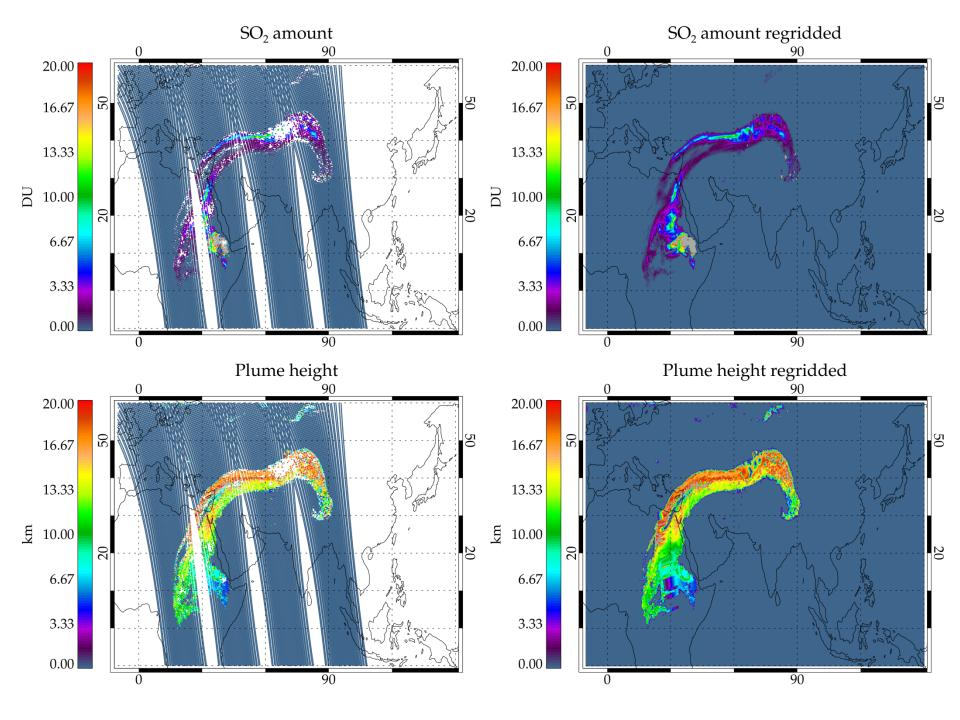
SO₂ thermal infrared spectra

Brightness temperature differences (BTD) between the clear atmosphere and the atmosphere with enhanced tropospheric/stratospheric (top/bottom) SO_2 containing a total column amount of 10DU (black line) or 100DU (red line) of SO_2 ;



IASI is sensitive to both the amount of SO_2 and the altitude of the plume => getting the altitude correct is important in order to get the correct *amount* of SO_2 , since the signal depends strongly on altitude.





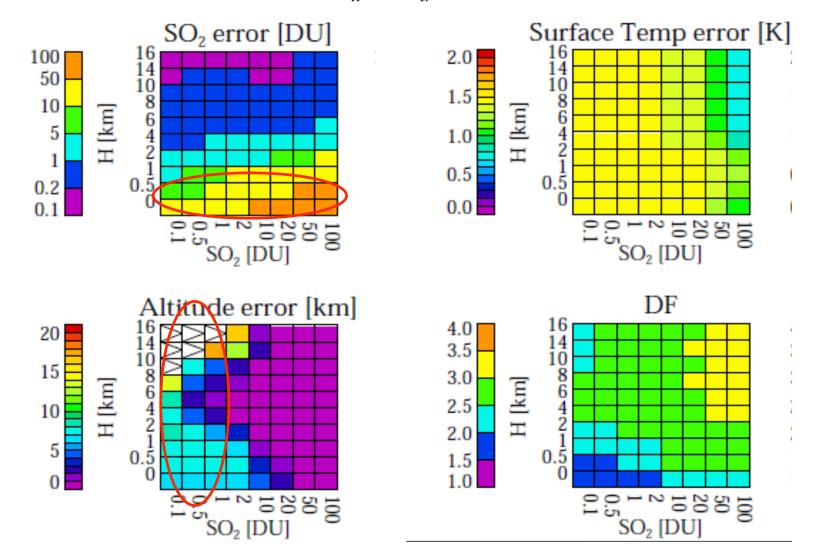
Error analysis

state obtained as:

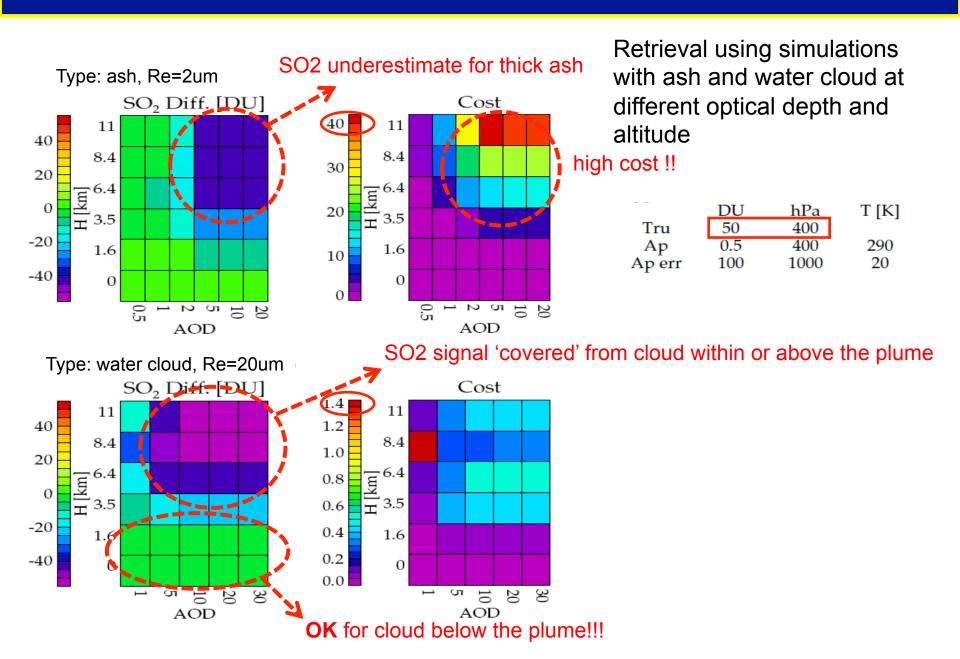
linear error on the 'true'
$$S_x = (K^T S_y^{-1} K + S_a^{-1})^{-1}$$
 state obtained as:
$$A = S_x S_x^{-1} - S_a^{-1}$$

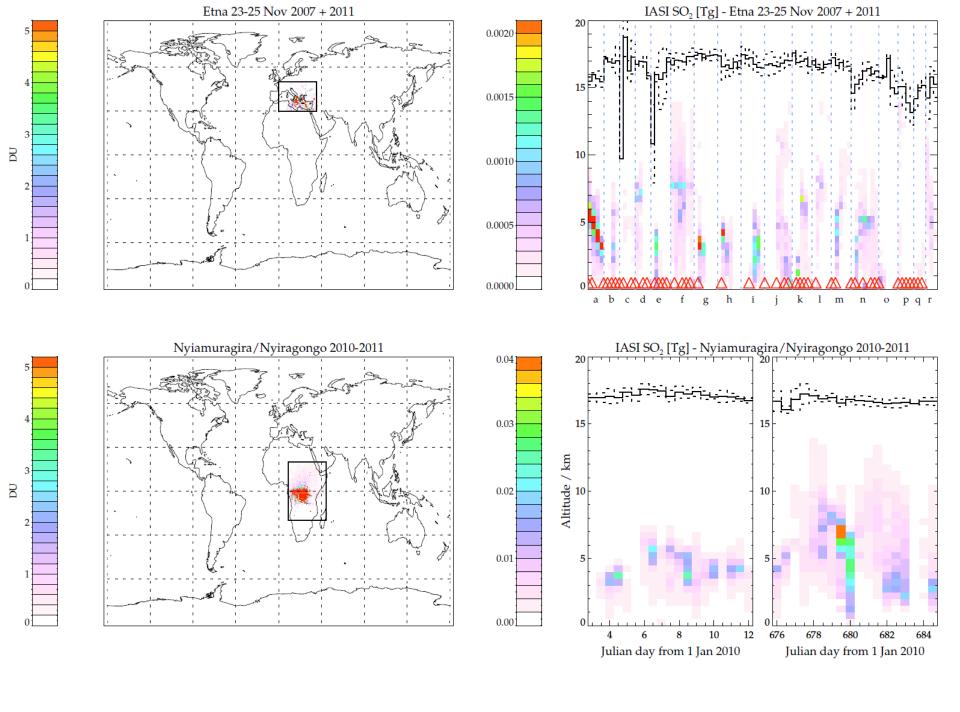
A priori values

SO2, H, s, Ts Xa=[0.5, 400, 100, 290] DXa=[100, 1000, 1, 20]



Sensitivity of retrieval to presence of ash and cloud



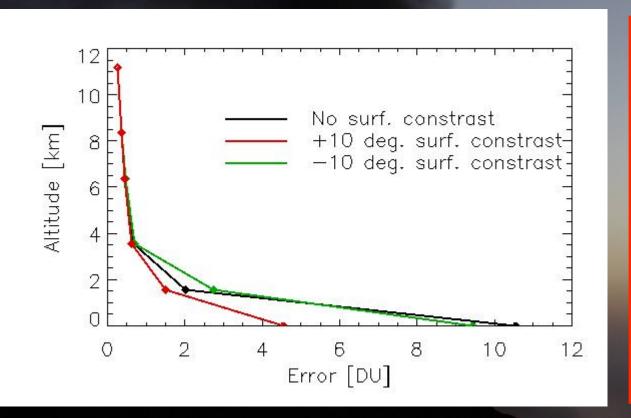


Minimum error estimate

surface contrast = skin temperature - temperature of the first atm. layer std. atm. profiles assumption: we know the altitude of the plume.

 $1 DU = 0.0285 g/m^2$

considering 60 overpass a month
=> error reduced of 1/sqrt(60)



SO₂ monthly errors [km] [g/m2] 11 9 10-4 8.3 13 10-4 6.4 17 10-4 3.5 24 10-4 1.5 73 10-4 388 10-4

SO₂ linear (v3)

It is mainly a 'measurements' of the SO2 signal

- All IASI archive 2007-2014 analysed
- NRT data processing

SO2 vertical profile, atmospheric profiles, Jacobian

Retrieve:

Assume:

SO2 column amount [DU]

Used for: (i) plume detection, (ii) identify where there is a signal

SO₂ iterative (all v1 and v3)

comprehensive error budget for every pixel

Require auxiliary data (ECMWF profiles), radiative transfer (RTTOV) called iteratively

Used for:

Volcanic plume,

study SO2 mass and vertical distribution

Retrieve:

(2)

SO2 column amount [DU] SO2 plume altitude [mb, km] (3) Low signal case: degassing, pollution Results are average in time

(monthly means)

Assume:

SO2 altitude

Retrieve:

SO2 column amount [DU]

(1)

(1) SO₂ linear retrieval (detection) theory

[Rodger 2000]

The optimal estimate of x taking into account total measurement error may be computed as:

$$\hat{\mathbf{x}} = \mathbf{x}_0 + (\mathbf{K}^T \mathbf{S}_y^{\mathsf{tot}^{-1}} \mathbf{K})^{-1} \mathbf{K}^T \mathbf{S}_y^{\mathsf{tot}^{-1}} (\mathbf{y} - \bar{\mathbf{y}})$$

$$\mathbf{G} = (\mathbf{K}^T \mathbf{S}_y^{\mathsf{tot}^{-1}} \mathbf{K})^{-1} \mathbf{K}^T \mathbf{S}_y^{\mathsf{tot}^{-1}}$$

Create a generalized error covariance S_y^{tot} that contains not only the instrument noise, but noises due to interfering gases and broadband scatterers (using IASI spectra only).

 S_y^{tot} is computed considering an appropriate ensemble of N measured spectra to construct an estimate of total measurement error variance-covariance S_y^{obs}

$$\mathbf{S}_y^{\mathsf{tot}} \approx \mathbf{S}_y^{\mathsf{obs}} = \frac{1}{N} \sum_{i=1}^N (\mathbf{y}_i - \bar{\mathbf{y}}) (\mathbf{y}_i - \bar{\mathbf{y}})^T$$

$$\bar{\mathbf{y}} = \frac{1}{N} \sum_{i=1}^{N} \mathbf{y}_i$$

[Walker, Dudhia, Carboni, Atmos. Meas. Tech., 2011]

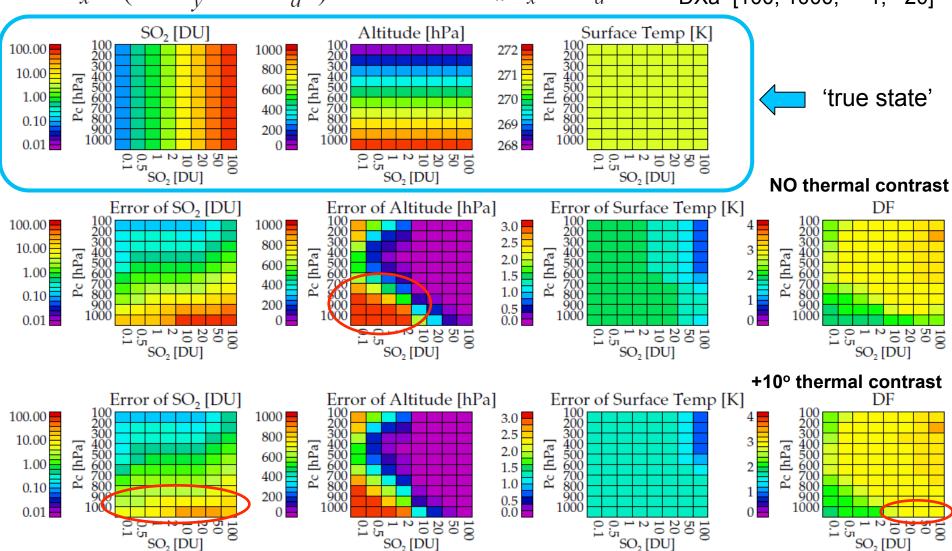
Error analysis

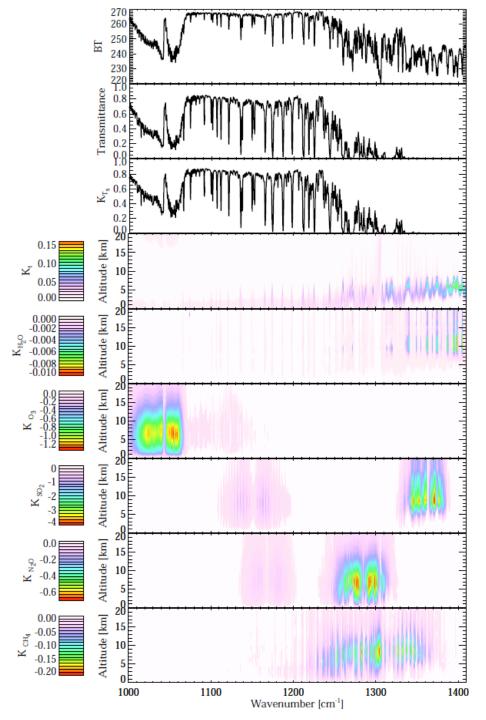
linear error on the 'true' state obtained as:

$$S_x = (K^T S_y^{-1} K + S_a^{-1})^{-1}$$
 $A = S_x S_x^{-1} - S_a^{-1}$

$$A = S_x S_x^{-1} - S_a^{-1}$$

SO2, H, s, Ts Xa=[0.5, 400, 100, 290] DXa=[100, 1000,





Radiative transfer model

Radiative transfer model based on RTTOV (plus SO2 coefficients computed by RAL using the line by line model RFM).

When used in proper retrieval RTTOV is driven by ECMWF profile interpolated at IASI pixel location and time

IASI FORWARD MODEL

The RTTOV output for a clean atmosphere (containing gas but not cloud or aerosol/ash) is combined with an ash layer using the same scheme as for the Oxford-RAL Retrieval of Aerosol and Cloud (ORAC) algorithm.

atm. above

 $I_{ac}^{\dagger} I_{ac}^{\dagger} T_{ac}$

aerosol layer

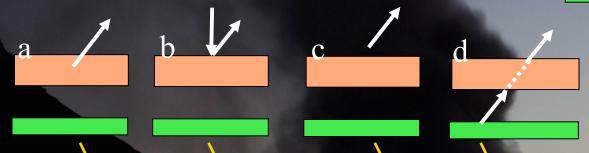
 $R_c,\,e_c,\,T_c^{}$, $B_c^{}$ The suffix 'c' refers to cloud aerosol layer

atm. belowe

 I_{bc} , I_{bc}

surface

 R_s , e_s , B_s

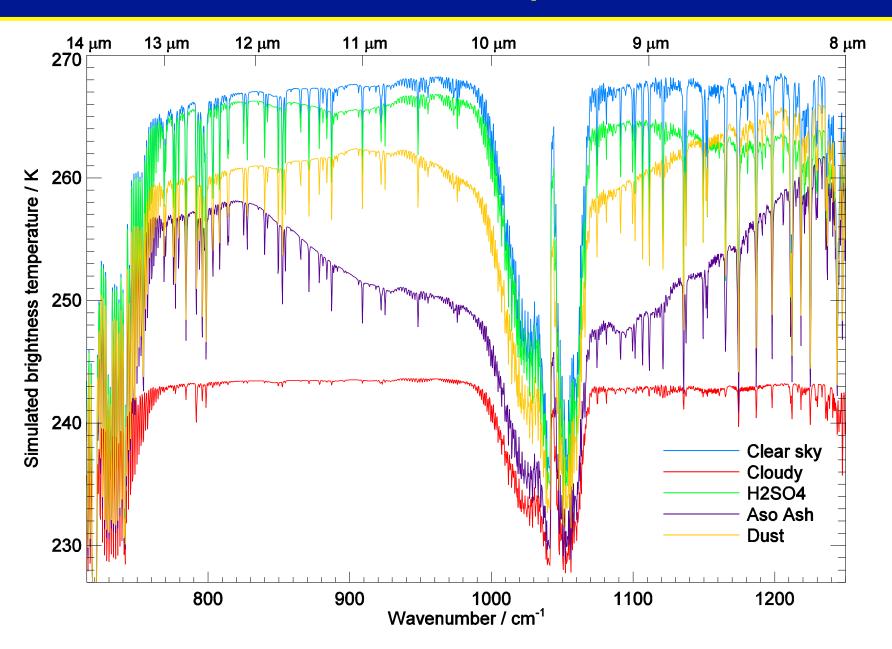


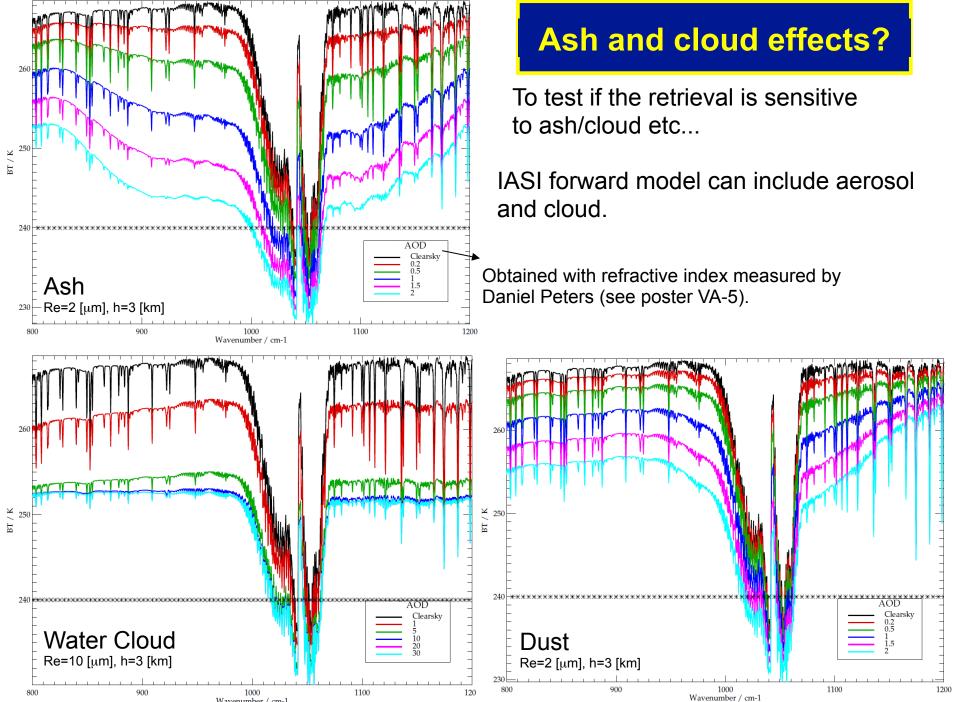
$$I_{\bullet}^{\uparrow} = B(\epsilon_c T_{ac} + I_{ac}^{\downarrow} R_c T_{ac} + I_{ac}^{\uparrow} + B_s \epsilon_s T_b T_c T_{ac} + I_{bc}^{\uparrow(atm)} T_c T_{ac}$$

LUTs for aerosol layer parameters (aerosol optical properties + DISORT)

Other atmospheric parameters (radiances above/below aerosol layer going up/down) are computed with RTTOV using ECMWF atmospheric profiles.

IASI simulated spectra

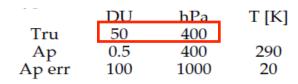


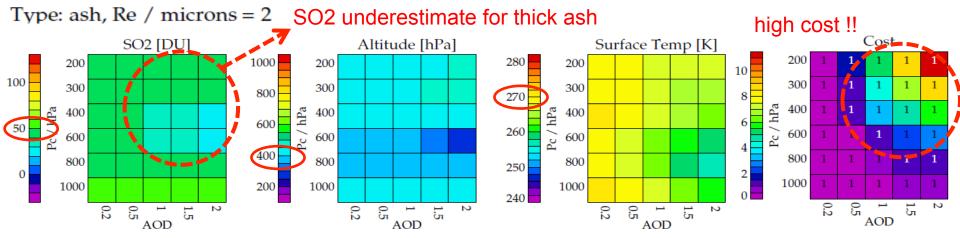


Wavenumber / cm-1

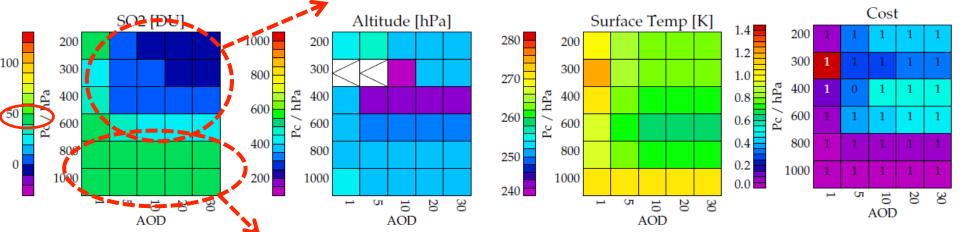
Sensitivity of retrieval to presence of ash and cloud

Retrieval using simulations with ash and water cloud at different optical depth and altitude

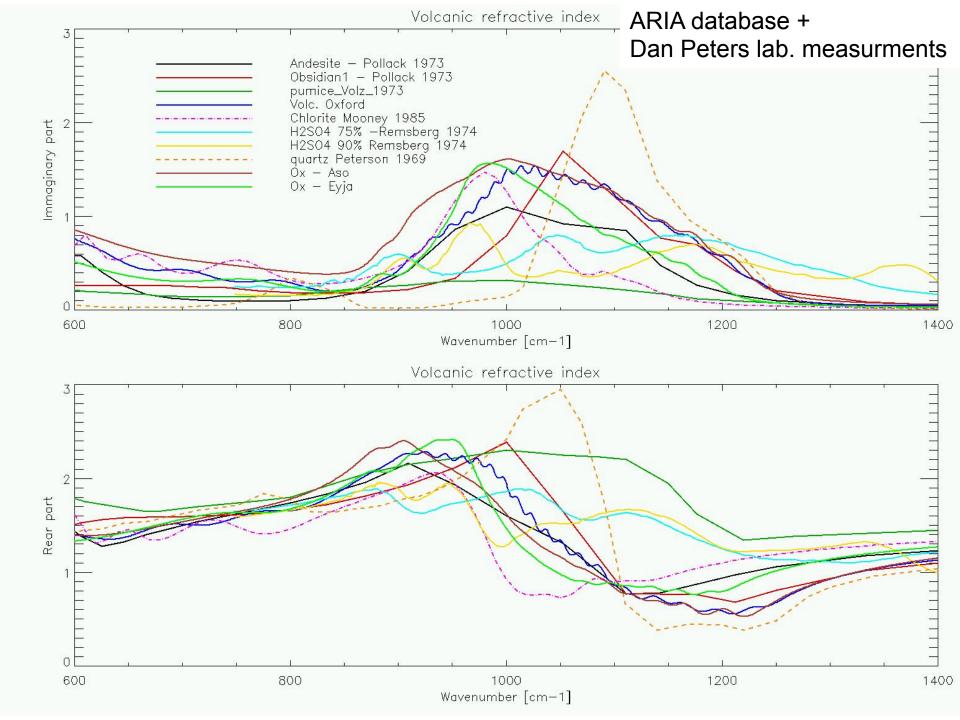




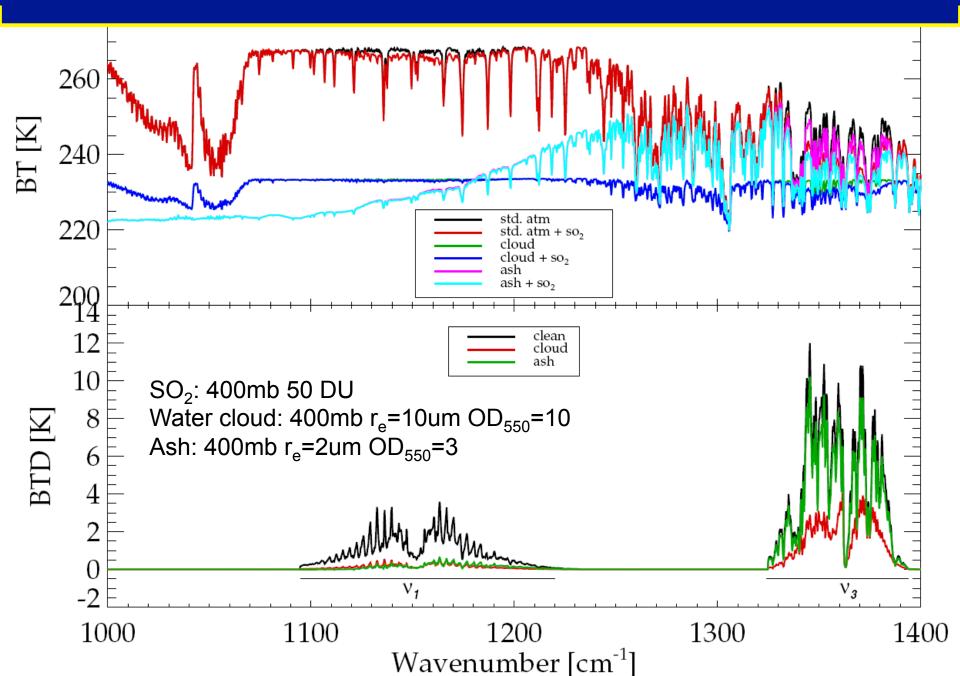
Type: wat, Re / microns = 15 SO2 signal 'covered' from cloud within or above the plume



OK for cloud below the plume!!!



Ash and cloud effects?



Nabro - starts 12 June 2011

- a) The MLS (microwave limb sounder) retrieval of SO₂,
- b) The corresponding MLS temperature profile retrieval,
- c) The IASI SO_2 mass as function of altitude and time (one column every half day). Each column of this plot is obtained summing all the SO_2 amounts (regridded), with retrieved altitude between the indicated vertical levels. In this way we can follow the evolution of the SO_2 plume in the vertical.

As measured from MLS (14th June) and IASI (13th June) the initial Nabro SO2 plume was injected at around 100m, and this altitude correspond to the minimum of the temperature profile as measured by MLS.

MLS data from: http://mls.jpl.nasa.gov/products/so2_product.php

